A METHODOLOGY FOR DETERMINING FLASH FLOOD

FORECASTING SYSTEMS FOR A COMMUNITY

By

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Submitted to the Faculty of the Graduate College of the Oklahoma State University in partial fulfillment of the requirements for the Degree of DOCTOR OF PHILOSOPHY July, 1977





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ACKNOWLEDGMENTS

I wish to sincerely express my gratitude and special respect to Dr. Richard N. DeVries, my major adviser, for his assistance and guidance through my study in pursuit of this degree.

I wish to thank my committee members, Dr. Anthony F. Gaudy, Dr. Don F. Kincannon, Dr. Marcia Headstream and Dr. Douglas Kent for their friendship and assistance and guidance through this course of study.

I wish to thank my loving wife, Jo Ann, for her encouragement and understanding during this period of our twenty-five year love affair.

I want to express my appreciation for the friendship and assistance from my student colleague Tortorelli.

I am most grateful to Mr. John M. Yates, Hydrologist-in-Charge, and fellow employees of the River Forecast Center, National Weather Service, Tulsa, Oklahoma, for their assistance during these years of my course of study.

I sincerely thank Mrs. Juanita Parsons for her patience, cheerful assistance and accurate typing of this dissertation.

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CHAPTER I

INTRODUCTION

Many floods occur in the United States each year, some caused from river flooding which may have a 15-hour to 10-day time-to-peak, while others will be caused by flash flooding which may have a very short peaking time. Flood forecasting requirements for a community, city, state or river basin have been determined in the past by (1) government reaction from a flood, (2) requests from flood prone areas and (3) cooperative programs with governmental agencies. The majority of flood forecast points now in existence in the United States are for the river floods, but the disasterous flash floods of the past six years have brought to issue that forecasting procedures should be available for all floodable communities. The populous movement to rural living and the increase of camping has made people more cognizant of the flood plain areas where they are now or waybe inhabiting. These areas. were always potential flood areas but heretofore were uninhabited. The occurrence in the past few years of heavy rains brings to issue the ever present threat that the Probable Maximum Storm does not exist and that it can occur in any locality. The functional Flash Flood Hydrologist must investigate these newly developing flood plains along with other little known basins to update the current Flood Forecasting Program, and to develop areas of forecasting responsibility.

A flash flood is considered to be a flood that occurs with very little warning and that constitutes an unusual event. In general, flash floods are defined as damaging floods that occur within four to six hours of the time that the causative rainfall occurs. Aside from artifically induced hazards such as dam failures, flash floods are generally the result of relatively intense rainstorms. Precise estimation of the critical rainfall intensity that is capable of producing flash flooding on any specified watershed is difficult, as many watershed and seasonal factors influence the hydrologic response, however, those areas susceptible to rainfall intensities of severe magnitude may be considered to be areas with potential hazards due to flash flooding.

If rainfall intensities of severe magnitude can be considered to be the causative mechanism of flash flooding, then almost all areas of the United States can be considered to be areas of possible flash flooding, as most areas of the United States have experienced rainfall intensities of severe magnitude. Even so, flash floods are most common in the arid and semi-arid regions of the west and southwest. This is due, in part, to the meteorological and physiographic conditions that frequently can lead to the development of large convective thunderstorm cells that are capable of producing large amounts of rainfall in short periods of time.

Formulation of practicable forecasting procedures requires the development and operation of the Flash Flood Forecast Program. The Flash Flood Forecast Program must provide communities in the United States with information concerning the flash-flood potential of streams in the communities and surrounding areas. In addition, when approaching storms or other conditions deem it necessary, flash-flood warnings are issued to appropriate community authorities and media so

that actions to reduce property damage and loss of life may be initiated. Three basic methods of providing flash-flood warnings have been used or proposed by State and governmental agencies. The first approach employs conventional flood forecast techniques at the community level. Under the guidance of the local hydrologic service of a responsible government agency, the community establishes a network of rainfall and river observation stations. As conditions warrant, information concerning rainfall rates, stream stages, and observed storm movement is collected by a warning representative. Flash-flood warnings are issued by the warning representative as necessary. If radar is available in the area, warnings may be based on radar tracking and rainfall measurement, as well as observational reports from the observer network.

The second approach involves use of a recently developed flashflood alarm mechanism. As rising stages in headwater streams or tributaries reach a predetermined height, an alarm located at a continuously occupied public authority is activated by telephone or radio signal. The trigger of the alarm system is an automatic stage measuring device located some distance upstream of the community. The positioning of the triggering device, both geographically and vertically, must be determined from consideration of expected warning times and required evacuation times.

The third approach is dependent upon the skill and alertness of the rainfall forecaster. Warnings of possible flash floods are issued on the basis of rainfall reports received by the meteorologist during the progress of the storms and the meteorologist's estimate of the continuing intensity of the storm. Telemetered rain gages can provide informa-

tion on storm intensities, and radar surveillance can provide information concerning time of onset, duration, areal extent, and intensity.

The Arkansas and Red River Basins in New Mexico, Oklahoma and Texas in that portion west of the 100^o longitude has been an area of little concern to flood forecasting. The general objective of this study is to find a method of determining the flash flood forecasting requirements of a community and then to apply this approach to the subject areas in New Mexico, Oklahoma and Texas. The objective will also be to determine the data input requirement for a forecasting system including equipment.

The specific objective of this research is to find an hydrologic model to predict the flash flood potential of a community based on the hydrological, geological and geographical variables for the Probable Maximum Storm.

This study will analyze the flooding potential of the following communities:

New Mex:	ico	Oklahoma	Texas
Cimarron	Valmora	Beaver	Amarillo
Clayton	Watrous	Guymon	Canadian
Folsom		Logan	Canyon
Raton			Channing
Springer			Palo Duro
Tucumcàri			Tascosa
Ute Park			

This study involves actual on-site experimental field trips into the proposed test communities. These experimental sessions have been made to determine and establish hydrologic parameters.

During this study a way of forecasting any flash flood location has been developed and is presented in the final operational form of tables. These forecast tables are designed to forecast flash floods from varied initial base flows and the table package should allow immediate evaluation of a flash flood threat.

This study also develops a relationship between the community and the possible number of deaths that could occur from any storm. The death relationship is shown in a Flood Potential Scale of one through ten (Table I).

The flash flood potential of a community is shown as a table with ratings of one to 10. The table is called the Flash Flood Potential Scale. This type of presentation of the flood threat is a new approach in the hydrologic field. The Flash Flood Potential Scale is similar to the Richter scale for earthquakes, except that the Flash Flood Potential Scale is shown in unit values of une through 10, rather than the power of 10.

Number one of the Flash Flood Potential Scale is equal to no significant flooding. Number 10 is equal to a catastrophic flood. Numbers one through five of the scale rate the increasing flood potential with no loss of life. Numbers six through 10 of the scale predict the increasing chance of loss of life (Table I).

TABLE I

FLASH FLOOD POTENTIAL SCALE

1	= No Significant Flooding.
2	= Some Street and Low Land Flooding.
3	= Street and Some Residential Flooding.
4	= Major Street and Residential Flooding.
5	= Stream Flooding and Property Loss, Very Light.
6	= Stream Flooding with Chance of Loss of Life and
	Property Loss, Light.
7	= Stream Flooding with Probable Loss of Life (0 to 50)
	and Property Loss, Moderate.
8	= Stream Flooding with Loss of Life (50 to 150) and
	Property Loss, Major.
9	= Stream Flooding with Loss of Life (150 to 400) and
2	Property Loss. Major.
10	= Catastrophic Flood.

Synopsis of Following Chapters

Chapter II contains a review of literature on hydrologic equations, Systems, Models and Multiple Regression Analysis.

Chapter III is concerned with the development of the Flash Flood Model, which includes the application of the MIT Catchment Model, HEC-2 program and Statistical Analysis.

Chapter IV is concerned with the Application of the Flash Flood Model. This Chapler presents the results of the Model in its application to the twenty-five original source basins and the four test basins.

Chapter V discusses the results of Chapter IV.

Chapter VI and VII are the Summary, Conclusions, and Suggestions for future study.

CHAPTER II

LITERATURE REVIEW

Hydrologic Equations, Systems and Model

Leo R. Beard (1) worked on a generalized evaluation of Flash-Flood Potential. The first criterion developed in this study is the flash-flood magnitude index. This index is defined as the ratio of the magnitudes of rare flood events to common flood events, and is indicative of the relative severity of rare flood events. Because of the relatively small variation in the observed skew coefficients for use in annual maximum stream flow frequency analysis, the standard deviation of the logarithms of annual maximum streamflows was considered to be an adequate estimation of the flash flood magnitude The second criterion developed was the flash-flood warning index. time index. This index is an inverse measure of the average warning time available during relatively rare flood events, and is, therefore, a direct measure of the intensity of expected flash-flood magnitudes. The warning time index of a location was defined as the average of the ratio of peak flow to 3-day flow, computed for the top 10% of the observed annual peak flows of the location. It was found that a larger value of flash-flood warning time index indicates less average warning time and higher intensity of flooding than does a smaller value.

Jerald F. McCain and Robert D. Jarrett (2) developed frequency studies for Colorado. Their study contains information of the 5, 10, 50, 100, 500 year floods at sites on natural-flow streams in Colorado. The report used the Log Pearson Type III method for fitting a frequency curve to gaging-station data and multiple-regression techniques for regionalization or transferring the results to ungaged basins. The three basic methods used were:

q

(2.1)

1. Flood Information at Gaged Sites.

2. Flood Information Near Gaged Sites on the Same Stream.

3. Flood Information at Ungaged Sites.

Flood Information at Gaged Sites (2)

$$Q_{T(W)} = \frac{Q_{T(S)} + Q_{T(R)}}{N + E}$$

where

Ε

Q_{T(W)} = the weighted discharge for recurrence interval T, Q_{T(S)} = the station value of the flood for recurrence interval T,

N = the number of years of station data used to compute

= the equivalent years of record for $Q_{T(R)}$.

$$Q_{T}(\nabla) = \frac{A_{U}}{A_{G}} Q_{T}(G)$$

where

QT(V) = peak discharge at ungaged site for recurrence

interval T,

QT(G) = weighted average discharge at gaged site for

recurrence interval T,

Flood Information at Ungaged Sites (2)

$$Y_t = ax_1 \qquad x_2$$

(2.3)

where

Y_t = a flood characteristic, either peak discharge or peak flood depth, for recurrence interval t;

 X_1, X_2 = basin and climatic parameters

a = regression constants; and

b₁, b₂ = regression coefficients

The research was developed from a flood frequency analysis of

(2.2)

gaging-station data and a multiple-regression analysis of flood characteristics and basin and climatic parameter of 258 gaged basins in Colorado and adjacent states. Annual peak discharges through September 1973 were fitted to the Log Pearson Type III distribution. Subjective appraisal of high and low outliners was made based on the reasonableness of the computed flood-frequency curves. High outliners cause frequency curves to estimate extremely large flood discharges, especially for large positive skewness. Low outliners caused large negative skewness but increased the standard deviation of the frequency distribution. Standard multiple regression techniques were used to develop equations by relating flood characteristics at gaged sites to basin and climatic parameters.

The resulting equations developed by McCain and Jarrett for Colorado and New Mexico for the Arkansas and Canadian River basins were:

		0.552 0.460	-	
Q ₅		75A S _B	(2	2.4)
Q ₁₀	=	0.528 0.336 144A S _B	(2	2.5)
Q ₅₀		$891_{A}^{0.482}$ $s_{B}^{0.154}$	(2	2.6)
Q ₁₀₀	=	$1770_{A}^{0.463}$ $s_{B}^{0.086}$	(2	2.7)
Q ₅₀₀	=	5770 _A ^{0.432}	(2	2.8)
Q _{PMS}	=	30,560 ^{0.375}	(2	2.9)
		-0.334		
D ₅	=	22. Ss	(2.	10)
D ₁₀	=	35.5 s ^{-0.462}	(2.	11)

$$D_{50} = 52.1 \text{ s}_{s}^{-0.500}$$
(2.12)

$$D_{100} = 59.3 \text{ s}_{s}^{-0.517}$$
(2.13)

$$D_{500} = 77.3 \text{ s}_{s}^{-0.533}$$
(2.14)

$$D_{PMS} = 156.2 \text{ s}_{s}^{-0.703}$$
(2.15)

where

Q = Peak discharge in cfs; A = Drainage area in square miles; n = Flood frequency; PMS = Probable maximum storm; D = Depth in feet.

V. B. Sauer (3) developed flood-frequency studies for rural and urban areas in Oklahoma and portions of adjacent states. The general form of the equation is:

$$Q_{X(U)} = \frac{7R_{X}Q_{2}(R_{L}-1)}{6} + \frac{Q_{X}(7-R_{L})}{6}$$
 (2.26)

where

 $Q_{X(U)}$ = Urban peak discharge for récurrence interval, X; R_L = Adjustment factor to account for the effect of urban development;

Q_X = natural peak discharge for recurrence interval, X; R_X = rainfall-intensity ratio for recurrence interval X. The resulting equation developed by Sauer for Oklahoma and portions of adjacent states were:

$$Q_{5} = 0.498A^{0.66} s^{0.40} p^{1.58}$$

$$Q_{10} = 1.081A^{0.67} s^{0.42} p^{1.44}$$

$$Q_{50} = 5.40A^{0.69} s^{0.47} p^{1.12}$$

$$Q_{100} = 9.14A^{0.70} s^{0.48} p^{1.01}$$

$$Q_{500} = 19.4A^{1.05} s^{0.58} p^{0.84}$$

$$Q_{2.21}$$

$$Q_{200} = 55.5A^{2.10} s^{0.82} p^{0.54}$$

$$Q_{2.22}$$

where

Q = Peak discharge in Cfs; n = Flood frequency; PMS = Probable maximum storm.

The standard error of prediction for these equations is on the order of ± 40 percent.

Wilbert O. Thomas, Jr. (4) developed techniques for estimating flood depths for Oklahoma streams. The purpose of the report was to present techniques for estimating flood depths for both natural and urban streams in Oklahoma.

The study by W. O. Thomas, Jr. (4) shows that for less than bankfull discharges a basin-wide relation exists between stream depth and discharge when discharge is of equal frequency of occurrence at all sites. It was proposed a general equation of the form:

$$D = CQ^{f}$$

(2.23)

where

D = average cross-section depth

C&F = constants for a given frequency.

Thomas (4) found this type of relation applicable for greater than bankfull discharges in New Jersey, and for simplicity modified the equation to:

$$h = C(Q_{2.33})^{f}$$
 (2.24)

where

h = height of the water surface above the average channel bottom Q_{2} $_{33}$ = mean annual flood discharge

C&f = constants for a given frequency

In this study by Thomas flood depths were determined at 132 gaging stations throughout the State and these were related to basin, climatic, and channel geometry characteristics by multiple regression techniques. The analysis indicated that contributing drainage area, A, and the 2-year 24-hour rainfall are the two most useful variables for estimating flood depths in Oklahoma and portions of adjacent states.

The regression equation used was in the following form,

$$D_{\rm X} = a {\rm A}^{\rm b} {\rm I}^{\rm c} \tag{2.25}$$

where

 D_X = peak flood depth in feet,

- A = contributing drainage area in square miles,
- I = the 2-year 24-hour rainfall in inches,

a = regression constant,

b,c = regression coefficients.

The following equations were defined and indicated by regression analysis:

D ₅	$= 0.53 \text{ A}^{0.24} \text{ I}^{1.60}$	(2.26)
D ₁₀	$= 0.85 \text{ A}^{0.22} \text{ I}^{1.40}$	(2.27)
D ₅₀	= 1.58 A 1.14	(2.28)
D100	= $1.95 A^{0.19} I^{1.06}$	(2.29)
D ₅₀₀	$= 2.85 \text{ A}^{0.17} \text{ 1}^{0.91}$	(2.30)
D _{PMS}	$= 6.25 A^{0.16} I^{0.75}$	(2.31)

where

D = Depth, in feet; n = Flood frequency; PMS = Probable maximum storm.

P. R. Jordan and T. J. Irza (5) developed a technique of determining magnitude and frequency of floods in Kansas. In their report floods were found to be related most significantly to the contributing drainage and the 2-year 24-hour rainfall. The scope of the study was limited to peak flows and does not consider the shape or volume of the flood hydrograph. Equations developed by this study are applicable to unregulated drainage basins in Kansas, ranging in area from 0.4 to 10,000 square miles.

The method of frequency analysis used to determine the N-year flood

peak at gaging stations was the fitting of a "Log Pearson Type III" distribution. This distribution uses three parameters--the mean, standard deviation, and skewness coefficient--calculated from the logarithms of the original data. Adjustments to the Log Pearson Type III frequency were made, where warranted, by one or both of the following conditions:

- Historical data--Information identifying the highest flood during a period longer than the period of gaging station operation.
- Low outliners--One or two exceptionally low-peak flows that have a large effect on the skewness coefficient, and thus effect the computed flood magnitudes for large recurrence intervals.

The final regression equations developed by Jordan and Irza are listed below:

where

Q = Peak discharge, in cfs; n = Flood frequency; PMS = Probable maximum storm.

Q ₅		$3.98 A_{C}^{0.548} P_{2}^{4.752}$	(2.32)
Q ₁₀	=	9.92 $A_{C}^{0.525} P_{2}^{3.591}$	(2.33)
Q ₅₀ /	=	47.6 $A_{C}^{0.523}$ $P_{2}^{2.821}$	(2.34)
Q ₁₀₀	#	83.8 $A_{C}^{0.524} P_{2}^{2.529}$	(2.35)
Q ₅₀₀	ŭ	202. $A_{C}^{0.524} P_{2}^{1.98}$	(2.36)
Q _{PMS}	n	1101. $A_{C}^{0.523} P_{2}^{0.56}$	(2.37)

16.

David M. Hershfield (6) in 1961 developed the Rainfall Frequency Atlas of the United States, referred to as Technical Paper 40. Until about 1953, economic and engineering design requiring rainfall frequency data was based largely on Yarnell's Paper (7) which contains a series of generalized maps for several combinations of durations and return periods. Yarnell's maps are based on data from about 200 first order Weather Service stations.

The data for Technical Paper 40 was divided into three categories. First, there was the recording-gage data from long-record first-order Weather Service stations; second the recording-gage data of the hydrologic network which are hourly reports, and finally, the very large amount of non-recording gage data with observations made once a day.

The factors considered in the construction of the isopluvial maps were availability of data, reliability of the return period estimates and the range of duration. There was much data available for construction of the 2-year 24-hour maps (Figure 1) and these are deemed most significant. The probable maximum precipitation (PMP) relationship uses a combination of physical model and several estimated meteorological parameters. The main purpose of the PMP method is to provide complete safety design criteria in cases where structure failure would be disastrous. The PMP relationship is based on a 10-square mile value of Hydrometeorological Report No. 33 (8) (Figure 2).

2-YEAR 24-HOUR RAINFALL (INCHES)



Figure 1. 2-Year 24-Hour Map, Rainfall Frequency Atlas of the United States, Technical Paper 40 (6)

PROBABLE MAXIMUM 6-HOUR

PRECIPITATION



Figure 2.

 Probable Maximum Precipitation, Rainfall Frequency Atlas of the United States, Technical Paper 40 (6) 19 .

Digital Computer Models

Weather Service River Forecast System

The National Weather Service River Forecast System is called NWS HYDRO 14 (9). The program was developed by the National Weather Service (NWS) and the basic portion of the system is that it was developed around the hydrologic cycle. The Model has a primary requirement input from a gaged location and, therefore, could not be easily allocated to an ungaged area.

Sacramento Model

The Sacramento Model was developed from the Stanford Model (10) and the NWS Hydro 14 Model. The Sacramento Model attempts to simulate streamflow by simulating all of the significant components of the hydrologic cycle in a more simplified manner. Burnash has tried to associate each variable in the Model with a recognizable counterpart in the physical world. This Model is one more developed to fit the larger, historic basins rather than the ungaged small basin where most flash flooding occurs.

MIT Catchment Model

The MIT Catchment Model (11) represents movement of water over the catchment surface and through the network (Figure 3). model recognizes that surface geometry is extremely irregular and impossible to represent in complete detail in either a physical or a



Figure 3. Equivalent Block Diagram of the Natural Catchment, MIT Catchment Model



Figure 4. Typical Natural Catchment MIT Catchment Model

mathematical model. A reductionist effect was used to replace the natural complexities with a number of simple elements such as overland flow planes, stream segments, pipe lengths, etc. A suitable combination of an appropriate number of these simple elements is assumed sufficient to model the behavior of an entire catchment.

Illustrated in Figure 4 is an example catchment. A possible combination of overland flow segments and stream flow segments would appear as a detailed model of this catchment. Consider the drawing of a simple catchment stream element as pictured in Figure 5. Rain falls on the overland flow surfaces of this simple catchment. If the surface is pervious, then the rain will soak (infiltrate) into the ground and there will be no runoff. If the intensity of the rainfall exceeds the infiltration rate, then there will be rain excess, and runoff toward the stream channel will occur. The water runs off in the form of sheet flow or overland flow, and the motion of this water is described mathematically by means of the kinematic wave approximation. It is noteworthy to realize that the kinematic wave theory was originally intended to be used to describe flow in channels, but with respect to overland flow; the water is considered to flow in very wide channels one unit of length (i.e., 1 foot) in width $(y/b \ll 1 \ y = depth, b = width).$

Some of the basic considerations governing the development phase of the MIT catchment model were:

- 1. That the model would be based on sound physical reasoning.
- 2. That the parameters were to be directly related to the physical characteristics of the catchment and insofar as possible to be directly measurable from map or field data.



Figure 5. Simple Conceptual Catchment

- 3. That the model would be subject to experimental verification.
- 4. That the model would be based on sound hydrologic parameters.
- 5. The model would be as simple as possible.
- 6. That the time period could be from a few minutes to several hours.
- 7. That historic rainfall or runoff data would be eliminated or minimized.
- 8. That the model would handle any size catchment.
- 9. That flow distributed over the surface of the catchment would be modeled as planes of overland flow.
- 10. That the flow from the overland planes would be collected by streamflow segments as lateral inflow and then passed downstream to other stream segments.

The MIT Catchment Model defines the following parameters:

<u>Overland Flow</u>. In the direct runoff of a vegetated catchment, the water trickles over, through, and under a highly irregular surface. The flow regime varies between laminar and turbulent, and the flow among the stems and debris of dense vegetation may at times resemble flow through porous media. It is clear, that a vigorous mathematical description of this phenomenon would require solution of the continuity and momentum equations at an exceedingly small scale, both areally and temporally. Although computer techniques for solving these equations on single elements of homogeneous surface are well established, the amount of field data needed to implement and verify the extension of these methods to the detailed variations of natural surfaces, the computer size, and the computational expense necessary for the solution of the resulting vast systems of equations are inconsistent with the insensitivity of catchment response to this microscale of variability.

The important characteristic of overland flow is that the water is distributed over a wide area at a very small average depth until it reaches a well-defined stream channel.

<u>Streamflow</u>. For modelling of the stream segments many routing techniques are available, ranging from solutions of the full nonlinear continuity and momentum equations through progressively simplified or linearized forms of these equations to simple parametric storage models. It would be desirable to use a form of these equations which is compatible with the overland flow model requirements and which would represent those nonlinearities important to the dynamic behavior of the catchment.

Lighthill and Whitham (12) in their comprehensive considerations of the fluid mechanics of flood movement in rivers, have separated the effects into dynamic and kinematic waves, both of which are initially present. They show that for Froude numbers less than 2, the dynamic component decays exponentially and the kinematic wave ultimately predominates. Woolhiser and Liggett (13) indicate that the rate of damping of the dynamic component will be large enough to justify neglecting the dynamic effects provided that
$$k = \frac{S_0L}{yF^2} > 10$$

where

 $S_0 = Slope of the stream;$

q = a y

L = Length;

Y = Depth of flow;

F = Froude number.

Use of the kinematic form of the unsteady flow equations allows particularly simple numerical solutions (since all disturbances propogate only in the downstream direction), while retaining some of the non-linear effects of the full dynamic form. The successful application by Wooding (14) of this approach to natural catchments ranging from 0.84 square miles to 3383 square miles has led to the adoption of the kinematic approach as the basic routing element.

The Kinematic Wave Equation. The kinematic wave equation for an overland flow segment is

$$\frac{\partial y}{\partial t} + \frac{\partial q}{\partial x} = (i - f) / 43200 \qquad (2.39)$$

(2.40)

(2.38)

where

y = the depth of flow (ft);

q - the rate of flow (cfs/ft);

t = time (sec);

x = distance along the segment (ft);

i = the rainfall intensity (in/hr);

f = the infiltration rate (in/hr).

In equation (2.39) both i and f may vary with x and t. The difference i - f may be treated as an effective rainfall rate (which by convention in hydrology is never negative), or the water remaining on the surface when f exceeds i may be permitted to continue to percolate into the soil. The fact that f may vary with x causes the model to simulate runoff only from those locations where i exceeds f.

The corresponding equation for the stream segments is

$$\frac{\partial A}{\partial t} + \frac{\partial Q}{\partial x} = q$$
 (2.41)

(2.42)

where

A = Cross-sectional area of flow (ft^2) ;

Q = Discharge rate (cfs);

q = Lateral inflow rate of overland flow (cfs/ft).

The above kinematic wave equations contain the parameters a_c , m_c , a_s and m_s which may be estimated from the Manning formula

$$q = \frac{1.49}{n_c} y_c^{5/3} s_c^{1/2}$$
(2.43)

as

$$a_{c} = \frac{1.49}{n_{c}} S_{c}^{1/2}$$
(2.44)

$$m_c = 5/3$$
 (2.45)

in the case of overland flow or from the Manning formula

$$Q = \frac{1.182}{n_{g}} \left(\frac{\sqrt{z}}{1 + \sqrt{1 + z^{2}}} \right)^{2/3} A^{4/3} S_{g}^{1/2}$$
(2.46)

as

$$a_{s} = \frac{1.182}{n_{s}} \left(\frac{\sqrt{z}}{1+\sqrt{1+z^{2}}}\right)^{2/3} S_{s}^{1/2}$$
(2.47)
$$m_{s} = 4/3$$
(2.48)

in the case of flow in a triangular channel.

The MIT catchment model solves the kinematic wave equations by numerical techniques. The details of these techniques have been carefully developed over a period of many years to the point where reliable procedures have been programmed to automatically assure the most economical solution of these equations. These numerical procedures discretize time in steps of Δt and distance in steps of Δx . Over any one time step, i and f are assumed to remain constant. Variables i and f also are assumed to remain constant over x. Variations with x of f can be represented in stepwise changes by a cascade of overland flow segments. Along any stream segment, the variable q changes continuously with time but is assumed constant over the length of the stream segment.

<u>Time of Concentration</u>. One of the most important considerations to reduce a catchment to a network of segments is the time it would take each segment to reach equilibrium when excited by a steady-state inflow. This time is called the time of concentration and may be derived from the kinematic wave equation. The time of concentration for an overland flow segment is

$$t_{c} = \left(\frac{L_{c}}{a_{c}(i_{e}/43200)}\right)^{-1/m}c$$
 (2.49)

where i is the rate of rainfall excess (in/hr). The time of concentration for a streamflow segment is

$$E_{s} = \left(\frac{L_{s}}{a_{s} q}\right)^{1/m} s$$
 (2.50)

where q is the equilibrium rate of lateral inflow to the stream from all adjacent overland flow segments and there is assumed to be no upstream inflow to the stream segment. The time of concentration

for a simple model having identical overland flow segments on each side of a stream segment would be

$$t_{l} = t_{c} + t_{s}$$

$$= \left(\frac{L_{c}}{a_{c}(i_{e}/43200)}\right)^{1/m}c + \left(\frac{L_{s}}{a_{s}(2i_{e}L_{c}/43200)}\right)^{1/m}s$$
(2.52)

Although the time of concentration is important to reduce the catchment to a minimum number of segments, it should be noted that natural rainfall events do not occur at constant intensity, and the time of concentration cannot be observed in the field.

Infiltration. Although infiltration research has been conducted for many years, there still is no generally recognized adequate quantitative model of natural infiltration. Considering the complex combination of soil characteristics, soil moisture conditions and other factors occurring in nature, this is not too surprising. Nevertheless, infiltration is important because it can influence not only the volume and intensity of direct runoff rates but the timing of the runoff hydrograph as well. The MIT catchment model has the capability of utilizing two different methods for determining infiltration, the Soil Conservation Service method (15) and Horton's method (16). Horton's method is a mathematical equation for defining the rate curve of infiltration capacity:



where

- $f_0 =$ initial rate of infiltration capacity (in/hr)
- e = base of natural logarithms
- k = constant depending primarily upon soils and vegetation \min^{-1}
- t = time from start of rainfall
- f = steady state infiltration capacity (in/hr)

See Figure 6.

The values of f_0 , f_c and k are given by the model user as data input and are placed on data card number 4. All previous areas are presumed to have the same infiltration properties. However, an overland flow segment does not have to be completely pervious. The percentage of imperviousness for each overland flow segment is specified in the data input and the model user can have a 100% impervious segment, a 100% pervious segment or any combination thereof.

In computing excess precipitation for any overland flow segment, the following equation is used:

EP = (RAIN)(IMP) + (EXCESS)(1-IMP)(2.54)

EP	excess precip	itation			
RAIN	amount of rain	n falling in s	segment	:	н
IMP	percentage of	overland flow	v segment	which is	impervious
(1-IMP)	percentage of	overland flow	segment	which is	impervious

33

(2.53)

EXCESS = present rate of precipitation excess for pervious parts

of the segment

The second method for determining infiltration is through the use of the Soil Conservation Service Method. The method is explained fully in the SCS National Engineering Handbook, Section 4, "Hydrology", United States Department of Agriculture, 1972 (15) but it will be briefly explained here, since it is the method used in this example.

The SCS has divided all the soils in the United States (including Puerto Rico) into four major hydrologic soil groups. Group A (low runoff potential) are soils having high infiltration rates even when thoroughly wetted and consist chiefly of deep, well - to excessively drained sands or gravels. These soils have a high rate of water transmission. Group B are soils having moderate infiltration rates even when thoroughly wetted and consist chiefly of moderately deep to deep, moderately well to well-drained soils with moderately fine to moderately coarse textures. These soils have a moderate rate of water transmission. • Group C are soils having slow infiltration rates when thoroughly wetted and consists chiefly of soils with a layer that impedes downward movement of water, or soils with moderately fine to fine texture. These soils have a slow rate of water transmission. Group D (high runoff potential) are soils having slow infiltration rates when thoroughly wetted and consists chiefly of clay soils with a high swelling potential, soils with a permanent high water table, soils with a clay pan or clay layer at or near the surface, and shallow soils over nearly impervious material. These soils have a very slow rate of water transmission.

These four major groups are then broken down further by land use and treatment, by hydrologic soil-cover and by antecedent moisture con-The land use and treatment class incorporates items such as ditions. crop rotation, contouring, terracing, pasture, range, farm woodlots, commercial forests, straight-row farming, roads and urban areas. The second class, that of hydrologic soil-cover, is self-explanatory. The third class is the antecedent moisture conditions. This class is divided into three areas. A soil can be AMC-I, a condition of watershed soils where the soils are dry but not to the wilting point, and when satisfactory plowing or cultivation takes place. AMC-II is the average cause for annual floods, that is, an average of the conditions which have preceded the occurrence of the maximum annual flood on numerous watersheds. AMC-III is the condition of the soil when heavy rainfall or light rainfall and low temperatures have occurred during the five days previous to the given storm and the soil is nearly saturated. In this example AMC-II was considered in all cases.

Thus, the major soil type, land use, and ground cover were determined for each overland flow segment of the catchment. Assuming that we use AMC-II, the SCS National Engineering Handbook was consulted and from this manual, a certain curve was found. The curve number is an arbitrary number used with the SCS method. Multiple Regression Analysis (17)(18)

If there are m variables to correlate, including one dependent and m-1 externally independent, the general equation for multiple linear regression is

$$Y = B_0 + B_1 X_1 + \dots + B_i X_i \dots + B_m X_m$$
 (2.55)

where B_0 is the intercept and B_i is the multiple regression coefficient of the dependent variable Y on the independent variable X with all other variables kept constant, m = number of variables.

The principal results for the multiple regression model (equation (2.55) can be shown in matrix form.

To express the multiple linear regression model

$$Y_{i} = B_{0} + B_{1}X_{11} + B_{2}X_{i2} + B_{m}X_{im} + \epsilon_{i}$$
 (2.56)

In the general form, the multiple regression model (4.3) is then

$$Y = XB + \varepsilon$$

nxl nx(m+1)(m+1)xl nxl (2.57)

where

Y is a vector of observations

B is a vector of parameters

X is a matrix of constants

 ε is a vector of independent normal random variables with expectation $E(\varepsilon) = 0$ and

Variance-covariance matrix $\sigma^2(\epsilon) = \sigma^2 I$.

The least squares normal equations for the general multiple linear regression (4.4) are:

$$(X'X)' = B + X' Y$$
 (2.58)
(m+1)x(m+1) (m+1)x1 (m+1)n nx1

and the least squares estimators are

$$b = ('X')^{-1} X'Y$$
(2.59)
(m+1)x1 (m+1) (m+1) m+1x1

Let the vector of the fitted values \hat{Y}_1 be denoted by \hat{Y} and the vector of the residual terms, $e_1 = Y_1$.

The fitted values are represented by

$$\hat{\mathbf{Y}} = \mathbf{X}\mathbf{b} \tag{2.60}$$

and the residual vector by

$$e = Y - \hat{Y}$$
(2.61)

The sums of squares for the analysis of variances are: Sums of squares total = SSTOT = $Y'Y - nY^{-2}$ (2.62) Sums of squares regression = SSR = $b'X'Y' - n\overline{\overline{Y}}^2$

Sums of squares error =
$$SSE = e'e = Y'Y - b'X'Y'$$
 (2.64)

(2.63)

The sum of squares total, as usual, has n-1 degrees of freedom associated with it. The sum of squares error has n-(m+1) degrees of freedom associated with it since m+1 parameters need to be estimated in the regression function for model (4.4). Finally, the sum of squares regression has m+1-1 = m degrees of freedom associated with it, representing the number of X variables $X_1 \dots X_m$ for which a coefficient has been estimated.

TableII shows these analyses of variance results, as well as the mean squares MSR and MSE:

(MEAN SQUARE REGRESSION) =
$$MSR = \frac{sum of square regression}{m}$$
 (2.65)

(MEAN SQUARE ERROR) = MSE =
$$\frac{\text{sum of square error}}{n-m+1}$$
 (2.66)

The expectation of MSE is σ^2 , as for simple regression. Street and Torrie (18) + stated that the expectation of MSR is σ^2 plus a quantity which is positive if any of the B_k (k = 1,, m) coefficients is not zero. For instance, when m+1-1 = m = 2, then

$$E(MSR) = \sigma^{2} + B_{1}^{2} \Sigma(X_{11} + \bar{X}_{1})^{2} + B_{2}^{2} \Sigma(X_{12} - \bar{X}_{2})^{2} + 2B_{1}B_{2} \Sigma(X_{11} - \bar{X}_{1})(X_{12} - \bar{X}_{2}) /2$$
(2.67)

Thus, if both B_1 and B_2 equal zero, $E(MSR) = \sigma^2$. Otherwise, $E(MSR) > \sigma^2$.

TABLE II

	· · · · · · · · · · · · · · · · · · ·	•	
Source of			
Variation	Sum of Square	DF	
REGRESSION	$SSR = b'X'Y' - n\overline{Y}^2$	m.	$MSR = \frac{SSR}{m}$
ERROR	SSE = Y'Y - b'X'Y'	n-m-l	$MSE = \frac{SSE}{n-m+1}$
TOTAL	SSTO = $Y'Y-nY^2$		

ANOVA TABLE FOR GENERAL LINEAR REGRESSION MODEL

The coefficient of multiple determination, denoted by \mathbb{R}^2 , is defined as follows:

 $R^2 = \frac{SSR}{SSTO} = 1 - \frac{SSE}{SSTO}$ (2.68)

It measures the proportionate reduction of total sum of squares variation in Y associated with the use of the set of X variables X_1, \ldots, X_m . The coefficient of multiple determination R^2 reduces to the coefficient of simple determination r^2 (simple regression) when m = 1; that is, when one independent variable is in the model (equation (2.68). Thus, for R^2 we have

$$0 \leq R^2 \leq 1.$$
 (2.69)

 R^2 assumes the value of 0 when all $b_k = 0$ (k = 1,, m). R^2 takes on the value 1 when all observations fall directly on the fitted

response surface; that is, when $Y_i = \hat{Y}_j$ for all i.

The coefficient of multiple correlation R is the positive square root of R^2 :

 $R = \sqrt{R^2}$.

(2.70)

CHAPTER III

DEVELOPMENT OF THE FLASH FLOOD MODEL

Flash Flood Model

To be able to walk across a drainage basin and from visual observation predict the flows and stage for a varied condition of runoff has been the dream of many hydrologists. The Flash Flood Model is an attempt to enable an hydrologist to be able to predict the maximum disaster and major floods that may occur to a community or to a basin by knowing a few basic basin parameters (Figure 7).

Today the urban and rural residents of our country are warned not to live in a flood plain. Many State and Federal agencies have stressed the point that the only real protection is to locate one's home or business above the 100 year flood plain. These governing bodies do not attempt to explain the possibility that a flood five times worse than the 100 year flood could occur. The development of this Flash Flood Model is an attempt to educate the citizenry and governing bodies that it is possible that a catastrophic flood could occur in their respective communities.

The specific purpose of the Flash Flood Model is to provide to the hydrologist/engineer a tool to aid him to better analyze a basin. Consider that the Probable Maximum Storm only occurs twice a year in the United States, therefore, where it does occur, the departure from normal hydrologic conditions is usually so great that any understanding of what could happen is beyond comprehension.

The Flash Flood Model first establishes within its files the average rainfall/runoff values. The index to moisture conditions within a basin is the universal application of the antecedent precipitation index P_a . The equation is:

$$P_a = b_1 P_1 + b_2 P_2 + B_3 P_3 + \dots + b_t P_t$$
(3.1)

where

In a day to day accounting of the index, there is advantage in assuming that b_t decreases with t according to a logarithmic recession,

Therefore

$$Pat = Paok$$

rather than as a reciprocal.

(3.2)

where

$$b_t = k^t$$

a = Coefficient

k = Exponent

o = Initial value of API.

then by letting t equal unit,

 $P_{a1} = kP_{a0}$ (3.3)

Therefore, the index for any day is equal to a constant k times the index of the day before. The value of k is a function of the geographic area. Experience has shown that the k value in the plains area of the United States is 0.90.

The Flash Flood Model is one main computer model with two subroutines. The sub-routines are the MIT Catchment sub-routine and the HEC-II sub-routine (Figure 7). The Flash Flood Model has within its program four basic data arrays, namely, DATA TSA, DATA TSB, DATA TSC and DATA TSD (Figure 13). These data sets relate to soil moisture conditions in areas of four different rumoff capabilities. DATA TSA is for 0.25 inches of runoff to flood an average stream in a particular area, DATA TSB is for 0.50 inches, DATA TSC is for 0.75 inches and DATA TSD is for 1.00 inch.

The four data sets were compiled from the rainfall/runoff curve relationship in Figure 13. These tables were keyed on the stream capacities of an area and the amount of rainfall required to cause stream flooding.



Figure 7. Flow Chart of the Flash Flood Model





The next major component of the model is the data requirement as shown in Figures 9 and 10.

The model will next interrogate the statistical equations to determine the positioning of the subject basin within the population of statistical data. The model will use the statistical equation developed from the "Best" 6 variables found by the maximum R-square improvement procedure. (A description of the statistical model is located further on in this chapter). In the "Best" procedure, the variables are RFPMS (rainfall/probable maximum storm), DENFLP (population density of flood plain), DRAINA (drainage area), RAINFAL (rainfall), SLOPCN (slope of channel), SLOPBAS (slope of basin).

The main equation used in this subroutine is:

DEATHS = -71.641 + 2.53RFPMS + 0.016DENFLP - 0.101DRAINA - 2.328RAINFL + 2.414SLOPCH + 0.728SLOPBA (3.4)

Next in line of computation is the development of a stage dis-"charge relationship called a rating. This rating employs the use of empirical formulas (2)(4)(5). The option will exist to develop the rating by the noted empirical formula or by other analytical means through use of the Corps of Engineers Hydrologic Engineer Center's Backwater Program called "HEC-II" (19).

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Figure 9. Data Requirements for Flash Flood Model

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Figure 10, Data Req Sheet 2.

Data Requirements for Flash Flood Model

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The majority of these empirical equations will be in the form:

$$Q_n = (Constant) (Drainage Area) (M) (Slope of Basin) (L) (3.5)$$

where

Q = Flow in CFS

n = Flood frequency

M = Exponent for drainage area

L = Exponent for slope of basin

or

Stage = (Constant) (Drainage Area) (M) (2 hour storm)+

Constant

where

Stage = Gage height in feet

n = Flood frequency

M = Exponent for drainage area

The model will next position the community within the Flash Flood Potential Scale. The program will determine the potential death number and then locate its number position in the scale between one to ten.

The model will next calculate hydrologic parameters from basic information supplied as noted in Table V. It is in this parameter calculation step of the model that the time to peak, unit graph peak, channel slope, overbank slope, 100 year flow, 100 year gage height, probable maximum storm flow depth, number of rainfall gages required and number of river stations required are determined.

(3.6)

The time of concentration is determined by the equation: (20)

$$T_{c} = \frac{((11.9)(RM)^{3.0})}{\Lambda h^{0.35}}$$
(3.7)

where

- $T_c = Time of concentration;$
- RM = Distance of stream, in miles;
- Δh = Difference in height of basin.

The time to peak is determined by the equation: (20)

$$T_{p} = ((Dur/2.0) + (0.6)(T_{o}))(.7)$$
 (3.8)

where

T = Time to peak; p = Time to peak; Dur = Rainfall duration in yours; T = Time of concentration.

The unit graph peak is determined by the equation: (21)

UG= ((DA) (640)) ((2/(
$$T_p+T_r$$
))

where

UG = Unit graph peak, in cfs; DA = Drainage Area, in miles; T_p = Time to peak; T_r = Time of recession. (3.9)

The model will next write all output on the disk, waiting for a later command to print results.

The model next investigates the input data for dam break information. If dam data is available, then flow from a total wash out break and a breach break will be determined in terms of flow and downstream stage.

Peak outflow from a complete reservoir washout is calculated by: (22)

$$Q_{\text{Res}} = (8/27)(B)(32.2)^{0.5}(DEPTR)^{1.5}$$
 (3.10)

where

DEPTR = Depth of Reservoir

The MIT Catchment Model (23)

The basic equations of the MIT Catchment Model were described in Chapter II. The MIT Catchment Model is used as a subroutine to the Flash Flood Model. It is used to calculate the time to peak and peak flow from a predetermined study storm, or operationally for unusual storm events.

The data sheet for compiling the data input for the MIT Catchment Model is shown in Figures 11 and 12.

The computer cards for the Flash Flood Model is listed in Figure 13.

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Figure

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Requirements

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Form #24-7327-2 Palaned in U.S.A. FORTRAN CODING FORM Punching Instructions Poge 1 MIT MODEL - DETERMINISTIC URBAN RUNOFF MODEL 01 2 Program Cord Form # Identification Grophic Ly ... Programmer Dote ي. ور Punch ---- C TOR COMMENT STATEMENT E FORTRAN STATEMENT 10 15 20 35 40 45 50 55 7980 5 6 7 25 30 60 65 ٠. PART 1 (TITLE, ONE CARD T- 2014) 1 in Drainage area/(avg. over-SWALE FLOW (Use width of channel one foot above bottom) ς ٢ S: C S INFL Thiessen coel 1.9 PART 3 (PASIN INFORMATION -- BAU, 2F1, 16, F8.9, 45.9, 5F2.0) L. Impervicus Ft9Pe Manning (1. paved) ın. O. pervi) . (INFILTRATION PASSANETERS 4F5.0) ***** ONLY TO PE USED WITH HORTON METHOD FART L Initial Initial Strady Event loss Infil R; Infil R; Fale depay

* A standard card form, IBM electro 888157, is available for punching source statuments from Uis form.

J. 23

5)ອ 2 N 2 of 2 Identification \$ Poge 8 3 Card Form # 2 Punching Instructions FORTRAN CODING FORM 2 FORTRAN STATEMENT · A standard card form, IBM electro 888157, is available for punching source statements from this form. Greenic Punch 8 ĩη .Sta. • WIT WODEL - DETERMINISTIC URBAN RUNOFF MODEL Dole 2 Sta. 4 . 8 PART 5 [1]2 DE BAENFALL DATA CARDS ---No. of [2]ar to data od]eratnebe 8 Sta. 3 = LOCATION Rainfall Lintenuj vy Linches/hr Sta. 2 (Sta. 2) • 2 --, C FOF COMMENT • 514141417 2 BW Programmer ttm. sunce start store rogram

Figure 12. Data Requirements for MIT Model (Sheet 2)

Prime Karlanda

Computer Program

Flash Flood Model

С

PROGRAM TABL REAL MLCITY DIMENSION DNAME(13) . CNAME(13) DIMENSION TSA(209), TSB(209), TSC(209), TSD(209), TSE(88) DIMENSION ZB(231) + TS(924) + ZEELV(2310) + ZA(2310) DIMENSION RRFF(11), FFFF(21), DATE(2), STATE(3) COMMON STAGE(7) . EV(5) . Q(7) . QX(10) . DATUM COMMON BASIN(13), REACH(13) COMMON RF(11) + FF(21) + EELV(21+11) DATA TSA/ .07. .81. 1.21. 1.66. 2.64. 3.71. 4.80. 6.98. 8.90. .521 1 .28. .03. .69. 1.06. 1.48. 2.42. 3.45. 4.55. 6.75. 8.66. 1 .21, .43. .92, 1.30, 2.21, 3.22, 4.30, 6.52, 8.41, .00, .15. .34 . .57. 1 .28. .49. .81, 1.17, 2.05, 3.05, 4.11, 6.35, 8.23, .00, .12, 1 .37, .00. .92, 1.71, 2.67, 3.72, 5.96, 7.81, .05. .19. .63. 1 .00. .291 .52. .80, 1.48, 2.40, 3.43, 5.68, 7.50, 1 .01, .13. .221 .67, 1.39, 2.29, 3.30, 5,49, 7,31, .58, 1.30, 2.25, 3.28, 5.40, 7,27, .49, 1.08, 1.82, 2.70, 4.95, 6.70, .00, .07. .40. .00. 1 1 .00. .00. .00. .11. .33, 1 .00. .00. .00. .10. .28. .00. .00. .01. .38. .00. .18. .88, 1.58, 2.42, 4.56, 6.28, 1 .76, 1.42, 2.23, 4.31, 5.96, 1 .00. .00. .00. .00. .10. .29. .00, .00. .00'. .00. .04. 1 .00. .00. .00, .00. .56, 1.10, 1.84, 3.78, 5.35, .00, .15. 1 .00. .00. .00. .00. .08. .47. 1.06. 1.78. 3.77. 5.40. 1 .00. .94, 1.63, 3.58, 5.19, .38. .00. .00, .00. .00. .00. .03. 1 .00. .00. .00. .29. .00. .00, .82, 1.49, 3.39, 4.98, .00. 1 .69, 1.47, 3.34, 5.05, .00, .00, .00. .00. .21, .00, .00. 1 .00. .00. .00, .00. .00. .00. .11. .55, 1.15, 2.76, 4,22, 1 .00. .00, .00. .00. .00. .00. .07. .46, 1.03, 2.65, 4.11/ 1 DATA TSB/ .00. .00. .00. .00. .00, .02, .00. .38, .92, 2.42, 3,83, 1 .00, .00. .00. .00. .00. .00, .00, .28, .80, 2.31, 3.73, 1 .89, 1.31, 1.82, 2.33, 3.38, 4.44, 5.49, 7.58, 9.55, 1 .21, .52. .17. .74. 1.12. 1.62. 2.13. 3.18. 4.25. 5.31. 7.43. 9.39. 1 .43. 1 .10. .32, .57. .89, 1.32, 1.78, 2.81, 3.89, 4.98, 7.13, 9.06, .77. 1.17. 1.61. 2.57. 3.64. 4.73. 6.92. 8.84. 1 .06, .26, .49. .39. .65, 1.01, 1,41, 2.34, 3.37, 4.46, 6.67, 8.57, .02. .18. 1 .85, 1.23, 2.12, 3.12, 4.19, 6.42, 8.31, 1 .00, .13, .30. .52. .69, 1.01, 1.85, 2.83, 3.89, 6.13, 7.99, .00. .23. .42. 1 .08. .00. .34. .59. .88, 1.63, 2.58, 3.62, 5.87, 7.71, .75, 1.41, 2.28, 3.31, 5.56, 7.37, .04, .17. 1 .48. .00, .00. .10. .25. 1 .70, 1.35, 2.18, 3.21, 5.46, 7.26, .44. 1 .00. .00, .08. .22. .00, .13. .31, .54, 1.15, 1.90, 2.82, 5.07, 6.83, .00. .01. 1 .00, .00. .00. .07. .25. .45, 1.01, 1.74, 2.61, 4.82, 6.55, 1 .85, 1.54, 2.38, 4.51, 6,21, .00. .00, .00, .00. .17. .36. 1 .00. .00. .74, 1.36, 2.17, 4.22, 5.89, 1 .00. .00. .10. .28. .63, 1.20, 1.96, 3.96, 5.56, 1 .00. .00, .00. .00. .05, .21, .00. .00. .15, .53, 1.07, 1.76, 3.71, 5.22, ,00, .00, .00, 1 .00, .00. .00. .00. .06. .41, 1.02, 1.83, 3.79, 5.46/ 1 .00. DATA TSC/ .00. .00. .00. .00. .00. .00. .33. .87, 1.73, 3.72, 5.51, 1 .00. .00, .00. .00. .00. .00. .26. .76, 1.51, 3.39, 5.03, 1 .69. 1.47. 3.34. 5.05. .00. 1 .00. .00. .00. .00. .00. .21. .00. 1 .00, .00. .00. .00. .00. .16. .60, 1.22, 3.02, 4.57,

Figure 13. Flash Flood Model

.34. .77, 1.22, 1.70, 2.21, 2.73, 3.76, 4.79, 5.82, 7.86, 9.85, 1 .60, 1,00, 1,44, 1,96, 2,47, 3,52, 4,57, 5,61, 7,69, 9,66, 1 .25, .92, 1.34, 1.85, 2.37, 3.42, 4.47, 5.53, 7.61, 9.59, .22. .54. 1 .16. .42, .71, 1.08, 1.58, 2.09, 3.14, 4.21, 5.27, 7.40, 9.35, 1 .61. ,12, .34. .92. 1.36. 1.83. 2.88. 3.95. 5.03. 7.18. 9,12. 1 .03. .78. 1.16. 1.63. 2.67. 3.73. 4.80. 6.94. 8.88. 1 .24. .49. .38. .00. .15. .66. .97. 1.43. 2.45. 3.51. 4.57. 6.70. 8.64. 1 .92, 1.30, 2.21, 3.22, 4.30, 6.52, 8.41, ,00, .15. .34. .57. 1 .00, .23. .72, 1.08, 1.94, 2.92, 3.98, 6.19, 8.07, .07. .43. 1 .00. .59, .93, 1.77, 2.73, 3.77, 5.96, 7.83, .01. .15. .35. 1 .00. 1 .00. .10. .25. .48, .75, 1.41, 2.28, 3.31, 5.56, 7.37, .00. .00, .07. .40. .22. .67, 1.39, 2.29, 3.30, 5.49, 7.31, 1 .00. .11. 1 .00, .00. .30. .54, 1.26, 2.15, 3.14, 5.28, 7.12, .00. .00,05. .22, .44. 1.09. 1.96. 2.94. 5.05. 6.89. 1 .00. .00. .14, .35, .93, 1.78, 2.73, 4.83, 6.65/ .00, .00. 1 DATA TSD/ .00. .00. .00. .00. .07. .26. .77, 1.59, 2.53, 4.61, 6.42, 1 .00. .00, .00, .00. .00, .66, 1.43, 2.43, 4.48, 6.33, 1 .16. .00. .00, .00. 1 .00. .00. .09. .54, 1.23, 2.13, 4.16, 5.96, .00. .00, .00. .00. .00. .08. .47, 1.06, 1.78, 3.77, 5.40, 1 .00, .00. .00. .03. .41. .97, 1.67, 3.54, 5.12, .00. .00. 1 .00, .00. .00. .00. .00. .00. .33, .84, 1.51, 3.28, 4.82, 1 1.00, 1.50, 2.00, 2.50, 3.00, 4.00, 5.00, 6.00, 8.00,10.00, 1 .50. .38. .83, 1.30, 1.79, 2.31, 2.82, 3,85, 4.87, 5.89, 7.91, 9.91, 1 1 .30. .69, 1.12, 1.58, 2.10, 2.61, 3.65, 4.69, 5.73, 7.78, 9.77, .22. .54 . -,92, 1.34, 1.85, 2.37, 3.42, 4.47, 5.53, 7.61, 9,59, 1 .81, 1.21, 1.71, 2.23, 3.28, 4.34, 5.54, 7.50, 9.47, .46. 1 .19, .15, .70, 1.04, 1.54, 2.05, 3.10, 4.16, 5.23, 7.36, 9.31, .41, 1 .64. .13, .36, .95, 1.40, 1.90, 2.96, 4.03, 5.10, 7.25, 9.19, 1 .04. .22, .45, .71. 1.09. 1.51. 2.47. 3.51. 4.60. 6.80. 8.71. 1 .01. .16. .36. .61. .96+ 1.35+ 2.27+ 3.29+ 4.37+ 6.59+ 8.48+ 1 .00. .14, .31, .54. .88, 1.26, 2.15, 3.16, 4.23, 6.46, 8.35, 1 .69, 1.01, 1.85, 2.83, 3.89, 6.13, 7.99, .61, .90, 1.66, 2.61, 3.66, 5.91, Ø.75, .00, .42. .08. 1 .23, .00. .04. .18, .35. 1 .00. .00, .13, .29. .51, .79, 1.54, 2.47, 3.50, 5.71, 7.55/ 1 DATA TSE/ .32, .00. .00, .00. .16. .61, 1.40, 2.40, 3.43, 5.52, 7.42, 1 .57, 1.19, 2.18, 3,21, 5,28, 7.18, .21. .00, .00. .00. .15. 1 .00. .00, .00. .19, .50, 1.10, 1.84, 2.73, 4.98, 6.73, 1 .10. .00, .87, 1.75, 2.76, 4.81, 6.70, .00, .00. .00. .11, .43, 1 .00, .00. .00. .00. .05. .35, .83. 1.51. 2.35. 4.46. 6.16. .00. .00, .00, .00. .03, .21, .66, 1.30, 2.09, 4.11, 5.77, 1 .00, .00, .00. .00. .00. .14, .57, 1.14, 1.90, 3.83, 5.45, 1 .00. .00. .00. .00. .00. .06. .41, 1.02, 1.83, 3.79, 5,46/ 1 DATA RRFF/ 1 .5, 1.0, 1.5, 2.0, 2.5, 3.0, 4.0, 5.0, 6.0, 8.0, 10.0/ DATA FFFF/ 1 1.0, 1.2, 1.4, 1.6, 1.8, 2.0, 2.2, 2.4, 2.6, 2.8, 3.0, 3.2, 3.4, 13.6, 3.8, 4.0, 4.2, 4.4, 4.6, 4.8, 5.0/ DO 4311 I = 1+11RF(I) = RRFF(I)4311 $D0 \ 4312 \ I = 1.21$ FF(I) = FFFF(I)4312 D0 5600 I = 1.209TS(I) = TSA(I)5600 DO 5601 I = 210.418 5601 TS(I) = TSB(I-209)Figure 13. (continued)

Flash Flood Model

	DU = 502 I = 4194527	
5602	1S(1) = 1SC(1-418)	
	DO 5603 I = 628+836	
5603	TS(I) = TSD(I-627)	
	DO 5604 I = 837,924	
5604	TS(I) = TSE(I-836)	
5559	READ(5,104,END=5560)BASIN	:
104	FORMAT(13A4)	
5561	PEADIS, 54 IPFACH	
5001	EOOMATIIZAUI	•
54		
	READISTICSTURIE	
105	FURMAT(2A4)	
	READ(5,1300)STATE	
1300	FORMAT(3A4)	·
	READ(5,4001)STATEN	,
4001	FORMAT(F2.0)	x
	READ(5,4002)RFPMS	
4002	FORMAT(F5.1)	
1	READ(5.4003)DENELP	
"003		
4000		
	READISIGISIONIUN	
4161	FURMAT(FIU.U)	
	READ(5,4004)DRAINA	
4004	FORMAT(F8.1)	·
	READ(5.4005)RAINFA	·
4005	FORMAT(F6.2)	
	READ(5,4006)SLOPCH	
4006	FORMAT(F6.1)	
	RFAD(5,4007)SLOPBA	
4007	FORMAT(F6.1)	
1001	READ (5.4008) SLOPST	
"		
4000		
	READ(31401211)	
4012	FURMAT(FD.2)	•
	READ(5:4044) IPP	
4014	FORMAT(F6.2)	
	READ(5+4015)UNITT	
4015	FORMAT(F8.2)	
	READ(5,4016)HIGH	
4016	FORMAT(F7.1)	
	READ(5,4017)ZERO	
4017	FORMAT(F7.1)	
	READ(5,4018)DUR	
4018	FORMAT(F4.1)	
1010	PEAD(5,4019)RM	
4013		•
	READISIGUOTANPL	
4000	FORMATIF6.2)	
	READ(5+6996)FSI	
6996	FORMAT(F5.2)	
	READ(5,6984)00PMS	
6984	FORMAT(F10.0)	
	DEATHS =((-71.6413)+(2.5355*R	FPMS)+(.0163*DENFLP)+(-0.1005*DRAINA)
	1+(-2.3284*RAINFA)+(-2.41*SLOP	CH)+(0.7283*SLOPBA))
	IF(SLOPSI-16,)8730,8731,8731	
8730	DEATHS = DENFLP**.4	
0.00	SCALF = 6.	
	Figure 12	
	Figure 13. (continued)	rlash rlood Model

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GO TO 4044 8731 CONTINUE MINRAF =0.15*PMS IF(STATEN-1,)4010,4011,4010 4011 Q(1) = 0.0Q(2)=75.*DRAINA**.552*SLOPBA**.460 Q(3)=144.*DRAINA**.528*SLOPBA**.336 Q(4)=891.*DRAINA**.482*SLOPBA**.154 Q(5)=1770.*DRAINA**0.463*SLOPBA**.086 Q(6)=5770.*DRAINA**0.432 0(7)=10020.+DRAINA++0.375 IF (QQPMS-1.)6989.6988.6988 6988 Q(7) = OQPMS6989 CONTINUE STAGE(1)=ZERO STAGE(2)=(22.*SLOPCH**(-.334))+ZERO STAGE(3)=(35.55*SLOPCH**(-.462))+ZERO+1.5 STAGE(4)=(52.1*SLOPCH**(-.50))+ZERO+3. STAGE(5)=(59.3*SLOPCH**(-.533))+ZERC+4.5 STAGE(6) = (77.3*SLOPCH**(-.553))+ZERO+6. STAGE(7)=(136.2*SLOPCH**(-.703))+ZERO GO TO 4013 4010 0(1)=0.0 Q(2)=0.50*DRAINA**0.66*SLOPCH**0.40*ANPC**1.58 Q(3)=1.08*DRAINA**0.67*SLOPCH**0.42*ANPC**1.44 Q(4)=5.40*DRAINA**0.69*SLOPCH**0.47*ANPC**1.12 Q(5)=9.14*DRAINA**0.70*SLOPCH**0.48*ANPC**1.01 Q(6)=24.4*DRAINA**0.72*SLOPCH**0.48*ANPC Q(7)=80.2*DRAINA**0.75*SLOPCH**0.75*ANPC IF(QQPMS-1.0)6987.6986.6986 6986 Q(7) = QQPMS6987 CONTINUE STAGE(1)=ZERO STAGE(2)=(0.53*DRAINA**0.24*TY**1.60)+ZERO D STAGE(3)=(0.83+DRAINA+*0.22*TY**1.40)+2ERO+1.5 STAGE(4)=(1.58*DRAINA**0.20*TY**1.14)+ZERO+2.5 STAGE(5)=(1.95*DRAINA**0.19*TY**1.0E)+ZERO+3. STAGE(6)=(2.85*DRAINA**0.17*TY**0.91)+ZERO+4. STAGE(7)=(6.25*DRAINA**0.16*TY**0.75)+ZERO 4013 CONTINUE DIFF=HIGH-ZERO IF(TPP-0.25)4020,4021,4021 4020 TC=((11.9*RM**3.)/DIFF)**0.35 TP=((DUR/2.)+(0.6*TC))*.7 XK=0.43+0.0003*DRAINA*640. TR=1.38*XK UNIT = DRAINA*640.*1.0*(2./(TP+TR)) GO TO 4022 TP=TPP 4021 UNIT = UNITTCONTINUE 4022 IF (DENFLP-1,)7710,7711,7711 DEATHS = 0.77.10 GO TO 4023 CONTINUE 7711 IF (DEATHS-2.0)4023.4034.4034 4023 IF((STAGE(2)-STAGE(1))-2.)4025.4026.4026

4026	IF((STAGE(2)-STAGE(1))-4.)4027.4028.4028
4028	IF((STAGE(2)-STAGE(1))-6.)4029,4030,4030
4030	1F((STAGE(2)-STAGE(1))-8.)4031.4032.4032
4032	1F((STAGF(2) - STAGF(1)) - 10.) + 033. + 034. + 034
4034	1F(DEATHS-2) 4035.4036
1036	
4030	
4030	IF (DEATHS-150.)4039;4040;4040
4040	1F (DEATHS-400.)4041.4042.4042
4042	IF(DEATHS-10000.)4043.4044.4044
4025	SCALE=1.
	GO TO 4044
4027	SCALE=2.
	GO TO 4044
4029	SCALE=3-
1022	
0031	
4051	
4035	SLALE=5.
	GO TO 4044
4035	SCALE= 6.
	GO TO 4044
4037	SCALE=7.
	GO TO 4044
4039	SCALE=8.
	GO TO 4044
4041	SCALE=9.
4011	
4045	
4044	CONTINUE - CONTRACT - CONTRACT
	IF (SLOPS1-5.0)//12.7/13.7/13
7712	DEATHS = (DRAINA*SLOPSI)**•5
	SCALE = 6.
7713	CONTINUE
	IF(DEATHS-1.)7715.7716.7716
7715	DEATHS = 0.0
7716	CONTINUE
	BAGANO = DBAINA**.5
	BVGANO = BM/6.
4010	$\mathbf{x} \mathbf{y} \mathbf{y} \mathbf{x} \mathbf{x} \mathbf{y} \mathbf{x} \mathbf{y} \mathbf{x} \mathbf{y} \mathbf{x} \mathbf{y} \mathbf{x} \mathbf{x} \mathbf{y} \mathbf{x} \mathbf{x} \mathbf{x} \mathbf{x} \mathbf{x} \mathbf{x} \mathbf{x} x$
4041	
4048	FFA = 1
	GO TO 4050
4049	FFA = 0.
4050	CONTINUE
	WRITE(6,4051) DATE
4051	FORMAT(1H1,////.66X,2A4)
	WRITE(6,4052)BASIN
4052	FORMAT(//,19X,13A4)
	VRITE(6,4053) REACH
4053	
7000	
	RKIILIDITUTTIN ALLAND DATENTINI TA PETTINTED AT A TA A SE A SELL
4054	FURNALIZZATIELUUD PUTENITAL IS ESTIMATED AL 14124 OF A SCALE
	WRITE(6,4055)
4055	FORMAT(//+15X)+FLOOD POTENTIAL SCALE IS LISTED BELOW;")

Figure 13. (continued) Flash Flood Model

WRITE(6.4056) 4056 FORMAT(/.18X. 1 = NO SIGNIFICANT FLOODING.) WRITE(6,4057) FORMAT(/+18X+'2 = SOME STREET AND LOW LAND FLOODING.") 4057 WRITE(6,4058) 4058 FORMAT(/,18X,'3 = STREET AND SOME RESIDENTIAL FLOODING.') WRITE(6,4059) FORMAT(/+18X++4 = MAJOR STREET AND RESIDENTIAL FLOODING. *) 4059 WRITE(6,4060) FORMAT(/.18X.'5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.') 4060 WRITE(6,4061) FORMAT(/+18X++6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND 4061 11) WRITE(6,4101) FORMAT(22X, 'PROPERTY LOSS, LIGHT.') 4101 WRITE(6,4062) 4062 FORMAT(/+18X+'7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 11) WRITE(6,4102) 4102 FORMAT(22X. 50) AND PROPERTY LOSS. MODERATE. 1) WRITE(6,4063) 4063 FORMAT(/+18X+'8 = STREAM FLOODING WITH LOSS OF LIFE (50 TO 150) AN 1D') WRITE(6,4103) 4103 FORMAT(22X, PROPERTY LOSS, MAJOR. ") WRITE(6.4064) 4064 FORMAT(/+18X+'9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400)*) WRITE(6,4104) FORMAT(22X, 'AND PROPERTY LOSS, MAJOR. ') 4104 WRITE(6,4065) FORMAT(/.18X.'10 = CATASTROPHIC FLOOD.') 4065 NDEAD = DEATHSWRITE(6,4105)NDEAD FORMAT(///,18X, POSSIBLE DEATHS FROM PROBABLE MAX STORM WOULD BE * 4105 1,15) WRITE(6,4066) FORMAT(///.35X, 'RIVER FORECAST CENTER') 4066 WRITE(6,4067) 4067 FORMAT(/.38X.'TULSA OKLAHOMA') WRITE(6.4068)DATE FORMAT(1H1,//////.66X,2A4) 4068 WRITE(6.4069)BASIN 4069 FORMAT(//.19X.13A4) WRITE(6,4070)REACH 4070 FORMAT(/+19X+11A4) WRITE(6,4071) FORMAT(///.18X.'PERTINENT DATA') 4071 WRITE(6.4072)DRAINA 4072 FORMAT(/+18X+'DRAINAGE AREA = '+F6.2+' SOUARE MILES') WRITE(6,4073)SLOPCH FORMAT(/.18X. 'CHANNEL SLOPE = '.F6.1.' FEET/MILE') 4073 WRITE(6,4074)SLOPSI FORMAT(/:18X: 'OVERLAND SLOPE = '.F7.1.' FEET/MILE') 4074 WRITE(6.4075)HIGH 4075 FORMAT(/.18X. MAXIMUM HEIGHT OF BASIN = ".F7.1." FT MEAN SEA LEVEL 11) WRITE(6.4076)ZER0

FORMAT(/.18X. ZERO OF LOCATION = '.F7.1. FEET HEAN SEA LEVEL'; 407.6 IF(FSI-1.0)6999.6998.6998 6998 FS = FSIGO TO 6997 6999 FS = STAGE(2) - ZERO6997 CONTINUE WRITE(6,4077)FS FORMAT(/.18X. FLOOD STAGE = '.F7.1. FEET') 4077 NNQ = Q(2)/10.WRITE(6,4112)NNO FORMAT(/+18X+'FLOOD STAGE FLOW = '+14+'0. CFS') 4112 xG = STAGE(5) - ZERONNXG = XGWRITE(G.4078)NNXG FORMAT(/+18X+'FLOW DEPTH FROM 100 YEAR FLOOD = '+16+', FEET') 4078 IGG = Q(5)/100. WRITE(6,4113)IGG FORMAT(/.18X.'100 YEAR FLOOD FLOW = '.IS.'00, CFS') 4113 PMS = (RFPMS*RAINFA)/100. WRITE(6,5110)PMS FORMAT(/.18X. PROBABLE MAXIMUN PRECIPITATION STORM = '.F5.2, 'INCH 5110 1ES!) PMSS = STAGE(7)-ZERO WRITE(6,4114)PMSS FORMAT(/+18X+'DEPTH FROM PROBABLE MAXIMUN PRECIPITATION STORM = "+ 4114 1F4.0, ' FEET') WRITE(6,4140)TP FORMAT(/+18X+'TIME TO PEAK = '+F4.1+' HOURS') 4140 NUNIT = UNIT/100. WRITE(6.4141)NUNIT FORMAT(/+18X+'UNIT GRAPH PEAK = '+14+'00. CFS') 4141 NOQQ = Q(7)/1000.WRITE(6,4115)NG00 FORMAT(/+18X+ PROBABLE MAXIMUM PRECIPITATION STORM FLOOD FLOW = . 4115 113. 000. CFS') NGAG = RAGANOWRITE(6.4130)NGAG FORMAT(//.18X. NUNBER OF RAINFALL GAGES REGUIRED = *.13) 4130 NRVRG = RVGANO WRITE(6.4131)NRVRG FORMAT(/.18X. NUMBER OF RIVER STATIONS REQUIRED = '.12) 4131 READ(5,101)DAM 101 FORMAT(F2.0) DAM BREAK PROGRAM С IF(DAM-1.)4121.106.106 **C** SRDAA = SURFACE ACRES 4121 READ(5:49)DNAME FORMAT(13A4) 49 READ(5.17)SRDAA 17 FORMAT(F7.2) DEPTH = MAX DEPTH OF RESERVOIR AT DAM С READ(5.18)DEPTM 18 FORMAT(F7.2) CAP = RESERVOIR CAPACITY С CAP = SRDAA*0.4*DEPTM READ(5.19)RESLN **C RESLN** = **RESERVOIR LENGTH**

FORMAT(F5.2) B = 2.*CAP/DEPTM*RESLN ORES = PEAK OUTFLOW OF RESERVOIR QRES = (8./27.)*8*32.2**.5*DEPTM**1.5 TK = CAP/ORES*60. TK = TIME IN MINUTES READ(5.60)TW FORMAT(F7.1) READ(5+61)DEPTH FORMAT(F7.1) CP = (TW/(2.*DEPTH))**2. QPQ = CP*4.*(DEPTH**2.5)CAPB = SRDAA*0.4*DEPTH READ(5.70)CAPBB FORMAT(F12.1) IF(CAPBB-1.)71.72.72 CAPB = CAPBBTKK = CAPB/OPQ*60. READ(5,20)CNAME FORMAT(13A4) READ(5,21)MLCITY 21 FORMAT(F6.2) DSSTAG = (DEPTH*0.65)/(MLCITY**0.35) TIM =((1.486*2.8)*((SLOPCH/5280.)**0.5))/.016 TIME = TIM*0.6812 TRAVL = TIME/(MLCITY) DAM BREAK PRINT OUT WRITE(6+3074)DATE FORMAT(1H1,////,66X,2A4) WRITE(6,3075)BASIN FORMAT(//.19X.13A4) WRITE(6,3076)REACH

FORMAT(/+19X+13A4) 3076 WRITE(6,3077)

CONTINUE

19

С

С

60

61

70

72

20

71

С

3074

3075

- FORMAT(///.14X. 'DAM BREAK DATA FOR THE BASIN') 3077
- WRITE(6,3078)DNAME 3078 FORMAT(/.14X. 'NAME OF DAM IS '.13A4)
- WRITE(6,4079)CNAME FORMAT(/+14X+ NAME OF CITY DOWNSTREAM FROM DAM IS ++20A4) 4079 WRITE(6,4080)SRDAA
- 4080 FORMAT(/+14X+'SURFACE ACRES = '+F10.D)
- WRITE(6,4081)DEPTM FORMAT(/+14X+*DEPTH OF THE WATER AT THE DAM = '+F6.2+' FEET') 4081 WRITE(6,4082)CAP
- FORMAT(/+14X+'RESERVOIR CAPACITY = '+F11.1+' ACRE-FEET') 4082 WRITE(6,4083) RESLN
- FORMAT(/.14X. RESERVOIR LENGTH = '.F5.1. MILES') 4083 NQUES = QRES/1000. WRITE(6,4084)NOUES
- FORMAT(/.14X. PEAK CFS FLOW OF TOTAL DAM BREAK = '.16, 000. CFS' 4084 WRITE(6,4085)TK
- FORMAT(/+14X+ RESERVOIR TIME TO PEAK AT SPILLWAY = "+F5.2+" MINUTE 4085 151) WRITE(6,4086)
- FORMAT(//+14X+ WEIR BREACH -- MOST PROBABLE WAY OF DAM FAILURE!) 4086 WRITE(6,4087)TW

4087 FORMAT(/+14X++WIDTH OF BREACH = ++F7+1+* FEET*) WRITE(6,4088)DEPTH FORMAT(/+14X+ DEPTH OF BREACH = "+F7.1+ FEET") 4088 NOPO = 0P0/1000. WRITE(6,4089)NOPO FORMAT(/.14X. PEAK FLOW AT BREACH = ".14. ODC. CFS") 4089 WRITE(6,4090)TKK FORMAT(/+14X++PEAK TIME IN MINUTES AT BREACH = ++F4.0++ MINUTES+) 4090 WRITE(6,4091) FORMAT(//.14X. DOWN STREAM EFFECT FROM BREACH!) 4091 WRITE(6.4092)DSSTAG 4092 FORMAT(/+14X++STAGE AT CITY DOWN STREAM WILL BE NEAR ++F6+2++ FEET 1') WRITE(6,4093)TRAVL FORMAT(/+14X, 'PEAK TRAVEL TIME TO CITY WILL BE = 'F4.1.' HOURS') 4093 106 CONTINUE READ(5,4120)MORDAM 4120 FORMAT(F2.0) IF (MORDAM - 1.)4121,4122,4122 4122 CONTINUE READ(5+1301)ZONE 1301 FORMAT(12) FSS = FSEV(1) = .5 + ZEROEV(2) = (FSS*.25) + ZEROEV(3) = (FSS*.5)+ZEROEV(4) = (FSS*.75) + ZEROEV(5) = FSS + ZEROCALL RATING QX(6) = QX(5) * 2. QX(7) = QX(5) * 3.QX(8) = QX(5) * 4.QX(9) = QX(5) * 5. QX(10) = QX(5)*6.READ(5.31)TZ FORMAT(F4.2) 31 READ(5.59)DATUM 59 FORMAT(F4.0) IF(TZ=,25)34,33,34 33 K = 1 DO 35 I = 1,231ZB(I)=TS(K)*UNIT*TZ K = K + 135 GO TO 301 IF(TZ-.50) 40.39.40 34 39 K = 232D0 41 I = 1.231ZB(I)=TS(K)*UNIT*TZ K = K + 141 GO TO 301 40 IF(TZ - .75)44,43,44 K = 46343 $D0 \ 45 \ I = 1.231$ ZB(I)=TS(K)*UNIT*TZ 45 $\mathbf{K} = \mathbf{K} + \mathbf{1}$ GO TO 301 K = 69444

Figure 13. (continued) Flash Flood Model
•	DO 46 I = 1,231	· ·
•	ZB(I)=TS(K)*UNIT*TZ	
46	K = K + 1	
301	CONTINUE	
	J = 1	
	DO 3501 K = 1.10	
	$D0 \ 3501 \ I = 1,231$	
	ZA(J) = ZB(I) + QX(K)	х Р
	J = J + 1	
3501	CONTINUE	
	DO 803 I=1,2310	
	DO 8033 K = $2,20$	
	1F(Q(K)-ZA(1))8033,345,306	
306	77HIG = Q(K)	
000	7710W = Q(K-1)	
	72FLF = STAGE(K-1)	
	777EL = STAGE(K)	•
	7777 = 77751 = 77515	
	200Y = (7x(y) = 77(0y)/(77y)c=77(0y)	,
	77Y7 = 77Y7 + 700Y	,
	$LLAAL \sim LLALTLUUA TOOD - TTUE + TTYT$	
	2000 = 2200 + 22000	
- 11 E	$\frac{10}{2000} = \frac{1000}{1000}$	
342	2555 = 151AUEINII	
		÷
8033	$\frac{1}{2} = \frac{1}{2} = \frac{1}$	•
307	ZEELV(1) = (ZSSS - ZERU)	
803	CUNIINUE	
	IF (DATUM-1.0750.51.51	
51	00.521 = 1.2510	
52	2EELV(1) = 2EELV(1) + 2ERU	
- •	FSS = FS	
50	K = 1	
	$D0 \ 6677 \ 1 = 1.21$	
•	$DO \ 6677 \ J = 1.11$	
	$EELV(I_{1}J) = ZEELV(K)$	
6677	K=K+1	
	BASEQ = QX(1)	
	CALL STAGG (BASEQ: ZERU: PASEF?	
	CALL TABLE (DATE + STATE + ZONE +	BASEQ+BASEF+FSS+ZERO+TP
	K = 232	
	D0 5312 1 = 1.21	
	D0 5312 J = 1.11	
	$EELV(I_{1}J) = ZEELV(K)$	
5312	K = K+1	
	BASEQ = QX(2)	
	CALL STAGG (BASEQ.ZERO.BASEF)	
	CALL TABLE (DATE, STATE, ZONE,	BASE3.BASEF.FSS.ZERD.TP
	K = 463	<i>.</i>
	DO 5313 I =1+21	
	DD 5313 J = 1.11	,
	EELV(I,J) = ZEELV(K)	
5313	K = K + 1	
	BASEQ = QX(3)	
	CALL STAGE (BASEQ, ZERO, BASEF)	
	CALL TABLE (DATE STATE ZONE)	BASE2+BASEF+FSS+ZERO+TP
* .	K = 694	
	D0 5314 I = 1,21	

Figure 13. (continued) Flash Flood Model

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.

D0 5314 J = 1.11 $EELV(I \cdot J) = ZEELV(K)$ 5314 K = K+1BASEQ = QX(4)CALL STAGG (BASEQ,ZERO,RASEF) CALL TABLE (DATE,STATE,ZONE, K = 925D0.5315 I = 1.21D0 5315 J = 1.11EELV(I,J) = ZEELV(K)5315 K = K+1BASEQ = QX(5)CALL STAGG (BASEQ.ZERO.RASEF) CALL TABLE (DATE.STATE.ZONE. K = 1156D0 5316 I = 1.21D0 5316 J = 1.11EELV(I,J) = ZEELV(K)5316 K = K+1 BASEQ = QX(6)CALL STAGG (BASEQ.ZERO, PASEF) CALL TABLE (DATE, STATE, ZONE, K = 1387D0 5317 I = 1.21D0 5317 J = 1,11 EELV(I,J) = ZEELV(K)5317 K = K+1 BASEQ = QX(7)CALL STAGG (BASEQ, ZERO, BASEF) CALL TABLE (DATE, STATE, ZONE, K = .1618DO 5318 I = 1,21DO 5318 J = 1.11EELV(I,J) = ZEELV(K)K = K + 15318 BASEQ = QX(8)CALL STAGG (BASEQ, ZERO, PASEF) CALL TABLE (DATE, STATE, ZONE, K = 1849D0 5319 I = 1.21D0 5319 J = 1.11EELV(I,J) = ZEELV(K)5319 K = K+1BASEQ = QX(9)CALL STAGG (BASEQ, ZERO, PASEF) CALL TABLE (DATE, STATE, ZONE, K = 2080D0 5320 I = 1.21D0 5320 J = 1.11EELV(I,J) = ZEELV(K)5320 K = K +1 BASEQ = QX(10)CALL STAGG (BASEQ.ZERO.BASEF) CALL TABLE (DATE STATE ZONE) GO TO 5559 CONTINUE 5560 STOP

BASE2+BASEF+FSS+ZERO+TP+RF)

BASE2+BASEF+FSS+ZERO+TP+RF)

BASE3 BASEF (FSS ZERO, TP, RF)

BASE2 + BASEF + FSS + ZERO + TP + RF)

BASE2 . BASEF . FSS . ZERO . TP . RF)

BASE2 . BASEF . FSS . ZERO . TP . RF)

BASE2+BASEF+FSS+ZER0+TP+RF)

Figure 13. (continued) Flash Flood Model

	END
	SUBROUTINE RATING
	COMMON STAFF(7).FV(5).R(7).DX(10).DATHM
	$\int \frac{\partial f}{\partial t} = \frac{1}{2} \int \frac{\partial f}{\partial t} = \frac{1}{2} \int \frac{\partial f}{\partial t} \int \frac{\partial f}{\partial t} = \frac{1}{2} \int \frac{\partial f}{\partial t} \int \frac{\partial f}{\partial t$
	IF = (S AGE(R) - EV(I) / I + (IZ) I S
15	HIGH = O(K)
	XLOW = G(K-1)
	DIFF = HIGH-XLOW
	HHI = STAGE(K)
	HHL = STAGE(K-1)
	HHH = HHI - HHL
	HA = (EV(I) - HHL) / HHH
	AFLOW = HA*DIFF
	QX(I) = AFLOW + XLOW
	60 10 144
12	$O_{X}(T) = O(K)$
• *	
14	
144	CONTINUE
	RETURN
	END
	SUBROUTINE STAGG (BASEQ,ZERO,BASEF)
	COMMON STAGE(7).EV(5).Q(7).QX(10).DATUM
	DO 8033 K = 2.7
4	IF(Q(K)-BASEQ)8033,345,306
306	ZZHIG = Q(K)
	ZZLOW = Q(K-1)
	ZZFLF = STAGE(K-1)
	727EI = STAGE(K)
	22AL = LELL = LELL = 22101 / 122HIG=77100
	- ZWWA -
	$ZZAAZ \rightarrow ZZAZ + ZZAZZA$
	60 10 307
345	2SSS = (STAGE(R))
	GO TO 307
8033	CONTINUE
307	BASEF = (ZSSS-ZERO)
	IF(DATUM-1,0)55,56,56
56	BASEF = BASEF + ZERO
55	CONTINUE
	RETURN
	FND
	SUBROUTINE TABLE (DATE.STATE.ZONE. BASEQ.BASEF.ESS.ZERO.
•	
	DIMENSION DATE(2), STATE(3)
	COMMON BASIN(13), REACH(13)
	COMMON RF(11)+F(21)+LELV(21+11)
	WRITE(6.1033)DATE
1033	FORMAT(1H1,////,66X,2A4)
	IF(TP-13.)56.57.57
56	WRITE(6:4700)
4700	FORMAT(1H +24X+'FLASH FLOOD FORECASTING TABLE')
	GD TO 58
57	WRITE(6,4701)

Figure 13. (continued) Flash Flood Model

FORMAT(1H .27X. FLOOD FORECASTING TABLE) 4701 58 . CONTINUE WRITE(6,1302)STATE 1302 FORMAT(//.55X.'STATE OF '.3A4) WRITE(6,1303)ZONE 1303 FORMAT(/:55X: FORECAST ZONE ::12) WRITE(6,1044)BASIN 1044 FORMAT(//.12X.13A4) WRITE(6.53)REACH 53 FORMAT(/+12X+13A4) WRITE(6,2001) FORMAT(1H .26X, 'CREST GAGE HEIGHT IN FEET') 2001 5303 FORMAT(///.25X.'CURRENT FLOW = '.F7.0.' CFS') FORMAT(/.25X. INITIAL STAGE = '.F7.1.' FT') 5304 WRITE(6.5304)BASEF WRITE(6,107) 107 FORMAT(1H .) WRITE(6,113)FSS FORMAT(1H ,46X, 'FLOOD STAGE = '+ F5.1,' FT.') 113 WRITE(6,108) ZERO FORMAT(1H ,46X, 'GAGE ZERO = ', F7.2, ' FT, -MSL') 108 WRITE(6,5024)TP FORMAT(1H +46X+*TIME TO PEAK = *+F4+1+* HOURS*) 5024 WRITE(6:59) FORMAT(1H , 'FLASH FLOOD') 59 WRITE(6,109) FORMAT(1H , 'GUIDANCE' , 19X, 'INCHES OF RAINFALL IN 3 HOURS') 109 WRITE(6,106) RF FORMAT(1H . VALUES*'.11F6.1) 106 WRITE(6,63) FORMAT(1H ,10X, '-----63 1----!) WRITE(6,305)(FF(I),(EELV(I,J),J=1,11),I=1,21) 305 FORMAT(1X+F4+1+* I*+11F6+1) WRITE(6,6969) FORMAT(1H . . * AVAILABLE FORM LOCAL WEATHER SERVICE OFFICE !) 6969 1112 FORMAT(1H0.) WRÏFE(6.1112) WRITE(6,110) FORMAT(1H0,28X, 'RIVER FORECAST CENTER') 110 WRITE(6,112) 112 FORMAT(1H .31X. 'TULSA OKLAHOMA') RETURN

Multiple Regression Analysis

In this study, a multiple regression technique is used as a way to find the number of deaths.

Annual peak flow of the drainage area will be expressed as a function of drainage basin characteristics and climatic conditions by using multiple regression analysis as an equation of the form

$$Y = F(X_1, X_2, X_3, ..., X_n)$$

where

Y = dependent variable death

F = the function, for example, linear function or logarithmic
function

 $X_1, X_2, X_3, \dots, X_n$ = drainage basin population characteristics and climatic conditions

The weighted least squares will be used to adjust the best fit of the equation when the residuals have different variances.

CHAPTER IV

APPLICATION OF THE FLASH FLOOD MODEL

Flood History of Source Data and Hydrologic Information of Study Area

Major Historic Death Storms

Folsom. The great disaster that struck Folsom, New Mexico (24), August 27, 1908, was the most destructive flood ever witnessed by people of northeast New Mexico. It was caused by a cloudburst west of town on the headwaters of the Dry Cimmarron. Just after a beautiful rain in the evening, the sun set upon a happy and prosperous little town of 800 inhabitants. The next morning it rose in a clear sky upon a scene of destruction, death, desolation and horror.

A lady who lived 8^o miles up the river telephoned the Central telephone office that the most terrific flood that had ever been known here was advancing upon the town. The telephone operator, Mrs. Rooke, faithfully warned all that she could warn of the impending danger. Her office was a small building which turned over as the flood struck it, extinguishing the light and carrying this brave and faithful lady to her death.

The water soon began to spread over the town in high rolling waves. The railroad bridge west of town held it in check for awhile, then it broke and let loose a mighty volume of water that swept

everything along with it. The stream was now nearly half a mile wide and was at least five feet deep in the streets and rushing along with a mad torrential velocity that picked up houses and floated them off like chips. Eighteen people lost their lives in this great northeastern New Mexico disaster.

Rapid City. The Rapid City Flood (25) that occurred on June 9, 1972, was one of the most destructive to occur in the midwest for many years. Rapid City, South Dakoka, with a population of 44,000, is located about 3,000 feet above sea level.

Rapid City takes its name from, and is divided by, Rapid Creek, a typical, boisterous mountain stream draining melted snow and rain waters from the rugged rocky crests and spruce-covered slopes of the central portion of the Black Hills. The creek has its headwaters near 7,140-foot Crooks Tower, just south of Cheyenne Crossing and 34 miles west of Rapid City. As its name suggests, its waters race 4,000 feet downhill to the city, controlled only by a dam at Pactola about midway between the city and the stream's origin. The rushing waters are only momentarily delayed as they pass through Canyon Lake at the western edge of the city. Then they accelerate as they race through the narrow gap, cut over the ages by the stream and framed by Dinosaur Hill on the right bank and "M" Hill on the left bank. The waters of the stream then flow along the main business district of the city on their way to join the Cheyenne River, 30 miles to the southeast on the prairie.

The flooding was not restricted to Rapid Creek but was equally serious to the north on Box Elder Creek, and to the south on Spring

and Battle Creeks, which also flow into the Cheyenne River. Fortunately, the latter two creeks do not travel through large population centers.

Fifteen inches of rain fell at Nemo on Box Elder Creek and 14.5 inches in about 5 hours near Sheridan Lake, located on the divide between Spring Creek and Rapid Creek, southwest of Rapid City. This set the stage for the great flood on Rapid Creek and the utter destruction of two-thirds of the City of Keystone on Battle Creek. (Figure 14).

The heavy sustained rainfall for a period of 3 to 6 hours was centered just to the west and northwest of Rapid City as shown in Figure 7. This precipitation averaged about four times the 6-hour amounts that are to be expected once every 100 years in the area. The resulting runoff produced record floods along Box Elder, Rapid, Spring, and Battle Creeks. Preliminary calculations by the U. S. Geological Survey indicate that Rapid Creek had a peak flow of about 31,200 cubic feet per second, 3 miles above Canyon Lake Dam at 10:45 p.m., and 50,600 cubic feet per second, more than 10 times the flow of any previous flood of record, in downtown Rapid City at 12:15 a.m. High-water marks have been used to establish a high stage of 15.5 feet at the Rapid City gage, which reads 9.0 feet when the creek is bank full.

It appears that the relatively small volume of water normally stored behind the Canyon Lake (about 192-acre-feet) would have contributed little to the downstream flooding. The flood waters above Canyon Lake Dam carried debris which clogged the spillway so that



Figure 14. Total Rainfall During Evening of June 9 into Morning of June 10, 1972 Rapid City, South Dakota (25) the reservoir pool temporarily became 11 to 12 feet deeper than normal. The total storage at this time was about 1,000 acre feet which is five times the normal capacity. Dam failure at 10:45 p.m. released this water causing a giant wave that devastated the urban area in its path and drowned more than 150 people during the next two hours.

<u>Big Thompson</u>. The Big Thompson Flood (26)(27)(28)(290 was one of the most costly flash floods in terms of both lives and property damage ever to occur in Colorado and the western United States. Only the disastrous Rapid City Flood is of comparable magnitude. Although destruction was the greatest in the Big Thompson Canyon, serious flooding occurred on the Cache la Poudre River and on several streams draining the adjacent foothills, including Soldier Canyon and Rist Canyon areas.

The storm cost 139 lives, with five persons still missing in mid-October, and an estimated \$28.8 million in damage. In the Big Thompson Canyon alone, 316 homes were totally destroyed and 73 received major damage; 56 mobile homes were lost and 52 businesses destroyed. Yet conditions could have been far more severe. Flood peaks on the mainstream and on North Fork were not synchronous. The North Fork peaked at 2140 MDT at approximately 8,700 cubic feet per second (c.f.s.) some 40 minutes after the mainstream had peaked at Drake at the confluence. Furthermore, Dry Gulch, a small watershed of approximately four square miles near the head of the Big Thompson, received extremely heavy precipitation. Its flow eroded 6,000 cubic yards of material from the base of Olympus Dam which holds back Lake



Figure 15. Before Big Thompson Flood, Natural Disaster Survey Report 76-1, National Oceanic and Atmospheric Administration (15)



Figure 16. After Big Thompson Flood, Natural Disaster Survey Report 76-1, National Oceanic and Atmospheric Administration (15) Estes. Had the dam failed, damage would have been increased by orders of magnitude (Figures 15 and 16).

Approximate timing of flood peaks has been determined, in part, from eyewitness accounts. The Big Thompson peaked about 8:00 p.m. at Glen Comfort, 9:00 p.m. at Drake, 9:30 p.m. at the Loveland power plant (about halfway between Cedar Cove and Midway), and 11:00 p.m. at the mouth of the canyon. The flood at Dry Gulch, near the head of the canyon, was in response to a late burst of rainfall, for its peak discharge was recorded at 10:30 p.m. The flood on the North Fork was due to heavy rainfall near Glen Haven. Its peak discharge near Drake came at 9:40 p.m. See Figure 17 for Big Thompson rainfall map.

Velocities associated with these high flows were also extreme. Velocities of 25 feet per second clearly were not uncommon. Stream banks and much of the U. S. Highway 34 were undercut and removed, trees and numerous structures including the Loveland power plant were swept away. The largest clast that can be documented to have moved has a maximum intercept of 22 feet. The largest structure to be removed was the nine-foot diameter pipe which spanned the mouth of the canyon.

Tremendous quantities of material were transported by the flood. It has been estimated by one worker that 40 percent of the flow volume consisted of debris. Most of the coarse debris was deposited in the canyon, while most of the flotsam was carried completely out of the canyon.

Flash floods are merciless destroyers. The incredible destructive power and speed with which large volumes of water rush down



Figure 17. Total Rainfall During Evening of July 31, 1976, Big Thompson Canyon, Colorado

mountain slopes and across canyon floors make them killers. The Big Thompson flash flood which struck on the evening of July 31, 1976, was no exception.

Hays, Kansas flood (18) in 1951 was a disastrous Hays Flood. flash flood that occurred on Big Creek tributary of the Smoky Hill River on the 22nd of May. The flood waters caused six deaths and very heavy property losses in parts of Hays, Kansas, where flood depths were reported as great as 15 to 20 feet. The disaster was the worst in the history of Hays with blocks of the city and campus of Fort Hays State College inundated. This flooding followed very intense rains which fell over a limited area centered about 4 miles west and upstream from Hays near Yocemento. A later bucket survey reveals that the 12-inch 48-hour isohyet encompassed about three square miles and the 8-inch isohyet about 25 square miles. The bulk of the rain appears to have fallen in a 3 or 4-hour period beginning around 8:00 p.m. on the 21st, although rain continued until about 7:00 a.m. on the 22nd, and from 1 to 2 inches fell in the same area the night of the 20-21st. Big Creek went over its banks at Hays at about 12:30 a.m. on the 22nd, rose very rapidly and apparently crested at about 2:00 a.m. Heavy damages resulted from the flash flood with preliminary estimates placed at near 2 million dollars.

Basic Storm Data

The Folsom, Rapid City, Big Thompson and Hays Flood were listed as the type typical in the west with which people of the area are

familiar. To create a base for study of a possible relationship between community rainfall and death, 21 additional floods were selected from a 37 year period, 1949 through 1976. The first condition for selection was loss of life. Next a variable geographic selection was maintained so that nearly all segments of the United States were represented. It was noted that many floods resulted in the death of from 2 to 4 people, but to vary the population of the data, only one or two floods were used when the death count was nearly the same.

Table III is a listing of the floods that were selected for the statistical analysis. All floods are of the flash flood type rather than river flood. These communities were selected at random.

Station Name and Location	Stream	Date	Data Number
Folsom, New Mexico	Dry Cimarron	8-27-1908	1
Hays, Kansas	Big Creek	5-21-1951	2
Merrill, Iowa	Floyd River	6-8-1953	3
Heber Springs, Arkansas	Little Red River	8-13-1957	4
Hamburg, Iowa	E. Nishnabotna Rvr.	7-1-1958	5 ·
Charleston, W. Virginia	Local Creek	7-30-1961	6
Sapulpa, Oklahoma	Euchee Creek	6-7-1962	7
Sanderson, Texas	Sanderson Canyon	6-11-1965	8
Greeley, Colorado	South Platte Basin	6-16-1965	9
Colorado Springs, Colorado	Cheyenne Creek	7-24-1965	10
Dallas, Texas	Sabine Basin	4-20-1966	11
Wheeling, W. Virginia	Local Creek	7-4-1969	12
Maury, Virginia	James River	7-4-1969	13
Nelson Count, Virginia	Huffman's Hollow	7-4-1969	14
Payson, Arizona	Mogdlon Run	9-20-1970	15
New Braunfels, Texas	Guadalupe River	5-11-1972	16
Rapid City, S. Dakota	Rapid Creek	6-9-1972	17
Logan County, W. Virginia	Buffalo Creek	2-26-1972	18
Las Vegas, Nevada	El Dorado Canyon	9-14-1974	. 19
San Antonio, Texas	Balcones Escarpment	11-23-1974	20
San Antonio, Texas	Balcones Escarpment	3-26-1973	21

LIST OF FLOODS FOR ORIGINAL DATA BASE - DEATH STORMS

TABLE III (Cont'd)

Station Name and Location	Stream	Date	Data Number
Cairo, Illinois	Local Creek	3-26-1973	22
Las Vegas, Nevada	Virgin River	7-10-1975	23
Nicholas County, Kentucky	Bobtown Creek	8-10-1975	24
Big Thompson Canyon, Colorado	Big Thompson Cr.	8-1-1976	25

LIST OF FLOODS FOR ORIGINAL DATA BASE - DEATH STORMS

To test the Flash Flood Model conclusions from the original 25 data base floods, data from four additional floods was compiled to be analyzed. (See Table IV).

TABLE IV

LIST OF FLOODS FOR TEST DATA BASE - DEATH STORMS

Station Name and Location	Stream	Date	Data Number
Colorado Springs, Colorado	Monument Creek	May 30-31, 1935	1
Enid, Oklahoma	Boggy Creek	October 12, 1973	2
Pueblo, Colorado	Fountain Creek	June 3-4, 1921	3
Tulsa, Oklahoma	Mingo Creek	June 1, 1976	4
			•

Statistical Model

Flood data was obtained from 25 historic flood sites throughout the United States. The collected data begins in 1949 and extends through a portion of 1976. In the search for data the main concern was to collect flood death information. The largest flood recorded during the period was the one on Rapid Creek in South Dakota and the smallest was a flash flood in Arizona where two persons died. The sampling was selected so as not to over emphasize any particular death count or regional location. The data collected was based more on the flash flood type of occurrence rather than the long time-topeak river flood (30-43).

The parameter which proved to be most significant for this study was found to be the RFPMS (Rainfall/Probable Maximum Storm) which is the proportion of Rainfall to Probable Maximum Storm. This most significant test was determined by using the statistical package called "Stepwise Regression" from the Statistic System Package Program (44). See Table IV for listing of input data for Statistical Model and Table V for definition of terms.

A procedure named STEPWISE (44) was used for statistical analysis. The STEPWISE procedure applied four techniques to find which variables of a collection of independent variables should most likely be included in a regression model.

Only one Statistical Analysis System (SAS) variable may be specified as the dependent variable. The last variable in the

VARIABLES statement is taken to be the dependent variable, and the other variables in the VARIABLES statement are taken to be independent variables.

The four techniques of STEPWISE used here are the following:

Forward Selection

This technique finds first the single-variable model which produces the largest \mathbb{R}^2 statistic. \mathbb{R}^2 is the square of the multiple correlation coefficient; it is expressed as the ratio of the regression sum of squares to the (corrected) total sum of squares. For each of the other independent variables, STEPWISE calculates an F-statistic reflecting that variable's contribution to the model were it to be included. If the F-statement for one or more variables has a significance probability greater than the specified "significance level for entry," then the variable with the largest F-statistic is included in the model. F-statistics are again calculated for the variables still remaining outside the model, and the evaluation process is repeated. Variables are thus added one by one to the model until no variable produces a significant F-statistic.

Backward Elimination

In this technique, calculations are first performed for a model including all the independent variables. Then variables are deleted one by one until all the variables remaining in the model produce "partial" F-statistics significant at the specified "significance level for staying in" (at each step, the variable showing the smallest contribution to the model is the one deleted).

Maximum R² Improvement

This technique was developed by calculating regressions on all possible subsets of the independent variables. This technique does not settle on a single model. Instead, it looks for the "best" onevariable model, the "best" two-variable model, and so forth. It finds first the one-variable model producing the highest R² statistic. Then another variable, the one which would yield the greatest increase in \mathbb{R}^2 , is added. Once this two-variable model is obtained, each of the variables in the model is compared to each variable not in the model. For each comparison, the procedure determines if removing the variable in the model and replacing it with the presently excluded variable would increase R², After all the possible comparisons have been made, the switch which produces the largest increase in \mathbb{R}^2 is made. Comparisons are made again, and the process continues until the procedure finds that no switch could increase R^2 . The two-variable model thus settled on is considered the "best" twovariable model the technique can find. The technique then adds a third variable to the model, according to the criteria used in adding the second variable. The comparing-and-switching process is repeated, the "best" three-variable model is discovered, and so forth. This technique differs from the STEPWISE technique in that here all switches are evaluated before any switch is made. In the STEPWISE technique, removal of the "worst" variable may be accomplished without consideration of what adding the "best" remaining variable would accomplish.

Minimum R² Improvement

This technique closely resembles the one just described. Here, though, when a switch is to be made, the switch which produces the smallest increase in \mathbb{R}^2 is the one actually performed. For a given number of variables in the model, the maximum \mathbb{R}^2 improvement technique and the minimum \mathbb{R}^2 improvement technique will usually produce the same "best" model. More models of a given size will be considered when the latter technique is applied.

The parameter which proved to be most significant for this study was found to be the RFPMS (Rainfall/Probable Maximum Storm) which is the proportion of Rainfall to Probable Maximum Storm.

Definition of terms used in the STEPWISE regression is shown in Table V. Originally, the study included five additional terms but these terms were found less significant than the eight chosen, because the five terms deleted were either directly related to or reflected in the eight variables chosen or the significance was minor.

The data input for the statistical model is shown in Table VI. This table lists the input data from the 25 original basins.

TABLE V

DEFINITION OF TERMS FOR ANALYSIS OF VARIANCE TABLE

Variable	Definition
DEATHS	= Number of persons killed (dependent variables)
DRAINA	= Drainage area in square miles
RAINFA	= Amount of rainfall, in inches
RFPMS	= Rainfall/probable maximum rainfall
DENTOT	= Population/size, total
DENFLP	= Population/size, flood plain
SLOPCH	= Slope channel feet/mile
SLOPBA	= Slope basin feet/mile
SLOPSI	= Overland slope feet/mile

In this stepwise regression procedure, several regressions are computed; the first one includes all eight basin and climatic characteristics: rainfall/probable maximum storm, population density of flood plain, drainage area, rainfall, total population density of community, slope of the channel, slope of the basin, and slope of the over land. A "backward elimination" computer program will make the first computation, eliminate the least significant variable, and recompute the regression, then continue the elimination process.

In the "Forward Selection, Stepwise" it was found that three

TABLE VI

DATA INPUT FOR STATISTICAL MODEL

	DEATHS	DRAINA	RAINFAL	PROMAXS	REPHS	DENFLPL	SLOPEHN	DENTOTA	SLOPBAS	SLOPS10	
3	18	50	8.0	22.0	36.0	3000	33.00	2000	101.00	300.00	
2	6	25	12.0	25.8	46.5	2400	7.00	3600	, S0'00	41.00	
3	14	22	13.0	25.1	51.8	2000	8,00	1210	25.00	40.10	
	3	60	8.0	30.1	26.5	2666	7.00	5025	21.00	37,00	
5	19	· 70	12.0	25.6	46.9	2800	8.20	2971	26.20	45.20	
6	22	30	5.0	28.0	17.9	2857	14.60	5881	23.00	42,00	
.7	2	7	8.0	28.9	27.7	500	50,00	. 200	110,00	260.00	
8	22	32	- 12.0	30.2	39.7	3400	10,00	. 3067	25,00	60.00	
5	11	26	8.0	20.2	39.6	1700	76.00	1208	210.00	430.00	
20	3	12	7.0	22.8	30.7	1500	110.00	500	370,00	625.00	
21	27	· 37	7.6	31.8	24.5	1875	, 1.30	1041	8,00	13.00	
2 ב	37	40	23.0	27.2	54.5	750	38.00	2272	66.00	170.00	
3 د	126	90	27.0	27.2	39.3	1111	80,00	18	89,01	35.02	
14	54	10	25.0	27.2	91.9	1167	48,00	227	96.00	190,00	
5د	23	28	5.8	10.0	58.0	5666	55,00	1500	140.00	360.00	
16	18	356	16.0	31.5	50.8	3100	2.00	4545	3.00	10.00	
27	236	100	15.0	19.8	75.8	2777	28.00	4870	120.00	320.00	
J.B	138	9 D	6.0	27.2	22.0	6700	24.00	4286	117,60	340.00	
29	9	23	3.5	17.4	20.1	2643	36,00	2222	240.00	410.00	
20	13	310	10.0	31.5	31.7	3100	1.00	4545	3.00	11.00	
21	· 4	260	13.0	31.5	41.0	2675	1,29	2500	3,30	12.00	
22	10	1.40	10.0	28.1	35.6	3000	0.30	4447	2,50	7.00	
23	2	20	5.1	17.4	29.3	1700	34.00	1666	140.00	\$50,00	
26	2	45	. 4.5	26.1	17.2	. 750	0.46	666	3.,40	6,00	
25	144	70	12.0	22.1	54.3	3083	25,40	1665	65.00	85,00	
			,			· • · •	0			Ş .	

variables were deemed significant at the 0.500 significance level. The three variables were RFPMS (rainfall/probable maximum storm), DENFLP (population density of the flood plain) and drainage area. The output from the analysis is shown in Table VII.

Regression equation based on the three variables would be

Deaths = -83.252 + (1.667) RFPMS + (0.283) DENFLP + (-0.115)

DRAINA (4.1) (Table VII)

In the "Backward Selection, Stepwise" it was found that two variables were deemed significant at the 0.100 significance level. The two variables were RFPMS (rainfall/probable maximum storm) and DEMFLP (population density of the flood plain). The output from the analysis is shown in Table VIII.

Regression equation based on the two "backward" variables would be:

Deaths =
$$-82.882 + (1.601)$$
 RFPMS + (0.021) DENFLP (4.2)

In the statistical analysis "Maximum R-Square Improvement" it was found that variable RFPMS (rainfall/probable maximum storm) is the best one variable found by the maximum R-square improvement procedure.

Regression equation based on the "Best" one variable would be:

Deaths = 18.393 + (1.276) RFPMS (4.3)

The output from the "Maximum R^2 Improvement" analysis is shown in Table IX.

TABLE VII

LINEAR MODEL EQUATIONS - DEATHS RELATED

TO EIGHT VARIABLES - FORWARD SELECTION

No. of Variables	Equation	Variables	Observed Significant Level
3	Deaths = -83.252 + 1.067 RFPMS	RFPMS	.0006
	+ 0.024 DENFLP -	DENFLP	.0059
	0.115 DRAINA	DRAINA	.2655

TABLE VIII

LINEAR MODEL EQUATIONS = DEATHS RELATED

.

TO EIGHT VARIABLES - BACKWARD SELECTION

No. of Variables	Equation	Variables	Observed Significant Level
2	DEATHS = -82.882 + 1.601 RFPMS	RFPMS	.0010
	+ 0.021 DENFLP	DENFLP	.0094

TABLE IX

TO EIGHT VARIABLES -- MAXIMUM R² IMPROVEMENT

No. of Best Variables		Equations	Var	iables	Observed Significant Level
2	DEATHS =	-82.882 + 1.607 R	FPMS	RFPMS	0.0010
		+0.021 DENFLP		DENFLP	0.0094
3	DEATHS =	-83.253 + 1.667 B	FPMS	RFPMS	0.008
		+0.024 DENFLP		DENFLP	0.0059
		-0.115 DRAINA		DRAINA	0.2659
4	DEATHS =	-81.251 + 2.292 R	FPMS	RFPMS	0.0391
		+0.023 DENFLP		DENFLP	0.0115
		-0.083 DRAINA		DRAINA	0.5195
		-2.606 RAINFL		RAINFL	0.5307
5	DEATHS -	-83.024 + 2.361 R	FPMS	RFPMS	0.0400
		+0.021 DENFLP		DENFLP	0.0489
		-0.097 DRAINA		DRAINA	0.5661
		-2.971 RAINFAL		RAINFAL	0.5150
		+0.003 DENTOTA		DENTOTA	0.6667
6	DEATHS =	-71.641 + 2.536 F	FPMS	RFPMS	0.0349
		+0.016 DENFLP		DENFLP	0.0897
		-0.101 DRAINA		DRAINA	0.5964
		-2.328 RAINFA		RAINFA	0.5967
		-2.414 SLOPCH		SLOPCH	0.1281
		+0.728 SLOPBA		SLOPBA	0.1365
7	DEATHS =	-63.577 + 2.974 B	FPMS	RFPMS	0.0279
		+0.017 DENFLP		DENFLP	0.0840
		-0.106 DRAINA		DRAINA	0.6185
		-4.396 RAINFL		RAINFA	0.628
		-2.721 SLOPCH		SLOPCH	0.0998
		+1.079 SLOPBA		SLOPBA	0.0998
		-0.141 SLOPSI		SLOPSI	0.5933

TABLE IX (Continued)

No. of Best Variables	Equations	Variables	Observed Significant Level
8	DEATHS = -66.316 + 2.968 RFP1	MS RFPMS	0.0320
	+0.019 DENFLP	DENFLP	0.1524
	-0.111 DRAINA	DRAINA	0.6176
	-4.352 RAINFA	RAINFA	0.5931
	+0.002 DENTOT	DENTOT	0.8227
	+1.069 SLOPBA	SLOPBA	0.1142
	-2.659 SLOPCH	SLOPCH	0.1234
	-0.140 SLOPSI	SLOPSI	0.5760

The minimum R-square improvement model was run on the previous listed eight variables. The results are much the same as the maximum R-square improvement model.

"Best" 1 variable

Deaths =
$$-18.393 + (1.276)$$
 RFPMS (4.4)

"Best" 2 variables

Deaths = -82.882 + (1.607) RFPMS +(0.021) DENFLP (4.5)

"Best" 3 variables

Deaths = -83.252 + (1.667) RFPMS + (0.024) DENFLP (4.6) + (-0.115) DRAINA

"Best" 4 variables

Deaths = -81.252 + (2.292) RFPMS + (0.023) DENFLP (4.7) + (-0.083) DRAINA + (-2.606) RAINFAL

"Best" 5 variables

Deaths = -83.824 + (2.361) RFPMS + (0.021) DENFLP (4.8) + (-0.097) DRAINA + (-2.771) RAINFAL + (0.003) DENTOT

"Best" 6 variables

"Best" 7 variables

"Best" 8 variables

The equation developed through the "best" of six in the minimum R^2 selection wasodetermined to be the most representative of the family of equations created. The "best" of six equations has as the independent variables; drainage area (DRAINA), amount of rainfall (RAINFA), the ratio of probable maximum rainfall to storm rainfall (RFPMS), density of flood plain (DENFL), slope of channel (SLOPCH), and slope of basin (SLOPBA). The selection was made this way due to (1) R^2 value of 0.56 and (2) the availability of source data.

Flash Flood Model

Application of the Flash Flood Model for the purpose of more realistic projection of flood flows is easily done. The Flash Flood Model is designed to solve many hydrologic problems when the data base is complete or partially complete. For this study measurements and calculations were made during field trips as if working in a laboratory on a miniature river basin. This study was, in fact, a laboratory research project on site for the creation of practical data. During the field trips it was found that many times certain basic measurements could not be obtained easily. As a result of this problem, several options were built into the program allowing the model to calculate a hard to determine parameter or to accept a specific input measurement. Items that are optional for entry in the Flash Flood Model are: time to peak, unit graph peak, flood stage, maximum probable storm and type of empirical equations to be used. There are many times when calculated data must be checked by historical information. One instance of the need of the selection capability is when, by local stream improvement, a natural flood stage has been revised due to the construction of a flow way.

The creation of the master files for Antecedent Precipitation Index (API) values (Data files TBA, TBB, TEC and TBD) makes it possible for their use by the design hydrologist. Through the use of these API files or tables, it is possible to create any soil moisture condition required of a design Engineer. The water balance characteristics of any hydrologic model are critical. The model must

predict what the soil moisture is as closely as possible. In this Flash Flood Model the system of percolation, soil moisture, storage, drainage, and evaporation characteristics are represented in the master API files. All variables in the water cycle are important to the hydrologic process. The main element of this, or any other flow model, is that, by its use it is possible to forecast or determine what the basin streamflow will be from a particular storm. The very nature of a flash flood will minimize certain hydrologic parameters while maximizing other parameters.

In the Flash Flood Model the geographic API files allow the modeler to select the logical file, or if a particular condition exists, the modeler may theoretically move the basin 200 miles east or west.

The determination of the time to peak and unit graph may be done by the model, but if possible a site investigation should be made to establish what the local community uses for a flood stage. The engineer must be careful to tie in theory with the local datum base, as a forecast scheme that would forecast for an unknown bridge or non-local datum control condition could result in a disastrous situuation rather than a good forecast. Many times two forecast schemes may be required for one community, one scheme for the natural creek bed in the rural areas, and one scheme for the improved channel section through town.

The Flash Flood Potential Scale, which is a subroutine output in the Flash Flood Model, is designed to present to local, state and Federal agencies what could occur from a flood caused by rains of a Probable Maximum Precipitation (PMP) Storm. This Flash Flood Potential Scale is divided in values of one through ten. The scale includes conditions from no flood to catastrophic floods. The scale is determined in the model by calculation of the possible deaths that may occur from the PMP storm. The Flash Flood Model may also be run with rainfall less than from a PMP storm, and if this is done, then a death count related to the subject storm will be given.

In this research of a forecast scheme for the test communities in the states of New Mexico, Oklahoma and Texas the design was to develop the Flash Flood Potential Scale for the PMP storm. The basic data from the design basin was compiled into the data requirements, as snown in Figure 6. A typical Flash Flood Potential Scale is shown in Table I. The Flash Flood Model was run and the resulting Flash Flood Potential Scale determinations are listed in Table X.

Flash Flood Potential Scales for each test community are listed in Appendix A.

A typical Flash Flood Potential Scale with calculated death count and Scale number is shown in Figure 18.

Flash Flood Forecasting schemes were developed for all design basins and are shown in Appendix B.

A typical Flash Flood Forecasting Scheme is shown in Figure 19.

TABLE X

POTENTIAL SCALE DETERMINATIONS

Location of Community	No. of Possible Deaths	Flash Flood Potential Scale Number
Cimarron, New Mexico	53	8
Clayton, New Mexico	76	8
Folsom, New Mexico	80	8
Raton, New Mexico	115	8
Springer, New Mexico	116	8
Tucumcari, New Mexico	125	8
Ute Park, New Mexico	104	8
Valmora, New Mexico	76	8
Watrous, New Mexico	112	8
Beaver, Oklahoma	0	. 6
Guymon, Oklahoma	61	8
Kenton, Oklahoma	117	8
Amarillo, Texas	27	6
Canadian, Texas	0	6
Palo Duro State Park, Texas	108	. 8
Tascosa, Texas	64	8

CIMARRON RIVER. CIMARRON. NEW MEXICO

EAGLE NEST LAKE TO CIMARRON

FLOOD POTENTIAL IS ESTIMATED AT 8 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

1 = NO SIGNIFICANT FLOODING.

2 = SOME STREET AND LOW LAND FLOODING.

3 = STREET AND SOME RESIDENTIAL FLOODING.

4 = MAJOR STREET AND RESIDENTIAL FLOODING.

- 5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.
- 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT.
- 7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.
- 8 = STREAM FLOODING WITH LOSS OF LIFE (50 TO 150) AND PROPERTY LOSS, MAJOR.
- 9 = STREAM FLOODING WITH LOSSOOF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.

POSSIBLE DEATHS FROM PROBABLE MAX STORM WOULD BE 53

Figure 18. Typical Flash Flood Potential Scale, Cimarron, New Mexico
FLASH FLOOD FORECASTING TABLE

STATE OF N MEXICO

FORECAST ZONE 2

CIMARRON RIVER, CIMARRON, NEW MEXICO

EAGLE NEST LAKE TO CIMARRON CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 1.2 FT

FLOOD STAGE = 4.8 FT. GAGE ZERO = 6400.00 FT.-MSL TIME TO PEAK = 2.4 HOURS

1.0 I 1.5 2.0 2.5 3.1 3.9 4.7 6.1 7.5 7.9 8.8 9. 1.2 I 1.5 1.8 2.3 2.9 3.6 4.4 5.8 7.2 7.9 8.8 9. 1.4 I 1.7 2.1 2.5 3.2 3.8 5.3 6.7 7.7 8.6 9. 1.6 I 1.3 1.6 1.9 2.3 2.9 3.6 5.0 6.4 7.6 8.6 9. 1.6 I 1.3 1.6 1.9 2.3 2.9 3.6 5.0 6.4 7.6 8.6 9. 1.8 I 1.2 1.5 1.8 2.2 2.7 3.3 4.7 6.1 7.5 8.4 9. 2.0 I 1.2 1.4 1.7 2.0 2.5 3.0 4.4 5.7 7.1 8.3 9. 2.2 I 1.2 1.3 1.5 1.8 2.2 2.7 4.0 5.3 6.7 </th <th>. 0</th>	. 0
1.2 1 1.5 1.8 2.3 2.9 3.6 4.4 5.8 7.2 7.9 8.8 9. 1.4 1 1.4 1.7 2.1 2.5 3.2 3.8 5.3 6.7 7.7 8.6 9. 1.6 1 1.3 1.6 1.9 2.3 2.9 3.6 5.0 6.4 7.6 8.6 9. 1.8 1 1.2 1.5 1.8 2.2 2.7 3.3 4.7 6.1 7.5 8.4 9. 2.0 I 1.2 1.4 1.7 2.0 2.5 3.0 4.4 5.7 7.1 8.3 9. 2.2 I 1.2 1.4 1.7 2.0 2.5 3.0 4.4 5.7 7.1 8.3 9. 2.2 I 1.2 1.3 1.5 1.8 2.2 2.7 4.0 5.3 6.7 8.2 9. 2.4 I 1.2 1.3 1.5 1.6 2.2 2.7 4.0 5.3 </td <td>.7</td>	.7
1,4 I 1.7 2.1 2.5 3.2 3.8 5.3 6.7 7.7 8.6 9. 1.6 I 1.3 1.6 1.9 2.3 2.9 3.6 5.0 6.4 7.6 8.6 9. 1.8 I 1.2 1.5 1.8 2.2 2.7 3.3 4.7 6.1 7.5 8.4 9. 2.0 I 1.2 1.4 1.7 2.0 2.5 3.0 4.4 5.7 7.1 8.3 9. 2.2 I 1.2 1.4 1.7 2.0 2.5 3.0 4.4 5.7 7.1 8.3 9. 2.2 I 1.2 1.3 1.5 1.8 2.2 2.7 4.0 5.3 6.7 8.2 9. 2.4 I 1.2 1.3 1.5 1.6 2.2 2.7 4.0 5.3 6.7 8.2 9.	.6
1.6 I 1.3 1.6 1.9 2.3 2.9 3.6 5.0 6.4 7.6 8.6 9. 1.8 I 1.2 1.5 1.8 2.2 2.7 3.3 4.7 6.1 7.5 8.4 9. 2.0 I 1.2 1.4 1.7 2.0 2.5 3.0 4.4 5.7 7.1 8.3 9. 2.2 I 1.2 1.3 1.5 1.8 2.2 2.7 4.0 5.3 6.7 8.2 9. 2.4 I 1.2 1.3 1.5 1.6 2.2 2.7 4.0 5.3 6.7 8.2 9.	.5
1.8 1 1.2 1.5 1.8 2.2 2.7 3.3 4.7 6.1 7.5 8.4 9. 2.0 1 1.2 1.4 1.7 2.0 2.5 3.0 4.4 5.7 7.1 8.3 9. 2.2 1 1.2 1.3 1.5 1.8 2.2 2.7 4.0 5.3 6.7 8.2 9. 2.4 1 1.2 1.3 1.5 1.7 2.1 2.5 3.6 5.0 6.4 8.1 8.	. 4
2.0 I 1.2 1.4 1.7 2.0 2.5 3.0 4.4 5.7 7.1 8.3 9. 2.2 I 1.2 1.3 1.5 1.6 2.2 2.7 4.0 5.3 6.7 8.2 9. 2.4 I 1.2 1.3 1.5 1.7 2.1 2.5 3.6 5.0 6.4 8.1 8.	.3
2.2 I 1.2 1.3 1.5 1.8 2.2 2.7 4.0 5.3 6.7 8.2 9. 2.4 I 1.2 1.3 1.5 1.7 2.1 2.5 3.6 5.0 6.4 8.1 8.	.2
2.4 Y 1.2 1.3 1.5 1.7 2.1 2.5 3.6 5.0 6.4 8.1 8.	. 0
	, 9
2.6 I 1.2 1.2 1.4 1.6 1.9 2.3 3.3 4.6 6.0 8.0 8 .	.8
2.8 I 1.2 1.2 1.3 1.5 1.9 2.2 3.2 4.4 5.8 7.9 8.	.7
3.0 I 1.2 1.2 1.4 1.7 2.0 2.9 4.0 5.3 7.8 8	• 5
3.2 I 1.2 1.2 1.3 1.6 1.9 2.7 3.8 5.1 7.7 8	• 4
3,4 I 1.2 1.2 1.2 1.2 1.5 1.7 2.5 3.5 4.7 7.5 8.	.3
3.6 I 1.2 1.2 1.2 1.4 1.6 2.3 3.2 4.4 7.2 8	. 1
3.8 I 1.2 1.2 1.2 1.2 1.3 1.5 2.1 3.0 4.1 6.8 8	. 0
4.0 I 1.2 1.2 1.2 1.2 1.2 1.4 2.0 2.8 3.8 6.5 7	. 8
4.2 I 1.2 1.2 1.2 1.2 1.2 1.3 1.8 2.7 3.9 6.6 7.	. 9
4.4 I 1.2 1.2 1.2 1.2 1.2 1.2 1.7 2.5 3.8 6.5 7.	. 9
4.6 I 1.2 1.2 1.2 1.2 1.2 1.2 1.6 2.3 3.4 6.1 7.	,7
4.8 I 1.2 1.2 1.2 1.2 1.2 1.2 1.5 2.2 3.4 6.0 7.	. 8
5.0 I 1.2 1.2 1.2 1.2 1.2 1.2 1.4 2.1 3.0 5.6 7.	• 5

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Figure 19. Flash Flood Forecasting Scheme Cimarron, New Mexico

The Flash Flood Model caculates unknown parameters and relists given known parameters in a format called the Pertinent Data Sheet. The Pertinent Data Sheet is calculated during each run of the Flash Flood Model. An example of the Pertinent Data Sheet is shown in Figure 20. During each run of the Flash Flood Model the computer will search the input data for dam break information. If data is found to operate the dam break routine, the Flash Flood Model will calculate the maximum flow from a complete washout of the dam or it will calculate a flow from a breached dam. In most cases a dam will fail by breach. The breach, in most situations, must be estimated in terms of width and depth. The Flash Flood Model calculates the breach flow in maximum flow at the dam face, and then will predict the height of the flood and travel time to the town downstream. See Figure 21 for an example of the dam break sheet.

CIMARRON RIVER. CIMARRON, NEW MEXICO

EAGLE NEST LAKE TO. CIMARRON

Figure 20. Pertinent Data Sheet, Cimarron, New Mexico

CIMARRON RIVER, CIMARRON, NEW MEXICO

EAGLE NEST LAKE TO CIMARRON

DAM BREAK DATA FOR THE BASIN NAME OF DAM IS EAGLE NEST NAME OF CITY DOWNSTREAM FROM DAM IS CIMARRON SURFACE ACRES = 2768. DEPTH OF THE WATER AT THE DAM = 89.00 FEET RESERVOIR CAPACITY = 98540.7 ACRE-FEET RESERVOIR LENGTH = 1.8 MILES PEAK CFS FLOW OF TOTAL DAM BREAK = 5626000. CFS RESERVOIR TIME TO PEAK AT SPILLWAY = 1.05 MINUTES

WEIR BREACH -- MOST PROBABLE WAY OF DAM FAILURE WIDTH OF BREACH = 70.0 FEET DEPTH OF BREACH = 60.0 FEET PEAK FLOW AT BREACH = 37000. CFS PEAK TIME IN MINUTES AT BREACH = 15.8 MINUTES

DOWN STREAM EFFECT FROM BREACH STAGE AT CITY DOWN STREAM WILL BE NEAR 13.92 FEET PEAK TRAVEL TIME TO CITY WILL BE = 1.2 HOURS

Figure 21. Dam Break Sheet, Eagle Nest Dam, New Mexico

MIT Catchment Model

The MIT Catchment sub-routine is used in the Flash Flood Model as a means of determining the unit graph storm or the peak flow from a particular storm. Its use is optional, either the peak flow and time to peak will be determined through hydrologic basin determination as shown in Chapter III, or the MIT Catchment through its more definite control of basin runoff characteristics. Whenever the user of the Flash Flood Model desires a more accurate determination of basin outflow than average utility control, the MIT Catchment Model is used for determination of flow hydrograph. See figure 22 for typical output from the MIT Catchment Model.

TYPICAL EXAMPLE MIT CATCHMENT MODEL DATA INPUT SUMMARY

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CIMARRON RIVER, CIMARRON, NEW MEXICO

NUMBER OF SEGMENTS = 3 DT = 5.000 MINUTES OUTPUT SAMPLING INTERVAL = 5.00 MINUTES NUMBER OF RAIN GAGES = 1 IOPT = 0 NUMBER OF STORMS = 1 NUMBER OF SEASONS = 0 INFILTRATION TYPE = SCS INFL

SEGMENT OF1	UPSTREAM	SEGNENTS	▲DJ	ACENT SEGRENTS	ттре 5	IPR	NDX 10	LENGTH IFEET1 7920.	SLOPE 0-2500	ROL Pari	JGHNESS METER 0.400	
of2 SV3		· ,	OF1	OF2	5	1 1	10 100	7920. 100320.	0.2300 0.0260		0,400 · 0,040	
					THIESSE	N COE	FFICI	ENTS FOR	RAIN GAG	Æ		
				SEGMENT OF1	1.0			· OTHER	PARAMETE NDD D	ERS:	SCS C	Ņ
			ja - 1	0F2 5¥3	1.0			1.	400 0	, Ũ	25.	

COMPUTATION SEQUENCE

INDEX	SEGMENT
1	OF 1
2	OF 2
В	S¥3

KINEMATIC CHANNEL PARAMETERS

SEGNENT	ALPHA	к
OF1	1.488	1.670
OF 2	1,488	1.670
543	2.094	1.330

Figure 22. Typical Example - MIT Catchment Model

TYPICAL EXAMPLE

MIT CATCHMENT MODEL

RAINFALL DATA SHEET

CIMARRON RIVER, CIMARRON, NEW MEXICO

RAINFALL DATA TIME INTENSITY (IN/HR) (MIN) 1 60.00 4.00 120.00 11.00 180.00 6.00 240.00 3.00

TYPICAL EXAMPLE MIT CATCHMENT MODEL OUTFLOW HYDROGRAPH

	CIMARRON	RIVER.	CIMARRON.	NEW MEXICO
TIME	OUTFLOW H	IYDROGRAF	HS FOR SE	SMENTS
(MIN)	0F1	OF2	SW3	
5.00	0.0	0.0	0.0	
10,00	-0.00	0,00	0.07	
15.00	0.00	0.00	1.59	
20.00	0.01	0.01	11.58	
25.00	0.02	0.02	47.51	
30.00	0.03	0.03	140.06	
35.00	0.04	0.04	334.19	
40.00	0.06	0.06	689,16	
45.00	0.09	0.09	1278.29	
50.00	0.12	0.12	2188.36	
55,00	0.15	0.15	3519,06	
60,00	0.18	0.18	5382,38	
65.00	0,30	0.30	8401.95	
70.00	0.44	0.44	13588,00	
75.00	0.61	0.61	21829,39	
80.00	0.80	0.80	34138.04	
85.00	1.01	1.01	51568.91	
90.00	1.23	1.23	75082.06	
95.00	1,48	1.48	105352.44	
100.00	1.74	1.74	142547.12	
105.00	1.99	1.99	186061.81	
110.00	2.22	2.22	234408.87	
115.00	2.41	2.41	285221.44	
120.00	2.55	2,55	335594.44	
125.00	2,50	2,50	380443.33	
130.00	2.43	2.43	415319.06	
135.00	2.34	2.34	439439.69	
140.00	2,24	2.24	453284.31	
145.00	2.14	2.14	458062.62	
150.00	3.04	2.04	455379.56	
155.00	1,96	1,96	447210.06	
160.00	1.87	1.87	435425.31	
165.00	1.80	1.80	421519.87	
170.00	1.74	1.74.	406795,31	
175.00	1.69	1,69	392215.06	
180.00	1.64	1.64	378390.50	•
185.00	1.54	1.54	364573.00	
190,00	1.45	1.45	350042.56	
195.00	1.,36	1.36	335229,56	
200.00	1.29	1.29	320440.50	

Figure 22. (continued) MIT Kinematic Catchment Model

TYPICAL EXAMPLE MIT CATCHMENT MODEL OUTFLOW HYDROGRAPH

•	CIMARRO	N RTVER.	CIMARRON. NI	EW MEXICO
TIME	OUTFLOW	HYDROGRAP	HS FOR SEGME	ENTS
(MIN)	OF1	OF 2	SW3	
205,00	1,23	1.23	305902.69	
210.00	1.17	1.17	291793.94	
215.00	1.12	1.12	278258,81	
220.00	1.07	1.07	265413.75	
225.00	1.03	1.03	253344.69	
230.00	0.99	0.99	242105.69	
235.00	0.96	0.96	231720.56	
240.00	0.93	0.93	222189.69	
245.00	0.84	0.84	212683.94	
250.00	0.76	0.76	202504.56	
255.00	0.70	• • 0,70	191869.81	
260.00	0.63	0.63	180974,50	
265.00	0.58	0.58	169993.44	
270.00	0.53	0.53	159085.69	
275.00	0.48	0.48	148394.25	
280.00	0.44	0.44	138044.12	
285.00	0,40	0.40	128138.25	
290.00	0.37	0.37	118754.69	
295.00	0.34	0.34	109944.69	
300,00	0.31	0,31	101734,19	
305,00	0.29	0.29	94126.94	
310.00	0.27	0.27	87109.62	
315.00	0.25	0.25	80656.00	
320.00	0.23	0,23	74732.62	
325.00	0.21	0.21	69302.06	
330,00	0.20	0.20	64325.63	
335.00	0.18	0.18	59765,41	
340.00	0.17	0.17	55585.66	
345.00	. 0.16	0.16	51752.96	
350.00	0.15	0.15	48236.52	
333.00	0.14	0.14	45008.09	
360.00	0.13	0.13	42042.00	
365.00	0.12	0.12	39314.86	
375 00		0.11	36803+45	
313.00	0.10	0,10	34474.46 - 39374 37	
300.00 XAS 00	0 NG	0.00	32354,3/ 30399 30	
202.00			00077,20 0858/ CO	
395.00	0.02 0.02	0.02.	20004,07 22907 59	
400.00	0.08	0.08 0.08	25355-85	

Figure 22. (continued) MIT Kinematic Catchment Model

TYPICAL EXAMPLE MIT CATCHMENT MODEL OUTFLOW HYDROGRAPH

	CIMARRON	RIVER, CI	MARRON. NE	W MEXICO
TIME	OUTFLOW H	IYDROGRAPHS	FOR SEGME	NTS
(MIN)	OF1	OF2	SW3	
405.00	0.07	0.07	23918.84	•
410.00	0.07	0.07	22586.68	
415.00	0.07	0.07	21350.46	
420.00	0.06	0.06	20202.09	
425.00	0.06	0.06	19134.21	
430.00	0.06	0,06	18140.15	
435.00	0.05	0.05	17213.84	
440.00	0.05	0.05	16349.77	
445.00	0.05	0.05	15542,95	
450.00	0.0_	0.05	14788.80	
455,00	0,05	0.05	14083,18	
460,00	0.04	0.04	13422,31	
465.00	0.04	0.04	12802.75	
470.00	0.04	0.04	12221,35	
475.00	0.04	0,04	11675.27	
480.00	0.04	0.04	11161.85	
485.00	0.03	0.03	10678.71	
490.00	0.03	0.03	10223.67	
495.00	0.03	0.03	9794.71	
500,00	0.03	0.03	9389,99	
505,00	0.03	0.03	9007.82	
510.00	0.03	0,03	8646.64	
515.00	0.03	0.03	8305,05	
520.00	0.03	0,03	7981.70	
525.00	0.03	0.03	7675.41	
530.00	0.02	0.02	7385,05	
535.00	0.02	0.02	7109,59	
540.00	0.02	0.02	6848.07	
545.00	0.02	0.02	6599,62	
550.00	0.02	0.02	6363.42	
555.00	0.02	0.02	6138.73	
560.00	0.02	0.02	5924,84	
565.00	0.02	0.02	5721,09	
570.00	0.02	0.02	5526,90	
575.00	0.02	0.02	5341.70	
580.00	0.02	0.02	5164.96	•
585.00	0.92	0.02	4996.20	
590.00	0,•02	0.02	4834.98	
595.00	0.02	0.02	4680.87	
600.00	0.02	0.02	4533.47	

Figure 22. (continued) MIT Kinematic Catchment Model

CHAPTER V

RESULTS

Flash Flood Model

Death Potential Scale

The data test bed was supplied by the original twenty-five locations. Statistical analysis was made and results indicated that a projection of possible deaths that could occur, could indeed be made. Once the equation with the best fit was determined, an analysis was made to see just how the developed equation would forecast the source data. Forecast values were determined for each input data point. It is shown that the residual value was as high as 92 and as low as -52 (Table VI) for a plotting of the relationship between the forecast value and the observed value. The statistical equation was applied back against the original input data to determine the relationship between forecast residual and observed deaths, (Figures 23 and 24). The final statistical equation (best of six variables) is not the ultimate equation. It appears that the developed equation has a bias to under forecast.

To further test the statistical equation an additional set of death data was applied against the Flash Flood Model. The second set of data was compiled from storms in the southwest. Two of the new storms are from Oklahoma and two are from Colorado. See Table XII







Figure 24. Observed Values vs. Residual - Death Count

TABLE XI

COMPARISON OF DEATHS

Station Location and Stream Name	Date		Observed No. of Deaths	Forecast No. of Deaths
Tulsa, Oklahoma Mingo Creek	June	'76	3	4
Enid, Oklahoma Boggy Creek	Oct.	'73	9	8
Pueblo, Colorado Fountain Creek	June	'21	78	60
Colorado Scprings, Colorado, Monument Creek	May	'35	18	22

for comparison data between reported deaths and forecast deaths. In the second set of data it shows that the confidence level is \pm 25%, which is better than the original data set. The second set of storms shows further the worth of the Flash Flood Potential in that predictions are close and certainly do give an indication of how great a disaster might occur. As stated for the test data of the original set and now for this second set of data, it appears that the statistical equation may have a bias to under forecast the death count. In the second set, the forecast death count for the flood in Tulsa and Colorado Springs is over forecasted, but not to the extent it under forecasted the Enid and Pueblo floods.

The subroutine for the development of the potential death count was, in the beginning, designed as a tool whereby through its use a community could be pre-warned of the approximate death count that could occur from the PMP storm. The data will seldom reach the limit of the PMP storm so the Flash Flood Model will accept any storm rainfall and then give a potential death count.

Flood Forecasting Schemes

Through the utilization of the Antecedent Precipitation Index master files as shown in lines 11 through 96 of Figure 11, the Flash Flood Model makes it possible to forecast any rise in any stream. The forecast tables are dependent on basin parameters, allowing quick "look-up" forecasts of floods and which permit on-the-spot updating as the storm progresses.

To present the reliability and worth of the Flash Flood tables,

(See Table XII). a comparison of forecasts was made. This table was compiled from recent rises in local streams over the past two years. The table lists the forecasts made by the Sacramento Model, the Existing Manual Model (the old data historic API method) and the Flash The Sacramento Model and Manual Model are currently Flood Model. used for public forecasts, the Flash Flood Model is to be used when time prohibits use of the big model and is to be used as a backup In the majority of the floods the Flash Flood Model was not system. checked or looked at until after the forecast was made. Many times the Flash Flood Model was not referred to at all. As stated, in the majority of cases, the Flash Flood Model was not referred to until the issuance of the forecast, but in the case of the two major differences, the Augusta, Kansas and Bartlesville, Oklahoma floods, the tables were used by the operational hydrologist. A statistical analysis was then run on the list of forecasts. The analysis of Variance Table indicated that the forecast of the Flash Flood Model holds 98% of the control of the true forecast, or the observed value.

The statistical analysis gave an equation:

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> True Forecast = (-1.167) + ((0.018) non-flash flood (5.1)model) +((0.999) flash flood model)

it was thus dictated to use 0.999% of the value of the Flash Flood Model and to use 0.018% of the non-Flash Flood Model schemes.

The Flash Flood Model will calculate the equipment needed to satisfy the data needs. The Model will also enable the hydrologist to predict dam breaks in terms of the flow at the breached dam and the stage and time of travel to the community downstream.

TABLE XII

Method of Forecast Sacra-Flash mento Manual Flood Observed Model Model Model Value Station - Stream - State Date Feet Feet Feet Feet 5-77 8 9.05 Canadian, Canadian R. Ok. 9.4 873 ms1 873 ms1 873 ms1 Waurika, Beaver Cr. Ok. 5-77 Beggs, Deep Fork R. Ok. 5-77 23 15 12.59 Sperry, Bird Cr. 28 23.5 23.12 0k 5-77 Owasso, Bird Cr. Ok. 5-77 30 24.5 24.95 Bartlesville, Caney R. 0k. 7-76 19 11.2 11 7-76 Augusta, Walnut R. Ks. 38 34. 33.8 El Dorado, Whitewater R. Ks. 7-76 26 24 23.5 Kingfisher, Uncle John C. Ok. 5-77 25 22 22.5 Kingfisher, Kingfisher C. Ok. 5-77 25 18 21.5 Copan, Caney R. 0k. 5-77 22 22 18.4 Ramona, Caney R. Ok. 5-77 25.5 26 23.5 Blackwell, Chikaskia R. Ok. 22 5-77 24 22.2 Murdock, Ninnescah R. 5-77 9.7 10 9.76 Ks. Blue, Blue R. Ok. 5-77 25 13.5 9.3 Caney, Caney R. Ok. 5-77 24 15.9 14.6

COMPARISON OF FORECASTS

CHAPTER VI

SUMMARY AND CONCLUSIONS

Summary

After analysis of 25 drainage basins by using the STEPWISE regression techniques relating deaths to eight drainage basin characteristics and climatic conditions, the following results were noted. The ratio of Rainfall to Probable Maximum Precipitation (RFPMS) was found to be the most significant in terms of both R^2 (coefficient of determination) and the observed significance level of the coefficient. Although the RFPMS was most significant, an equation developed in the minimum \mathbb{R}^2 was used as it included six basin, population and climatic conditions. In this equation the variables RFPMS, Rainfall (RAINFL) and slope of channel (SLOPCH) were deemed more significant than the density of flood plain (DENFLP), drainage area (DRAINA) and slope of the basin (SLOPBA). The potential death count was determined for the new seventeen test basins. The calculated death count was then set against the step limits of flood potential scale and a scale number was determined for each basin.

Each of the seventeen test basins were investigated for flood stage, unit graph flow, 100 year flow stage and flow, depth of flow from Probable Maximum Precipitation Storm, the number of index rain

gages and stream gages and where there is an upstream dam, the downstream time-to-peak and stage were determined. For each basin a set of flash flood tables were generated to further present to the city officials the flood potential of their community. (45)

An operational Flash Flood Model had never been developed in the United States, but with the continued community development in the flood plain some method of quick alert had to be developed. The Flash Flood Model is designed so that it can be used to create flood parameters for a community long before any flood occurs, or it may be used operationally during a flood.

The Flash Flood Potential Scale was developed from twenty-five original floods. The accuracy was tested against the original twentyfive locations and also against four additional locations. The Flash Flood Potential Scale should only be used as an indication or trend of the number of deaths which could occur from a particular storm.

Conclusions

The objective of this research was to create a way of predicting the possible deaths and damage that could occur due to the Probable Maximum Precipitation Storm. To enable practical application of the model, a Flood Potential Scale was developed to assist the hydrologist or layman in determining what relation a particular basin has to other basins of similar characteristics.

Through utilization of the Flash Flood Model on 17 test basins the model was found to be a most useful tool for the members of the community. Before, when a flood occurred in a basin, few had any

idea of the great disaster that could befall a community. The Flash Flood Model will allow the City Engineer or Civil Defense Director of a local city to understand the magnitude of a flood which could occur.

The model will help to create individual data sets for many basins, which up to now were believed to be non-floodable.

The model will also help the responsible Federal, State, County or City officials in development of a flood plain. Now it can be presented to the community an estimate of what the ultimate, probable stage could be expected in a particular reach or at a specific location. Along with estimate of probable maximum stage, the model will determine the time-to-peak and unit graph peak, the 100-year flow and stage and for each location a flash flood table will be developed with varying initial stages. The Flash Flood Table will allow on the spot forecasting for any storm, at any initial antecedent precipitation index

The model will also treat the sudden dam break and will develop the flow wave and route the flow down stream in terms of stage and travel time.

The Flash Flood Model and Flash Flood Potential Scale is original in the basic hypothesis, in concept and application. Application is now being made by some Engineering firms.

CHAPTER VII

SUGGESTIONS FOR FUTURE STUDY

The following suggestions for future study would be useful for flash flood forecasting in the high plains and the entire United States.

- Compile an ultra complete falsh flood vs deaths data set with 300 or more points covering 100 years of data.
- Use a multiple regression technique to predict flows at different storm frequencies.
- 3. Use additional data input of miles of roads in flood plains, time of storm and time period since last major rise.
- To investigate the resultant deaths of the same community or one similar in floods of ample warnings versus no warning at all.
- 5. To completely combine the momentum, kinematic and HEC-2 program into one conclusive basin design program.

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APPENDIX A

FLASH FLOOD POTENTIAL SCALES

PERICA CREEK, CLAYTON, NEW MEXICO

TOTAL LOCAL AREA

FLOOD POTENTIAL IS ESTIMATED AT 8 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

1 = NO SIGNIFICANT FLOODING.

2 = SOME STREET AND LOW LAND FLOODING.

- 3 = STREET AND SOME RESIDENTIAL FLOODING.
 - 4 = MAJOR STREET AND RESIDENTIAL FLOODING.
 - 5 = STREAM FLOODING AND FROPERTY LOSS, VERY LIGHT.
 - 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT.
 - 7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50.) AND PROPERTY LOSS. MODERATE.
 - B = STREAM FLOODING WITH LOSS OF LIFE (50 TO 150) AND PROPERTY LOSS, MAJOR.
 - 9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.

DRY CIMARRON: FOLSOM, NEW MEAICO

TOTAL DRAINAGE AREA

FLOOD POTENTIAL IS ESTIMATED AT 8 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

1 = NO SIGNIFICANT FLOODING.

2 = SOME STREET AND LOW LAND FLOODING.

3 = STREET AND SOME RESIDENTIAL FLOODING.

4 = MAJOR STREET AND RESIDENTIAL FLOODING.

5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.

- 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT.
- 7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.
- B = STREAM FLOODING WITH LOSS OF LIFE (50 TO 150) AND PROPERTY LOSS, MAJOR.
- 9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.

RAILROAD CANYON, RATON, NEW MEXICO TOTAL DRAINAGE AREA

FLOOD POTENTIAL IS ESTIMATED AT 8 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

1 = NO SIGNIFICANT FLOODING.

2 = SOME STREET AND LOW LAND FLOODING.

3 = STREET AND SOME RESIDENTIAL FLOODING.

4 = MAJOR STREET AND RESIDENTIAL FLOODING.

5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.

- 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT,
- 7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE,
- 8 = STREAM FLOODING WQTH LOSS OF LIFE (50 TO 150) AND PROPERTY LOSS, MAJOR.
- 9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.

CIMARRON RIVER, SPRINGER, NEW MEXICO 90 SQUARE MILES ABOVE SPRINGER, TOTAL = 2673

FLOOD POTENTIAL IS ESTIMATED AT 8 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

- 1 = NO SIGNIFICANT FLOODING.
- 2 = SOME STREET AND LOW LAND FLOODING.
- 3 = STREET AND SOME RESIDENTIAL FLOODING.
- 4 = MAJOR STREET AND RESIDENTIAL FLOODING.
- 5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.
- 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT.
- 7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.
- B = STREAM FLOODING WITH LOSS OF LIFE (50 TO 150) AND PROPERTY LOSS, MAJOR.
- 9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.

CONCHAS CANAL, TUCUMCARI, NEW MEXICO

TOTAL DRAINAGE' AREA

FLOOD POTENTIAL IS ESTIMATED AT 8 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

- 1 = NO SIGNIFICANT FLOODING.
- 2 = SOME STREET AND LOW LAND FLOODING.
- 3 = STREET AND SOME RESIDENTIAL FLOODING.
- 4 = MAJOR STREET AND RESIDENTIAL FLOODING.
- 5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.
- 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT.
- 7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.
- B = STREAM FLOODING WITH LOSS OF LIFE (50 ℃TO 1505 AND PROPERTY LOSS, MAJOR,
- 9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.
- 10 = CATASTROPHIC FLOOD.

CIMARRON RIVER, UTE PARK, NEW MEXICO

EAGLE NEST TO UTE PARK

FLOOD POTENTIAL IS ESTIMATED AT 8 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

1 = NO SIGNIFICANT FLOODING.

2 = SOME STREET AND LOW LAND FLOODING.

3 = STREET AND SOME RESIDENTIAL FLOODING.

4 = MAJOR STREET AND RESIDENTIAL FLOODING.

- 5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.
- 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT.
- 7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.
- 8 = STREAM FLOODING WITH LOSS OF LIFE² (50 TO 150) AND PROPERTY LOSS, MAJOR.
- 9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.

WOLF CREEK, VALMORA, NEW MEXICO

TOTAL DRAINAGE AREA

FLOOD POTENTIAL IS ESTIMATED AT 8 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

1 = NO SIGNIFICANT FLOODING.

2 = SOME STREET AND LOW LAND FLOODING.

- 3 = STREET AND SOME RESIDENTIAL FLOODING.
- 4 = MAJOR STREET AND RESIDENTIAL FLOODING.

5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.

- 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT.
- 7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.
- B = STREAM FLOODING WITH LOSS OF LIFE (50 TO 150) AND PROPERTY LOSS, MAJOR.
- 9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR,

10 = CATASTROPHIC FLOOD.

MORA RIVER, WATROUS, NEW MEXICO 90 SQUARE MILES ABOVE WATROUS, TOTAL = 670

FLOOD POTENTIAL IS ESTIMATED AT 8 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

- 1 = NO SIGNIFICANT FLOODING.
- 2 = SOME STREET AND LOW LAND FLOODING.
- 3 = STREET AND SOME RESIDENTIAL FLOODING.
- 4 = MAJOR STREET AND RESIDENTIAL FLOODING.
- 5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.
- 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT,
- 7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.
- 8 = STREAN FLOODING WITH LOSS OF LIFE (50 PTO 150) AND PROPERTY LOSS, MAJOR.
- 9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.

NORTH CANADIAN RIVER AT BEAVER, OKLAHOMA

LOCAL AREA, GUYMON TO BEAVER

FLOOD POTENTIAL IS ESTIMATED AT 6 OF A SCALE OF 1 TO 10

1

FLOOD POTENTIAL SCALE IS LISTED BELOW;

1 = NO SIGNIFICANT FLOODING.

2 = SOME STREET AND LOW LAND FLOODING.

3 = STREET AND SOME RESIDENTIAL FLOODING.

4 = MAJOR STREET AND RESIDENTIAL FLOODING.

5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.

6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LICHT.

7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.

B = STREAM FLOODING WITH LOSS OF LIFE (50 TO 150) AND PROPERTY LOSS, MAJOR.

9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) _ AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.

POSSIBLE DEATHS FROM PROBABLE MAX STORM WOULD BE

NORTH CANIDIAN RIVER AT GUYMON, OKLAHOMA LOCAL AREA, STATE HY 95 TO BEAVER

FLOOD POTENTIAL IS ESTIMATED AT 8 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW:

- 1 = NO SIGNIFICANT FLOODING.
- 2 = SOME STREET AND LOW LAND FLOODING.
- 3 = STREET AND SOME RESIDENTIAL FLOODING.
- 4 = MAJOR STREET AND RESIDENTIAL FLOODING.
- 5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.
- 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT.
- 7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.
- 8 = STREAM FLOODING WITH LOSS OF LIFE (50 TO 150) AND PROPERTY LOSS, MAJOR.
- 9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS: MAJOR.
- 10 = CATASTROPHIC FLOOD.

LOCAL CREEK, KENTON, OKLAHOMA

TOTAL DRAINAGE AREA

FLOOD POTENTIAL IS ESTIMATED AT 8 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW:

1 = NO SIGNIFICANT FLOODING.

2 = SOME STREET AND LOW LAND FLOODING.

3 = STREET AND SOME RESIDENTIAL FLOODING.

4 = MAJOR STREET AND RESIDENTIAL FLOODING.

5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.

- 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT.
- 7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.
- B = STREAM FLOODING WITH LOSS OF LIFE (50 TO 150) AND PROPERTY LOSS, MAJOR.
- 9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.
LOCAL CREEK, AMARILLO, TEXAS

FLOOD POTENTIAL IS ESTIMATED AT 6 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

- 1 = NO SIGNIFICANT FLOODING.
- 2 = SOME STREET AND LOW LAND FLOODING.
- 3 = STREET AND SOME RESIDENTIAL FLOODING.
- 4 = MAJOR STREET AND RESIDENTIAL FLOODING.
- 5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.
- 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT.
- 7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.
- 8 = STREAM FLOODING WITH LOSS OF LIFE (50 TO 150) AND PROPERTY LOSS, MAJOR.
- 9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.

POSSIBLE DEATHS FROM PROBABLE MAX STORM WOULD BE / 27

CANADIAN RIVER, CANADIAN, TEXAS BELOW SANFORD RESERVOIR

FLOOD POTENTIAL IS ESTIMATED AT 6 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

1 = NO SIGNIFICANT FLOODING.

2 = SOME STREET AND LOW LAND FLOODING.

3 = STREET AND SOME RESIDENTIAL FLOODING.

4 = MAJOR STREET AND RESIDENTIAL FLOODING.

5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.

6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT.

7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.

8 = STREAM FLOODING WITH LOSS OF LIFE (50 TO 150) AND PROPERTY LOSS, MAJOR.

9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.

POSSIBLE DEATHS FROM PROBABLE MAX STORM WOULD BE

D

PALO DURO CANYON STATE PARK NR. CANYON. TEXAS

FLOOD POTENTIAL IS ESTIMATED AT B OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

1 = NO SIGNIFICANT FLOODING.

2 = SOME STREET AND LOW LAND FLOODING.

3 = STREET AND SOME RESIDENTIAL FLOODING.

4 = MAJOR STREET AND RESIDENTIAL FLOODING.

5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.

- 6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT,
- 7 = STREAM FLOODING WITH PROPABLE,LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE,
- 9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.

POSSIBLE DEATHS FROM PROBABLE MAX STORM WOULD BE 108.

LOCAL CREEK, BOYS RANCH, TEXAS

TOTAL DRAINAGE AREA

FLOOD POTENTIAL IS ESTIMATED AT 8 OF A SCALE OF 1 TO 10

FLOOD POTENTIAL SCALE IS LISTED BELOW;

1 = NO SIGNIFICANT FLOODING.

2 = SOME STREET AND LOW LAND FLOODING.

3 = STREET AND SOME RESIDENTIAL FLOODING.

4 = MAJOR STREET AND RESIDENTIAL FLOODING.

5 = STREAM FLOODING AND PROPERTY LOSS, VERY LIGHT.

6 = STREAM FLOODING WITH CHANCE OF LOSS OF LIFE AND PROPERTY LOSS, LIGHT,

7 = STREAM FLOODING WITH PROPABLE LOSS OF LIFE (0 TO 50) AND PROPERTY LOSS, MODERATE.

8 = STREAM FLOODING WITH LOSS OF LIFE (50 TO 150) AND PROPERTY LOSS, MAJOR.

9 = STREAM FLOODING WITH LOSS OF LIFE (150 TO 400) AND PROPERTY LOSS, MAJOR.

10 = CATASTROPHIC FLOOD.

POSSIBLE DEATHS FROM PROBABLE MAX STORM WOULD BE 64

APPENDIX B

FLASH FLOOD FORECAST SCHEMES

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STATE OF N MEXICO

FORECAST ZONE

PERICA CREEK, CLAYTON, NEW MEXICO

.

TOTAL LOCAL AREA CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 0.7 FT

.

FLOOD STAGE = 2.7 FT. GAGE ZERO = 4980.00 FT.-MSL TIME TO PEAK = 1.6 HOURS

FLASH	FLC	00										
GUIDA	NCE				INCHES	OF R.	AINFALL	: IN 3	HOURS			•
VALUE	S*	0.5	1.0	1.5	2.0	2.5	3.0	4.0	5.0	6.0	8.0	10.0
1.0	I	1.3	2.3	3.9	5.3	6.1	6.9	8.4	9.3	9,6	10.4	11.2
1,2	I	1,2	2.0	3.2	5.0	5,8	6.8	8,1	5.2	9.6	10.4	11.1
1.4	I	1.0	1.7	2.5	3.9	5.3	6.0	7.6	9.0.	9,5	10.3.	11.0
1.6	ï	0.9	1.5	2.2	3.4	5.1	5.8	7.2	8.8	9.4	10.2	10.9
1.8	ĩ	0.7	1.2	1.9	2.8	4.5	5.4	6.9	8.4	9.3	10.1	10.8
2.0	ĩ	0.7	1.1	1.6	2.3	3.7	5.2	6.5	8,0	9.2	10.0	10.7
2.2	ĩ	0.7	0,9	1.4	2.0	3,0	4.5	6.1	7.6	9,0	9.9	10.6
2.4	I	0.7	0.8	1.2	1.7	2.5	3.9	5.8	7.2	8.7	9.8	10.5
2,6	I	0.7	0.7	1.0	1.5.	2.2	3.3	5.4	6,8	8.3	9.7	10.4
2.8	- I	0.7	0.7	0.9	1.4	2.1	3.0	5,4	6.6	8.2	9.6	10.3
3.0	I	0.7.	0.7	0.7	1.1	1.6	2.4	5.1	6.2	7.6	9.5	10.2
3.2	I,	0.7	0.7	0.7	0.9	1.5	2.1	4.5	6.0	7.3	9.4	10.1
3.4	I	0.7	0.7	0.7	0.7	1.2	1.8	3.7	5.6	6.9	9.3	9.9
3.6	1	0.7	0.7	0.7	0.7	1.0	1.6.	3.2	5.4	6.6	9.2	9.8
3.8	I	0.7	0.7	0.7	0.7	0.8	1.3	2.7	5.1	6.3	9.1	9.7
4.0	I	0.7	0.7	0.7	0.7	0.7	. 1.1	2.3	4.8	6.0	8.9	9.5
4.2	I	0.7	0.7	0.7	0.7	0.7	0.9	2.0	4.6	6.1	9.0	9.6
4.4	I	0.7	0,7	0.7	0.7	0.7	0.7	1.7	3.8	5.9	8.9	9.7
4.6	I	0.7	0.7	0.7	0.7	0.7	0.7	1.5	3.3	5.6	8.4	9,5
4.8	I	0.7	0.7	0.7	0.7	0.7	0.7	1.3	3.0	5.5	6.3	9.5
5.0	ĩ	0.7	0.7	0.7	0.7	0.7	0.7	1.2	2.6	5.2	7.9	9.3
* AVA	ILAE	LE FOR	M LOCI	L WEA	THER SE	RVICE	OFFICE					

STATE OF N MEXICO

FORECAST ZONE 1

DRY CIMARRON, FOLSOM, NEW MEDICO

TOTAL DRAINAGE AREA CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 1.1 FT

FLOOD STAGE = 4.6 FT. GAGE ZERO = 6390.00 FT.-MSL TIME TO PEAK = 2.3 HOURS

FLASH	FLO	000		•								
GUIDA	NCE			,	INCHES	S OF R.	AINFALL	L'IN 3	HOURS		/	
VALUE	S*	0.5	1.0	1.5	2.0	2.5	3.0	4.0	5.0	6.0	8.0	10.0
1.0	I	1.5	2.0	2.6	3.3	4.1	4.9	6.5	7.5	8,1	9.1	10.1
1.2	ī	1.4	1.8	2.4	3.0	3.8	4.6	6.2	7.4	8.0	9.0	10.0
1.4	I	1.3	1.,7	2.1	2.6	3.3	4.1	5.7	7.3	7.8	8.9	9,9
1.6	I	1.2	1,6	1.9	2.4	3.1	3.8	5,3	6.9	7.7	8.8	. 9.7
1.8	I	1.2	1.4	1.8	2.2	2.8	3,4	4.9	6.5	7.5	8.7	9.6
2.0	I	1.1	1.4	1.6	2.0	2.5	3.2	4.6	6.1	7.4	8.5	9.5
2.2	I	1.1	1.3	1.5	1.8	2.3	2.8	4.2	5.7	7.3	8.4	9,3
2.4	1	1.1	1.2	1.4	1.7	2.1	2.6	3.8	5.3	6.9	8.3	9.2
2.6	I	1.1	1.1	1.3	1.6	1.9	2.4	3.4	4.8	6.4	8.1	9.0
2.8	I	1.1	1.1	1.3	1.5	1.9	2.3	3.3	4.7	6.3	8.0	8.9
3.0	I	1.1.	1.1	1.2	1.4	1.7	2.0	3.0	4.2	5.7	7.9.	8.7
3.2	I	1.1.	- 1.1	1.1	1.3	1.6	1.9	2,8	4.0	5.4	7.7	8.6
3.4	I	, 1.1	1.1	1.1	1.1	1.4	1.7	2.5	3,7	5.0	7.6	8.4
3.6	I	1.1	1.1	1.1	1.1	1.3	1,6	2.4	3.4	4.7	7.4	8.3
3.8	I	1.1	1,1	1.1	1.1	1.2	1.5	2.2	3.1	4.3	7.3	8.1
4.0	I	1.1	1.1	1.1	1.1	1.1	1.4	2.0	2,9	4.0	7.0	.7.9
4.2	I	1.1	1,1	1.1	1.1	1.1	1.2	1.8	2.8	4.1	7.2	8.D
4.4	I	1.1	1.1	1.1	1.1	1.1	1.1	1.7	2.6	4.0	7.0	8.1
4.6	I	1.1	1.1	1,1	1.1	1.1	1.1	1.6	2.4	3.6	6.5	7.8
4.8	I	1.1	1.1	1,1	1.1	1.1	1.1	1.5	2.3	3,5	6.5	7.8
5.0	I	1.1	1.1	1.1	1.1	1.1	1.1	1.4	2.1	3.1	6.0	7.6
* AVA	TLAE	BLE FO	RM LOCA	AL WEA	THER SE	RVICE	OFFIC	F				

STATE OF N MEXICO

FORECAST ZONE 1

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RAILROAD CANYON, RATON: NEW MEXICO

TOTAL DRAINAGE AREA CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 0.7 FT

				•				FLDÓ GAGE TIME	D STAGE ZERD = TO PEA	= : 6780, K_=	2, 00 FT 1,7 H	9 FT. MSL OURS
FLASH	FLO	00		•	THEFE	05 0		· • • • •	100000			· •
GUIDAN	ICE		- 0		INCHES	OF K	AINFALL	· 1N 3	HOURS		• •	
VALUES	; * .	0.5	1,0	1.5	2.0	2.5	3.0	4.0	5.0	6.0	8.0	10.0
1.0	r	1.2	1.9	2.8	4.1	5.4	. 5.9	7.0	8.0	8.9	9.9	10.3
1.2	1 T	1.1	1.7	2.4	3.5	5.1	5.7	6.8	7.8	8.8	9.8	10.3
1.4	1	1.0	1.5	2.1	2.8	4.1	5.4	6.4	7.5	8.5	. 9.7	10.2
1.6	Ŷ	0.9	1.3	1.9	2.5	3.7	5.1	6.2	7.2	8.3	9.7	18.2
1.8	Ť	0.A	1.1	1.5	2.2	3.1	4,4	5.9	7.0	8.0	9.6	10.1
2.0	ŝ	0.7	1.0	1.4	1.9	.2.7	3.8	5.7	6.7	7.8	9 6	10.0
2.2	T	0.7	0.9	1.3	1.7	2.3	3.1	5.5	6.4	7.5	9.5	9.9
2.4	Ť	0.7	0.8	1.1	1.5	2.1	2.8	5.1	6.2	7.2	9.3	9.9
2.6	Ť	0.7	0.7	1.0	1.3	1.8	2.5	4.4	5.9	6.9	9.0	9,8
2.8	Ť	0.7	0.7	0.9	1.2	1.8	2.4	4.2	5.8	6.8	8.9	9.8
3.0	Ŷ	0.7	0.7	0.7	5.0	1.4	2.0	3.6	5.5	6.4	5.6	9.7
3.2	Ť	0.7	0.7	0.7	0.9	1.3	1.8	3.1	5.3	6.2	8.4	9.6
3.4	Ŷ-	0.7	0.7	0.7	0.7	1.1	1.6	2.7	4.8	5.0	8.1	9.5
3.6	Ť	0.7	0.7	0.7	0.7	1.0	1.4	2.4	4.3	5.8	7.8	9.3
3.8	Ŧ	0.7	0.7	0.7	0.7	0.8	1.2	2.2	3.8	5.6	7.5	9.0
4.0	Ť	0.7	0.7	0.7	0.7	0.7	1.1	2.0	3.3.	5.4	7.3	8.7
4.2	T	0.7	0,7	0.7	0.7	0.7	0.9	1.7	3.2	5.4	7,4	8.9
4.4	Ť	0.7	0.7	0.7	0.7	0.7	0.7	1.5	2.8	5.3	7.3	9.0
4.6	Ť	0.7	0.7	0.7	0.7	0.7	0.7	1.3	2.5	4.8	7.0	8.5
4.8	Ŧ	0.7	0.7	0.7	0.7	0.7	0.7	1.2	.2.3	4.6	6.9	E.6
5.0	Ť	0.7	0.7	0.7	0.7	0.7	0.7	1.1	2.1	3.8	6.6	· 8.1
* AVAT	Î AB	LE FOR	MLOC	AL WEAT	THER SE	RVICE	OFFICE					

STATE DE N MEXICO

FORECAST ZONE 1

CIMARRON RIVER. SPRINGER. NEW MEXICO

90 SOUARE MILES ABOVE SPRINGER. TOTAL = 2673 CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 1.4 FT

FLOOD	57	AGE	=		5.6	FT.
GAGE	ZER	0 =	5770.	00	FT	MSL
TIME	10	PEAN	(=	1.9	HOL	JRS

FLASH		000			INCHES	OF RI	ATNEAL	1 · TN 3	HOURS			
VALUE	S*	0.5	1.0	1.5	2.0	2.5	3.0	4.0	5.0	6.0	8.0	10.0
1.0	I	2.2	3,3	4.7	6.2	7.8	8.8	10.0	11.2	12.4	13,9	14.3
1.2	I	2,0	3.0	4.1	5.5	7.1	8.5	9.8	11.0	12.2	13.9	14.3
1.4	I	1.B	2.6	3.6	4.7	6.2	7.6	9,3	10.6	11.9	13.8	14.2
1.6	I	1.6	2.4	3.2	4.2	5.7	7.1	9.0	10.3	11.6	13.7	14.2
1.8	I	1.5	2.1	2.8	3.8	5.1	6.5	8.8	10.0	11.2	13,4	14.1
2.0	1	1.4	1.9	2.5	3.3	4.5	5.9	8.5	9.7	10.9	13.2	14.1
2.2	ĭ	1.4	1.7	2.3	3.0	3.9	5.1	7.9	9,3	10.6	13.0	14.0
2.4	1	1.4	1.6	2.0	2.7	3.6	4.6	7.2	9.1	10.3	12.7	13.9
2.6	I	1.4	1.4	1.8	2.3	3.2	4.2	6.5	8.7	9.9	12.5	13.9
2.8	ĩ	1.4	1.4	1.7	2.2	3.0	4.0	6.3	8.6	9.8	12.4	13.8
3.0	I	1.4	1.4	1.4	1.9	2.6	3.4	5.6	8.0	9.3	12.0	13.6
3.2	I	1.4 -	1.4	1.4	1.7	2.3	3.1	5.1	7,5	9.1	11.7	13.3
3.4	I	1.4	1.4	1.4	1.4	2.0	2.7	4.5	6.9	8.8	11.3	13.0
3.6	I	1.4	1.4	1.4	1.4	1.8	2.4	4.1	6.3	8.6	11.0	12.8
3.8	I	1.4	1.4	1.4	1.4	1.6	2.2	3.7	5.8	8.2	10.7	12.5.
4.0	I	1.4	1.4	1.4	1.4	1.4	2.0	3.4	5.3	7.6	10.4	12.1
4.2	I	1.4	1.4	1.4	1.4	1.4	1.6	2.9	5.2	7.8	10.5	12.4
4.4	1	1.4	1.4	1.4	1.4	1.4	1.4	2.6	4.6	7.5	10.4	12.4
4.6	I	1.4	1.4	1.4	1.4	1.4	1.4	2.4	4.2	6.8	10.0.	11.9
4.8	I	1.4	1.4	1.4	1.4	2.4	1.4	2.2	3.9	6.7	9,9	11.9
5.0	I	1.4	1.4	1.4	1.4	1.4	1.4	2.0	3.6	5.9	9.6	11.4
* AVA	Y1 01	NE FOR	M LOC.	AL NEA	THER SE	RVICE	OFEIC	7				

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STATE OF N MEXICO

FORECAST. ZONE 2

CONCHAS CANAL. TUCUMCARI. NEW MEXICO

TOTAL DRAINAGE AREA - CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 0.7 FT

FLOOD STAGE = 2.9 FT. GAGE ZERO = 4100.00 FT.-MSL TIME TO PEAK = 2.3 HOURS

FLASH	FLASH FLOOD													
GUIDA	NCE				AINFAL	LIN 3	HOURS							
VALUE	S *	0.5	1.0	1.5	2.0	2.5	3.0	4.0	5.0	6.0	8.0	10.0		
3 . D	T	1.8	3.5	5.6	6.7	8.0	9.2	10.0	10.7	11.3	-12.2	12.3		
1.2	Ŷ	1.6	2.8	5.1	6.2	7.5	8.8	9.9	50.6	11.2	12.2	12.3		
1.4	Ť	1.2	2.3	4.0	5.6	6.7	7.9	9.7	10.4	11.0	12.1	12.3		
1.6	Ť	1.0	2.0	3.2	5.3	6.4	9.5	9.6	10.2	10.9	12 1	12.3		
1 0	Ť	0.8	1.6	2.6	4.5	5.9	7 0	9 2	10.0	10.7	12 0	12 2		
2.00	T	0.7	1 4	2.2	7.5	5 5	2 5	ו C 8 7	4 Q	10.5	31 0	12 2		
2.0	r r	0.7	4 1	1 0	2.0 2.8	1.8	5.J	1,20	97	1010 10 L	15 7	10 0		
2 4	т. Т	07	.0.9	1.6	2.U	4.0	5.6	7.5	9.6	10.2	11 5	12 2		
	i T	0.7	0.7	1 2	2.1	2.0	5.0	7 0	. 9 1	10.0	14 2	10 0		
2.0	1	0.7	0.7	1+6	C+U	2.2	0 e C 1 - 5	7.0	2.1	10.0	44 2	10 0		
<	1	0.7	0.7	7.1	1.0	~+7 0 7	7.0	5.0	0,7	2.7	77*2	16.6		
3.0	I	0.7	0.7	0.8	2.4	2.0	5.6	5.5	8.2	9.1	11.1	12.1		
3.2	ĭ	0.7	0./	0.7	1.1	2.0	5.9		7.8	9.6	10.9	11.9		
3.4	I	0.7	0.7	0.7	0.7	1.6	2.5	5.5	7.3	9.3	10,7	11.7		
3.6	I	0.7	0.7	0.7	0.7	1.2	2.1	5.1	6.8	8.8	10.5	11.5		
3.8	I	0.7	0.7	0.7	0.7	1.0	1.8	4.3	6.4	8.4	10.4	11.3		
4.0	I	0.7	0.7	0.7	0.7	0.7	1.5	3.5	6.1	7.9	10.2	11.1		
4.2	1	0.7	0.7	0.7	0.7	0.7	1.0	2.7	6.0	8.0	10.3	11.3		
4.4	I	0.7	0.7	0.7	0.7	0.7	0.7	2.4	5.6	7.8	10.3	11.3		
4.6	Ţ	0.7	0.7	0.7	0.7	0.7	0.7	2.0	5.3	7.2	10.1	11.0		
4.8	ï	0.7	0.7	0.7	0.7	0.7	0.7	1.8	4.8	7.1	10.0	11.0		
5.0	ī	0.7	0,7	0.7	0.7	0.7	0.7	1.5	4.1	5.5	9.8	10.8		
* AVA	ILA	BLE FO	RM-LOCA	AL WEAT	THER SE	RVICE	OFF1C	Ε.						

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STATE OF N MEXICO

FORECAST ZONE 2

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CIHARRON RIVER. UTE PARK. NEW MEXICO

EAGLE NEST TO UTE PARK CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 1.0 FT

FLOOD STAGE = 4.1 FT. GAGE ZERO = 7405.00 FT.-MSL TIME TO PEAK = 1.6 HOURS

FLASH	FLC	OD		:	****			· •*				
VALUE	NCE S*	0.5	1.0	1.5	2.0	2.5	JINFALL 3.0	4.0	5.0	6.0	8.0	10.0
1.0	Ţ	1.4	2.1	2.8	3.6	4.6	5.7	7.1	7.7	8.4	9.8	10.9
1.2	ī	1.4	1.9	2.5	3.2	4.2	5.3	6.9	7.6	8.3	9.7	10.8
1.4	ī	1.2	1.7	2.2	2.8	3.6	4.6	6.6	7.4	8.1	9.5	10.6
1.6	ñ	1.1	1.5	2.0	2.6	3.3	4.2	5.1	7.2	7.9	9.4	10.5
1.8	ĩ	1.1	1.4	1.8	2.3	3.0	3.8	5.7	7.0	7.8	9.2	10.4
2.0	ĩ	1.0	1.3	1.6	2.1	2.7	3.5	5.2	6.9	7.6	9.0	10.2
2.2	I	1.0	1.2	1.5	1.9	2.4	3,0	4.7	6.7	7.4	8.8	10.0
2.4	r	1.0	1.1	1.4	1.7	2.2	2.8	4.3	6.2	7.2	8.7	9.9
2.6	I	1.0	1.0	1.2	1.5.	2.0	2.5	3.8	5.6	7.0	8.5	9.6
2.8	I	1.0	1.0	1.2	1.5	1.9	2.4	3.7	5.4	6.9	8.4	9.6
3.0	Ī.	1.0	1.0	1.1	1.3	1.6	2.1	3.3	4.B	6.7	8.1	9.3
3.2	- I	1.0	1.0	1.0	1.2	1.5	1.9	3.0	4.5	6.2	8.0	9.1
3.4	1	1.0	1.0	1.0	1.0	1.4	1.7	2.7	4.1	5.8	7,8	8.9
3.6	I	1.0	1.0	1.0	1.0	1.2	1.6	2.5	3.7	5.3	7.6	8.7
3.8	1	1.0	1.0	1.0	1.0	1.1	1.4	2.3	3.4	4.9	7.4	8.5
4.0	I	1.0	1.0	1.0	1.0	1.0	1.3	2.1	3.1	4.5	7.3	8.2
4.2	I	1.0	1.0	1.0	1.0	1.0	1.1	1.8	3.0	4.7	7.3	8.4
4.4	I	1.0	1.0	1.0	.1.0	1.0	1.0	1.7	2.7	4.5	7.3	8.4
4.6	I	1.0	1.0	1.0	1.0	1.0	1.0	1.5	2.5	4.0	7.1	8.1
4.8	I	1.0	1.0	1.0	1.0	1.0	1.0	1.4	2.4	3.9	7.0	8.1
5.0	I	1.0	1.0	1.0	1.0	1.0	1.0	1.3	2.2	3.4	6.8	7.8
* AVA	ILAB	LE FOR	M LOC.	AL WEAT	THER SE	RVICE	OFFICE					

STATE OF N MEXICO

FORECAST ZONE 1

WOLF CREEK, VALMORA, NEW MEXICO

TOTAL DRAINAGE AREA CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 1.4 FT

FLOOD STAGE = 5.8 FT. GAGE ZERO = 6315.00 FT.-MSL TIME TO PEAK = 3.0 HOURS

FLASH	FLO	DD			• •	•						
GUIDA	NCE				INCHES	OF R	AINFALL	: IN 3	HOURS			
VALUE	S*	0.5	1.0	1.5	2.0	2.5	3.0	4.0	5.0	6.0	8.0	10.0
٥. ٢	T	1.9	2.6	3.4	4.3	5.4	6.3	A. 1	9.0	9.7	10 9	12.1
1.2	Ť	1.8	2.4	3.0	3.9	4.9	6.0	7.8	8.9	9.5	10.2	12 0
1.4	T	1.7	2.1	2.7	3.4	4.3	5.3	7.1	B.7	9.4	10.6	11.8
1.6	Ť	1.6	2.0	2.5	3.1	4.0	4 9	6.7	8.6	9.2	10.5	11 7
1 8	Ť	1 5	1 8	2 3	2 8	3 6	4 5	63	8 1	D 0	10 0	11 5
2.0	T	1.4	1.7	2.1	2.6	3.3	. 1 1	6.0	7.7	8.9	10.2	11 X
2.2	т Т	1.4	1.6	1.9	2.3.	2.9	3.6	5 1	7.2	8.7	10.2	11 2
2.4	T .	1 4	5 5	1.8	2.2	2.7	र २ २ २	5.0	6 . R	8.5	-9 9	11.2
2 6	Ϋ́	1 4 5	1 4	1.7	2.0	25	x 1	4 5	6.2	8 0	97	10 8
2.8	т Т	1 4	3 4	1 6	1 9	2.0	3,1	- 	6.2	7 8	9 6	10.7
2.0	T T	1 4	5 4	1 6	1.7	2 1	3.6	7.9	55	7 3	2.0	10.5
3.0 x 2	*	1.1	1 a 7		1.1	2 0	2.0	.3.7	5.2	1.2	7.T	10.7
3.2		1.1	1 - 7	1.4	1.6	.2.0	2.4	- J. D	5.2	6.0	5.5	10.5
2.4	1	1.44	1.7	1.44	1.4	1.0	2.2	3.5	4.0	D • 4	2+1	10.1
3.6	1	1.4	1.4	1.4	1.4	1.1	2.0	3.0	4.4	6.1	8.9	3.9
3.8	I	1.4	1.4	1.4	1.4	1.6	1.9	2.8	4.0	5+7	8,7	. 9.7
4.0	I	1.4	1.4	1.4	1.4	1.4	1.8	2.6	3.8	5.2	8.6	9.5
4.2	I	1.4	1.4	1.4	1.4	1.4	1.6	2.3	3.6	5.4	8.6	9.6
4.4	I	1.4	1.4	1.4	1.4	1.4	1.4	2.2	3.3	5.2	8.6	9.7
4.6	I	1.4	1.4	1.4	1.4	1.4	1.4	2.0	3.1	4.7	8.1	9.4
4.8	I	1.4	1.4	1.4	1.4	1.4	1.4	1.9	2.9	4.6	8.1	9.4
5.0	I	1.4	1.4	1.4	1.4	1.4	1.4	1.8	2.7	4.1	7.5	9.1
* AVA	ILAB	LE FOR	H LOCA	L WEA	THER SE	RVICE	OFFICE				•	

STATE OF N MEXICO

FORECAST ZONE 1

MORA RIVER. WATROUS. NEW MEXICO

- 90 SQUARE MILES ABOVE WATROUS, TOTAL = 670 CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 1.4 FT

FLOOD) ST	AGE.			5.5 FT.
GAGE	ZEF	:0 =	6450.	00	FTMSL
TIME	TO	PEAN	(= _	3.2	HOURS

FLASH	FLO	000		•		•					•	-
GUIDA	NCE				INCHES	OF R	AINFALL	: IN 3	HOURS			
VALUE	S*	0.5	1.0	1.5	2.0	2.5	3.0	4.0	5.0	6.0	8.0	10.0
1.0	т	1.9	2.7	3.7	4.7	6.D	7.1	8.6	9.4	10.2	11.8	13.0
1.2	Ť	1.8	2.5	3.3	4.3	5.5	6.6	8.5	9.3	10.1	11.7	12.9
3 4	Ť	1 6	2.2	2.9	3.7	4. A	5.9	8.1	9.0	9.9	11 5	12.7
5 6	T	1.5	2 0	2.6	3.4	4 4	5.5	7.6	8.8	9.7	11 3	52 6
3 - C	T T	1 1	1.8	2 4	3.0	u n	5.0	7 1	86	9 5	77.5	10 5
2.0	¥	5 11	1 7	2 1	5.0	7.U 7.L	5.0	6 6	ន ដ		10 0	10 2
2.00	T T	1 4	1 6	5.0	2 4 7	3.0	4°.5	6 0	г «О	9.0	10.7	12.1
5 h	Ť	5 H	1 5	1 0	2 3	2 0	2 6	5 6	7 6	9,0 9,0	10.5	11 0
217	1	5 0	1.0	1.0	2.1	2.1	7 2	5.0	7 0	0.0	10.1	1107
2.00	1	4 1	1.1	5 6	2.0	2:0	3.5	J. D	1.0	0.0	10.0	11.1
2.0	1	- 1 - H	1	4.0	1	2.0		***0	0.0	0.0	10.2	11.0
3.U 7 0	1	1.4	1,4	1.4	2	2.2	2.0	4.J	6.1	0.1	2.2	11.3
5.2	1	1.4	1.4	1.44	1.6	2.0	2.5	4.0	5.8	· · · · ·	9.1	11.1
3.4	I	1.4	1.4	1.4	1.4	1.8	2.3	3.6	5,3	7.2	9.5	10.8
3.6	I	1.4	1.4	1.4	1.4	1.6	2.1	3.3	4.9	6.7	9.3	10.5
3.8	I	1.4	1.4	1.4	1.4	1.5	1.9	3.0	4.5	6.3	9.1	10.3
4.0	1 .	1.4	1.4	1.4	1.4	1.4	1.8	2.7	4.1	5.8	. 8.9	10.0
4.2	I	1.4	1.4	1.4	1.4	1.4	1.5	2,4	4.0	6.0	8.9	10.2
4.4	I	1.4	2.4	1.4	1.4	1.4	1.4	2.2	3.6	5.8	8.9	10.3
4.6	I	1.4	1.4	1.4	1.4	1.4	1.4	2.0	3.3	5.3	8.6	9.9
4.8	Ţ	1.4	1.4	1.4	1.4	1.4	1.4	1.9	3.1	5.2	8.6	9.9
5.0	ī	- 1.4	1.4	1.4	1.4	1.4	1.4	1.8	2.9	4.5	8.4	9.5
AVA *	TLAF	BLE FOR	N LOCA	AL WEA	THER SE	RVICE	OFFICE		• • •		- • •	

AVAILABL

147

STATE OF OKLAHOMA

FORECAST ZONE 20

NORTH CANADIAN RIVER AT BEAVER. OKLAHOMA

LOCAL AREA, GUYMON TO BEAVER CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 11.3 FT

.

					•		• ,	FLOD GAGE TIME	D STAG ZERO TO PE	E = = 236B AK =	15. .00 FT 8.9 H	0 FT. MSL DURS
FLASH	FLC	DOD										
GUIDAN	СE				INCHE	S OF RA	AINFAL	L: IN 3	HOURS			
VALUES	*	0.5	1.0	1.5	2.0	2.5	3.0	4.0	5.0	6.0	8,0	10.0
1.0	۲	11.8	12.5	13.3	14.3	15.2	15.6	16.6	17.5	18.3	19.2	20.0
1.2	Î	11.7	12.3	13.0	13.9	15.0	15.5	16.4	17.3	18.2	19.1	19.9
3.4	Ŧ	11.5	12.0	12.6	13.3.	14.3	15.2	16.1	17.0	18.0	19.0	19.8
1.6	ĩ	11.4	11.9	12.4	13.1	34.0	15.0	15.9	16.8	17.7	18.9	19.7
1.8	Ť	11.3	11.7	12.2	12.8	13.6	14.5	15.7	16.6	17.5	18.8	19.5
2.0	Ť	11.3	11.6	12.0	12.5	13.2	14.1	15.5	16.3	17.3	18.7	19.5
2.2	Ť	11.3	11.5	11.8	12.2	12.9	13.6	15.2	16.1	17.0	18.6	19.4
2.4	Ŷ	11.3	11.4	11.7	12.1	12.6	13.3	15.0	15.9	16.8	18.5	19.2
2.6	÷ r	11.3	11.3	11.5	11.8	12.4	13.0	14.5	15.6	16.5	18.4	19.1
2.8	Ť	11.3	11.3	11.5	11.8	12.3	12.9	14.4	15.5	16.4	18.3	19.1
3.0	Ŧ	11 3	17 3	11.3	11.6	12.0	12 5	13 9	15 3	16.1	18 0	18 9
3 2	T	11 3	11.3	11.3	11.0	11 8	12 3	13.6	15 1	15 9	17 R	18.8
<u>ү.</u> г	T.	31 3	11.3	11.3	15.3	11.7	121	13.0	14 8	15.7	17.5	18 4
3.4	1 4	11 7	39 3	11 7	11 3	11 5	1, 9	13.0	14.0	15.5	17 2	18 5
3 8 ·	Ť	11 3	11 3	11.3	11.3	11.0	11 2	12.7	14 0	15.3	17 1	10.0
5.0 L 0	ч ·	11 7	11 3	11.00 11 X		24.17 79 X	11.0	12 5	12 7	10.0	16 R	18 0
1.U U 2	7	11.5	11.0	11 7	11.3	11.7	11.0	12.2	13.1	15 2	16.0	10.2
7.C	1. •	11.0	11.U	11 7	49 7	11 X	11.7	10 0	12.0	1202	10.2	10.0
1.1	Å.	11.0	11.0	11.0	11.5	11.0	11.3	12.0	13,3	10.1	10.7	10.0
7.6	1	11.3	77.0	11 1	2 . T . C . T L	11.0	11.5	11.2	12.0	14.0	10.0	10.0
7.0 E 0	* T	11.3	11.0	11.2	11.5	77.0	14.0	11.0	107	14.1	10.0	10.0
		11.0	71.0	11.3	11.2	CONJOL TT°2	11.3	11.6	16.1	14.1	10.2	11.5
₹ AVAl	LAt	sil ru	nn LUL	AL WLA	INCK S	LKVILL	したたまし	- <u>E.</u>				

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STATE OF OKLAHOMA

FORECAST ZONE 21

NORTH CANIDIAN RIVER AT GUYMON, OKLAHOMA

LOCAL AREA, STATE HY 95 TO BEAVER CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 1.7 FT

FLOOD STAGE = 6.9 FT. . GAGE ZERO = 2970.00 FT.-MSL TIME TO PEAK = 4.4 HOURS

GUIDA	NCE				INCHES	OF R	AINFALL	IN 3	HOURS			
VALUE	S*	0.5	1.0	1.5	2.0	2.5	3.0	4.0	5.0	6.0	8.0	10,0
1.0	1	2.1	2.7	3.3	4.1	5.0	5,9	7.6	8.9	10.0	12.1	12.0
1.2	ĩ	2.0	2.5	3.1	3.7	4.6	5,6	7.3	8.7	9.9	11.0	11.9
1.4	l	1.9	2.3	2.8	3.3	4.1	4.9	6.8	8.2	9.6	10.8	11.8
1.6	1	1.8	2,2	2.6	3.1	3.8	4.6	6.4	7.5	9.3	10.7	11.7
1.8	1	1.8	2.1	2.4	2.9	3.6	4.3	5.9	7.6	9.0	10.6	11.5
2.0	I	1.7	2.0	2.3	2.7	3.3	3.9	5.5	7.2	8.6	10.5	11.4
5.5	r	1.7	1.9	2.1	2.5	3.0	3.6	5.1	6.8	8.2	10.3	21.3
2,4	I	1.7	1.8	2.0	2.3	2.8	3.3	4.7	6.4	7.9	10.2	11.1
2.6	I	1.7	1.7	1.9	2.2	2.6	3.1	42 3	5.8	.7.5	10.1	D1.0
2.8	ĩ	1.7 ~	1.7	1.9	2.1	2.5	3.0	4.2	5.7	7.3	10.0	10.9
3.0	I	1.7	1.7	1.8	2.0	2.3	2.7	3.8	5.2	6.8	9.8	10.7
3.2	I	1.7	1.7	1.7	1,9	2.2	2.5	3.6	4.9	6.4	9.4	10.6
3.4	I	1.7	1.7	1.7	1.7	2.0	2.4	3.3	4.5	6.0	9.0	10.4
3.6	I	1.7	1.7	1.7	1.7	1.9	2.2	3.1	4.2	5.6	8.7	10.2
3.8	I	1.7	1.7	1.7	1.7	1.8	2.1	2.9	3.9	5.3	8.3	10.1
4.0	I	. 1.7	1.7	1.7	1.7	1.7	2.0	2.7	3.7	4.9	8.0	9.9
4.2	1	1.7	1.7	1.7	1.7	1.7	1.8	2,5	3.6	5.0	8.1	10.0
•4.4	I	1.7	1.7	1.7	1.7	1.7	1.7	2.3	3.3	4.8	8.0	10.0
4.6	ĩ	1.7	1.7	1.7	1.7	1.7	1.7	2.2	3.1	4.5	7.6	9.7
4.8	I	1.7	1.7	1.7	1.7	1.7	1.7	2.1	3.0	4.4	7.5	9.7
5.0	- I	1.7	1.7	1.7	1.7	1.7	1.7	2.0	2.8	3.9	7.1	9.1
* AVA	TIAP	UE FOR	RM LOCI	M WEA	THER SE	RVICE	DEETCE					

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STATE OF OKLAHOMA

FORECAST ZONE 21

LOCAL CREEK, KENTON, OKLAHOMA

TOTAL DRAINAGE AREA CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 1.9 FT

FLOOD STAGE = 3.8 FT. GAGE ZERO = 4290.00 FT.-MSL TIME TO PEAK = 1.9 HOURS

FLASH	FLO	00										
GUIDA	NCE	•		• •	INCHES	OF R	AINFALL	_ IN 3	HOURS			
VALUE	S*	0.5	1.0	1.5	2.0	2.5	3.0	4.0	5.0	6.0	8.0	10.0
1.0	٣	3.0	4.6	6.5	7.3	8.3	9.3	10.8	11.3	11.8	12.7	13.5
1.2	ĵ	2.8	4.1	6.0	6.9	7.9	8.9	10.8	11.2	11.7	12.7	13.5
3.4	Ŷ	24	35	5.1	6.5	7.3	8 2	10 1	11.1	. 11.6	12 5	17.1
5 6	л Т	5 5	3.0	0 5	< 1 < 1	7 0	7 9	- 0 - 7	11.1	35 0	10 0	17 7
1.0	Ť	2 6	2.2	7 . J	5 1	674	75		10 6	11 7	10 2	12.0
200.	л. Т	5 0	2.0	2.2	5.4	6 1	7.5	5.5	10.0	44 0	10 0	17.2
2.0	۲. ۲	1+2	2.0	2.4	4.0	5 7	1.1	0.7	20.2	37 9	12.2	10.1
6. a C	1	1.7	2 3	2.1	4.1	5.1	6.7	7 0	10.2	11.1	12+1	12.7
6.44	1	1.7	2,1	2.0	5.6	D+1	5.4	1.9	7.1	11.0	12.0	75.8
2.6	1	1.9	1.9	2.4	3.2	4.4	6.0	1.5	9.2	10.8	11.8	12.6
2.8	I	1.9	1.9	2.3	3.0	4.2	5.7	7.4	9.0	10.8	11.8	12.6
3.0	I	1.9	1.9	1.9	2.6	3,5	4.8	7.0	8.5	10.2	11.6	12.4
3.2	r	1.9	1.9	1.9	2.3	3.2	4.2	6.7	8.1	9.8	11.5	12.3
3.4	I	1.9	1.9	1.9	1.9	2.8	3.7	6.4	7.7	9.4	.11.3	12.1
3.6	Ĩ.	1.9	1.9	1.9	1.9	2.4	3.3	6.0	7.4	9.0	11.2	12.0
3.8	I	1.9	1.9	1.9	1.9	2.2	3.0	5.3	7.1	8.6	11.1	11.8
4.0	I	1.9	1.9	1.9	1.9	1.9	2.7	4.7	6.8	8.2	11.0	11.7
4.2	1	1.9	1.9	1.9	1.9	1.9	2.2	4.0	6.7	8.3	11.0	11.8
4.4	T	1.9	1.9	1.9	1.9	1.9	1.9	3.6	6.4	8.1	11.0	11.8
4.6	Ţ	1.9	1.9	1.9	1.9	1.9	1.9	3.2	6.1	7.7	10.8	11.6
4.8	т.	1.9	1.9	1.9	1.9	1.9	1.9	3.0	5.7	7.6	10.8	11.6
5.0	Ť	1.9	1.9	1.9	1.9	1.9	1.9	2.7	5.1	7.1	10.5	11.4
* AVA	TIAR	FFOR	MIOCA	WEA	THER SE	RVICE	DEEIC	-				

STATE OF TEXAS

FORECAST ZONE 1

CANADIAN RIVER: CANADIAN: TEXAS

BELOW SANFORD RESERVOIR CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 8.5 FT

FLOOD STAGE = 17.1 FT. GAGE ZERO = 701.50 FT.-MSL TIME TO PEAK = 7.0 HOURS

FLACH	51.01	on		· ·				•	••• •••			00110
GUIDANCE INCHES OF RAINFALL: IN 3 HOURS												
VALUES*		0.5	1.0	1.5	2.0	2.5	3.0	4.0	5.0	6.0	8.0	10.0
			· · · · · ·					a,		~ ~ ~ ~ ~ ~ ~		
1.0	1	8.9	9.4	10.0	10.7	11.5	12.4	14.1	15.8	17.2	18.3	19.4
1.2	I	8.8	9,2	9.8	10.4	11.2	12.0	13.8	15.5	17.1	18.3.	19.3
1.4	1	8.7	9.1	9.5	10.0	10.7	11.5	13.2	14.9	16.7	18.1	19.1
1.6	1	8.6	9.0	9.3	9.8	10.5	11.2	12.8	14.5	16.3	18.0	19.0
1.8	I	8.6	8.8	9.2	9.6	10.2	10,9	12.4	14.1.	15.9	17.9	18.9
2.0	I	8.5	.8.8	9.0	9.4	9.9	10.6	12.0	13.7	15.4	17.7	18.7
2.2	I	8.5	8.7	8.9	9.2	9.7	10.2	11.6	13.2	14.9	17.6	18.6
2.4	I	8.5	8.6	8.8	9.1	9.5	10.0	11.2	12.8	14.5	17.4	18.4
2.6	I	8.5	8.5	8.7	9.0	9.3	9.8	10.9	12.3	14.0	17.3	18.2
2.8	I.	8.5.	8.53	8.7	8.9	°9,3	9.7-	10.8	12.1	13.8	17.2	18.2
3.0	I	8.5	8.5	8.6	8.8	9.0	9.4	10.4	11.7	13.2	16.9	17.9
3.2	I	8.5	8.5	8.5	8.7	9.0	9.3	10.2	11.4	12.8	16.5	17.8
3.4	I	8.5	8.5	8.5	8.5	8.8	9.1	.9.9	11.1	12.4	15.9	17.6
3.6	I	8.5	8.5	8.5	8.5	8.7	9.0	9,8	10.8	12.1	15.5	17.4
3.8	I	8.5	8.5	8.5	8.5	8.6	8.9	9.6	10.5	11.8	.15.0	17.3
4.0	Ι.	8.5	8.5	8.5	8.5	8.5	8.8	9.4	10.3	11.4	14.6	17.1
4.2	I	8.5	8.5	8.5	8.5	8.5	8.6	9.2	10.2	11.5	14.8	17.2
4.4	I	8.5	8.5	8.5	8.5	8.5	6.5	9.1	10,0	11.4	14.6	17.2
4.6	ĩ	8.5	8.5	8.5	8.5	.8.5	8.5	.9.0	9.8	11.9	14.1	16.8
4.8	I	8.5	8.5	8,5	8.5	8.5	8.5	8.9	9.7	11.0	14.0	16.8
5.0	I	8.5	8.5	8.5	8.5	8.5	8.5	8.8	9.5	10.5	13.5	16.0
* AVA	ILABI	LE FOR	M LOC	AL WEA	THER S	ERVICE	OFFIC	E	-			

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STATE OF TEXAS

FORECAST ZONE 4

PALO DURO CANYON STATE PARK NR. CANYON, TEXAS

TOTAL DRAINAGE BASIN ABOVE GAGE CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 0.7 FT

FLOOD STAGE = 3.0 FT. EAGE ZERO = 2780.00 FT.-MSL TIME TO PEAK = 2.8 HOURS

								-				
FLASH		000						· • •				
GUIDA	INCE				INCHES	OF R	AINFALL	16 2	HUURS			
VALUE	S*	0.5	1.0	1.5	2.0	2.5	3.0	4.0	5.0	6.0	8,0	10.0
	_						ی دست می مواند موجد ده. دم					
1.0	I	0.9	1.4	1.9	2.6	3.5	4.5	6.6	8.4	9.1	10.6	11.8
1.2	1	8.0	1.2	1.7	2.3	3.2.	4.1	6.2	8.2	9.0	10.4	11.7
1.4	I	0.7	1.1	1.5	2.1	2.9	3.7	5.8	7.7	8.8	10.3	11.5
1.6	1	0.7	1.0	1.4	1.9	2.6	3.4	5.4	7.4	8.7	10.2	11.4
1.8	I	0.7	0,9	1.2	1.6	2.2	2.9	4.7	6.7	8.4	9.9	11.1
2,0	I	0.7	0.8	1.0	1.4	1.9	2.6	4.1	6.1	8.1	9.7	10.9
2.2	I	0.7	0.7	0.9	1.3	1.7	2.3	3.9	5.9	7.9	9.6	10.8
2.4	I	0.7	0.7	0.7	1.0	1.5	2.1	3.7	5.9	7.8	9.5	10.8
2.6	I	0.7	0.7	0.7	1.0	1.4	1.9	3.2	4.9	6.7	9.2	10.4
2.8	I.	0.7	0.7	0.7	0 . B	1.2	1.6	2.8	4.4	6.2	9.0	10.1
3.0	I	0.7	0.7	0.7	0.7	1.0	1.4	2.5	4.0	5.8	8,8	9.9
3.2	I	0.7	0.7	0.7	0.7	0.8	1.2	1.6	3.6	5.4	8.6	9.7
3.4	I	0.7	0.7	0.7	0.7	0.7	1.1	2.0	3.3	5.0	8.4	9.5
3,6	Υ.	0.7	0.7	0.7	0.7	0.7	0.9	1.8	3.2	4.8	8.4	9.5
3.8	I	07	0.7	0.7	0.7	0.7	0.8	1.6	2.9	4.5	8.3	9.4
4,0	I	0.7	0.7	0.7	.0.7	0.7	0.7	1.4	2.6	4.2	8,0	9.2
4.2	I	0.7	0.7	0.7	0.7	0.7	0.7	1.2	2.3	4.1	7.9	9.3
4.4	I	0.7	0.7	0.7	0.7	0.7	0.7	1.0	2.0	3.4	6.8	8.7
4.6	I	0.7	0.7	0.7	0.7	0.7	0.7	0.9	1.8	3.1	6.6	8.7
4.8	I	D.7	0.7	0.7	0.7	0.7	0.7	0.8	1.6	2.9	6.2	8.5
5.0	I	0.7	0.7	0.7	0.7	0.7	0.7	0.7	1.4	2.6	6.0	8.4
* AVA	ILA	BLE FO	RM LOCA	AL WEA	THER SE	RVICE	OFFICE	-				

STATE	OF	TEXAS

FORECAST ZONE 3

LOCAL CREEK, BOYS RANCH, TEXAS

TOTAL DRAINAGE AREA CREST GAGE HEIGHT IN FEET

INITIAL STAGE = 1.2 FT

	·				. ·		FLOO GAGE TIME	D STAG ZERD TO PE	E = = 3180 AK =	5. .00 FT 2.5 H	D FT. MSL DURS
FLASH FL	DOD										
GUIDANCE				INCHES	OF R	AINFAL	L IN 3	HOURS			
VALUES*	0.5	1.0	1.5	2.0	2.5	3.0	q .0	5.0	6.0	8.0	10.0
1.D T	2.2	3.6	5.2	6.9	8.0	8.7	10.1	11.4	12.4	13.0	13.6
1.2 .1	2.0	3.2	4.6	6.1	7.8	8.4	9.8	11.2	12.4	13.0'	13.6
9.4 T	1.7	2.7	3.9	5.2	6.9	8.0	9.4	10.8	12.0	12.9	13.5
1.6 1	1.5	2.4	3.4	4.7	6.3	7.7	9.0	10.5	11.8	12.8	13.4
1.8 T	1.3	2.0	3.0	4.2	5.7	7.3	8.7	10.1	11.5	12.8	13.3
2.0 T	1.2	1.8	2.6	3.6	5.0	6.6	8.4	9.8	11.2	12.7	13.2
2.2 1	1.2	1.6	2.3	3.1	4.3	5.7	8.1	9.4	10.8	12.6	13.1
2.4 I	1.2	1.4	2.0	2.8	3.9	5.2	7.8	9.0	10.4	12.5	13.1
2.6 1	1.2	1.2	1.7	2.4	3.4	4.5	7.3	8.6	10.0	12.4	13.0
2.8 T	1.2	1.2	1.6	2.2	3.2	4.4	7.1	8.5	9.9	12.4	12.9
3.0 T	1.2	1.2	1.3	1.8	2.6	3.7	6.3	8.1	9.4	12.1	12.8
3.2 I	1.2	1.2	1.2	1.6	2.4	3.3	5.7	7.9	9.1	11.9	12.7
3.4 T	1.2	1.2	1.2	1.2	2.0	2.9	5.0	7.6	8.8	11.5	12.6
3.6 1	1.2	1.2	1.2	1.2	1.7	2.5	4.6	7.1	8.5	11.2	12.5
3.8 T	1.2	1.2	1.2	1.2	1.5	2.2	4.1	6.5	8.2	10.9	12.4
4.0 T	1.2	1.2	1.2	1.2	1.2	1.9	3.6	: 5.9	7.9	10.6	12.3
4.2 T	1.2	1.2	1.2	1.2	1.2	1.5	3.1	5.7	8.0	10.7	12.4
4.4.7	1.2	1.2	1.2	1.2	1.2	1.2	2.7	5.1	7.9	10.6	12.4
4.6 I	1.2	1.2	1.2	1.2	1.2	1.2	2.4	4.6	7.6	10.1	12.1
4.8 I	1.2	1.2	1.2	1.2	1.2	1.2	2.2	4.3	7.6	10.1	12.1
5.0 1	1.2	1.2	1.2	1.2	1.2	1.2	2.0	3.9	6.5	9.6	.11.6
* AVATIA	BIE EOI	RM LOCA	I NEA	THER SE	RVICE	OFFIC	F			• -	

VITA

John Francis Sheridan

Candidate for the Degree of

Doctor of Philosophy

Thesis: A METHODOLOGY FOR DETERMINING FLASH FLOOD FORECASTING SYSTEMS FOR A COMMUNITY

Major Field: Civil Engineering

Biographical:

- Personal Data: Born in Natick, Massachusetts, April 23, 1929, the son of Thomas J. and Margaret L. Sheridan. Married to Jo Ann Dearston 1953, two sons, John H. T. and Jeffery J. Sheridan.
- Education: Graduated from Natick High School, Natick, Massachusetts, in 1946; received a Bachelor of Science degree in Petroleum Engineering from the University of Tulsa, 1953; completed the requirements for the Master of Science degree December, 1975, from Oklahoma State University, Stillwater. Completer the requirements for the Doctor of Philosophy degree in Civil Engineering at Oklahoma State University, Stillwater, in July, 1977.
- Professional Experience: For the period 1953 through 1961 worked in the fields of Petroleum, Mechanical and Civil Engineering. Since 1961 have worked as a Hydrologist with the National Weather Service, River Forecast Center, Tulsa, Oklahoma.
- Professional Organizations: Student member of American Society of Petroleum Engineers; Registered Engineer, State of Oklahoma, Registered Land Surveyor, State of Oklahoma, Sigma XI and Chi Epsilon.