

MAPPING OF SALINE SOILS AND VEGETATION
BY REMOTE SENSING TECHNIQUES

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CHAPTER I

INTRODUCTION

General

Less than one hundred years ago physicists abandoned the view that heat is a substance and agreed that it is a form of energy transferable from one body to another by radiation, conduction or convection - or a combination of these three processes. Radiation of heat is the process of importance in remote sensing.

Earth scientists, space scientists and life scientists frequently have to detect and identify various objects and conditions in the physical universe solely from information collected by sensors that are remotely situated with respect to the area being investigated. This is called remote sensing. Energy that is either emitted or reflected from the objects of interest is recorded by the sensors, either in photographic form or in a form which can be readily reconstituted into a photo-like image.

The energy which is detected by the sensors moves with the constant velocity of light in a harmonic wave pattern and is defined to be in the electromagnetic spectrum. Distinguished by the methods used to produce and detect radiation, this spectrum has been divided historically into various regions called spectrum bands. Effectively differentiated by their interaction with matter, these bands are radio waves, microwaves,

infrared, visible, ultraviolet, x-rays, and gamma rays. This spectrum is shown in Figure 1.

Forasmuch as the intensity of radiation is simply the number of photons per unit time and area, various interactions may occur when a photon hits solid matter. According to basic physical principles, energy and mass are conserved, and energy is (1) transmitted, that is propagated through the solid matter, (2) reflected, that is returned unchanged to the medium, (3) absorbed, giving up its energy largely into heating matter, (4) emitted, or more commonly re-emitted by the matter as a function of temperature and structure, at the same or different wavelength, (5) scattered, that is deflected to one side and lost ultimately to absorption or further scattered (5).

Problem

Many soils in arid areas of the world are affected by high water tables and resultant soil salinity. Detection of the saline areas and the degree of salinity in the rooting zones is one of considerable interest to agricultural workers and engineers. The detection of these saline soils in early stages, in a precise and economical way, has been a problem which has been partially solved by the use of remote sensing techniques. It has been proved that obtaining information in more than one portion of the electromagnetic spectrum is sometimes advantageous. This method of reconnaissance is known as "multiband spectral reconnaissance" by means of which two or more remotely situated cameras or other sensing devices, each specially suited for the sensing of energy in its own spectral band, can be used together to provide information of the area that no single sensor can provide.

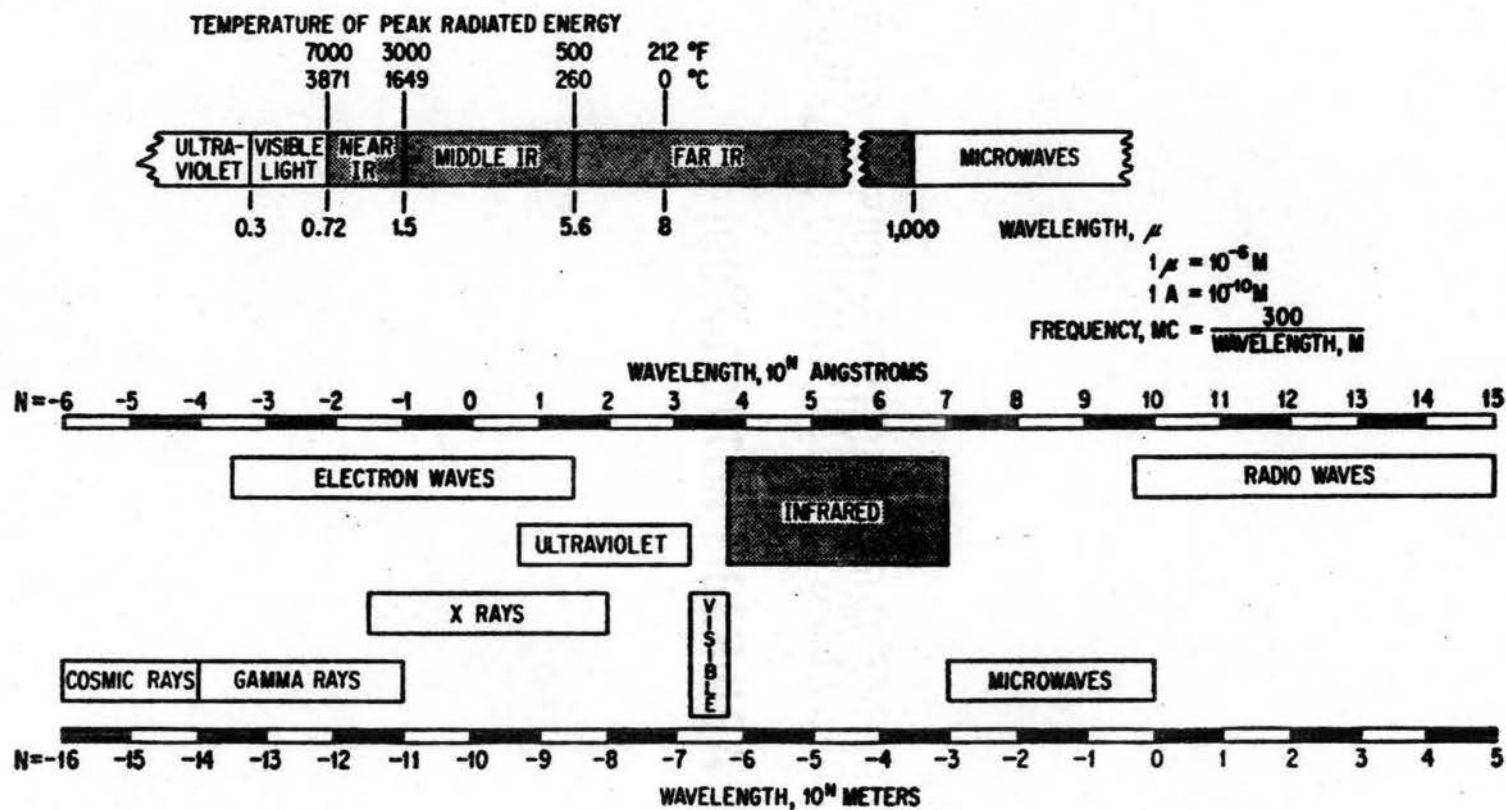


Figure 1. Position of Infrared Radiation in the Electromagnetic Spectrum

Purpose and Scope of Investigation

It is the general purpose of the research described herein to determine the type of remote sensing devices for which the problem of detecting saline soils can best be solved. This objective is to be obtained by: (1) using vegetation as indicators of salinity and determining the best film-filter combination for showing maximum contrasts in photographs of vegetation grown in different levels of salinity, and (2) directly measuring emitted energy from soils as they are affected by variations in moisture content, salt content and changing depth to free water table.

In this investigation wheat and salt grass were used as indicators and were grown in the greenhouse in soil at different levels of salinity. Wheat represents the major type of economic crop and salt grass is the major type of natural vegetation in the area under investigation. The wheat was grown in silty clay with salinity ranging from 0 ppm to 10,000 ppm, while the salt grass was grown in sand and silty clay with salinity ranging from 0 ppm to 160,000 ppm. Samples taken from the vegetation were tested in a DU-2 spectrophotometer with a reflectance attachment. Reflectance was recorded in the ultra-violet, visible and near infrared regions of the electromagnetic spectrum.

Silty clay and sand were used for determining the influence of salt content, moisture content, and changing water table on emissivity of soil. The soils, with salinity varying from 0 to 160,000 ppm, were placed in one-foot square panels and kept outdoors in the sunlight. The emissivity was measured at different periods of the day with a Precision Radiation Thermometer (PRT-5) radiometer, manufactured by Barnes Engineering Company, which measures the emitted radiation in the 8 to 14 micron window.

Silty clay and sand were also placed in an apparatus specially devised to simulate changing in water table, for saline and non-saline water, from 0 to 4 feet. The emissivity was measured for these soils in the same manner as for the soils in the panels.

CHAPTER II

REVIEW OF LITERATURE

Nature and Characteristics of Infrared Radiation

Infrared is that portion of the electromagnetic spectrum bounded on one side by the visible portion of the spectrum and the other side by micro wavelengths used for high resolution regions (Figure 1). Human eyes respond to light in the visible portion of the electromagnetic spectrum. However, to a scientist, one of the most useful regions on the spectrum is beyond that of the human eyes response - that is, the infrared region. Within this region, it has been established that all matter with temperature above absolute zero continuously emits and absorbs electromagnetic radiation. This emission is a result of the continual motion of the elementary charged particles within the substance. Whereas the thermal motion of the electrons and protons increases with temperature, it follows that the continuous radiation from a particular sample must increase with temperature. Consequently, almost everything in man's surroundings, and man himself, radiate energy in the infrared portion of the electromagnetic spectrum (10) (20).

The primary source of natural radiation is the sun; and, as solar energy impinges upon the surface of the earth, it is either absorbed or reflected. Objects with temperatures above absolute zero which absorb radiation tend to increase in temperature, then re-emit this energy

mostly in the infrared portion of the spectrum. Objects which absorb all incident radiation and emit more radiation than any other object at the same temperature are termed blackbodies; however, since a perfect blackbody does not exist, all materials are considered in terms of degree of "blackness", as gray bodies.

Today infrared, or IR, is recognized as electromagnetic energy of wavelengths between 0.7 and $1,000\mu$ (0.7 to $1,000 \times 10^{-4}$ cm). Its realm in the electromagnetic spectrum is further divided into three specific regions - near, middle, and far - with each region having radiation detectors peculiar to its location in the pattern of wavelengths. In the near infrared region, wavelengths from 0.72 to 1.5μ , radiation can be detected by photoemissive cells, photoconductive cells, photovoltaic cells and infrared photographic emulsions. Photoconductive and photovoltaic cells, together with thermal detectors at the longer wavelengths, are commonly used detectors in the middle infrared region or wavelengths from 1.5 to 20μ . Detection in the far infrared region, wavelengths from 20 to $1,000\mu$, is usually made by thermal detectors which measure the variation of some physical parameter with temperature.

Dictated by its domain in the electromagnetic spectrum, infrared exhibits some of the characteristics of both visible light and microwave radio waves. Infrared wavelengths near the visible light region can be optically focused and directed by lenses and mirrors or dispersed by prisms, whereas, infrared near the radio wave region can be propagated through some materials that are opaque to visible light (8).

Laws of Radiation

Since all objects with temperature above absolute zero radiate electromagnetic energy, the total amount of emitted radiation increases as the object's absolute temperature increases. This is shown by the Stefan-Boltzmann Law which states that the total energy (of all wavelengths) radiated from an object is given by the equation $E = \sigma(T^4 - T_o^4) e_t$

where E = energy radiated, in watts per sq cm

T = absolute temperature of the radiating body, in deg. K

T_o = absolute temperature of the radiating surrounding surfaces,
in deg. K

e_t = total emissivity (unit of blackbody source)

$$\sigma = 5.673 \times 10^{-12} \text{ watt cm}^{-2} \text{ } ^\circ\text{K}^{-4}$$

If T is greater than T_o , the error introduced by neglecting T_o is less than .5 percent and the equation becomes

$$E = \sigma T^4 e_t = 5.673 \times 10^{-12} T^4 e_t \quad (8) \quad (11).$$

The emissivity factor (e_t) in the Stefan-Boltzmann Law is the ratio of an object's total radiant emittance to that of the total radiant emittance of a blackbody at the same absolute temperature. For instance mirrors have emissivity of approximately .05 while a lamp black is approximately .95. Therefore, the smoother and shinier the surface of an object, the amount of radiation which will be absorbed is smaller.

The Stefan-Boltzmann Law shows how much energy is radiated at a given temperature, but it does not give any information about the distribution of energy. This is given by Planck's Law (8) (11)

$$E_{\lambda} d\lambda = C_1 \lambda^{-5} (e^{(C_2/\lambda T)} - 1)^{-1}$$

where $C_1 = 3.732 \times 10^{-12} \text{ watt cm}^2$

$$C_2 = 1.436 \text{ cm } ^\circ\text{K}$$

$$e = 2.718$$

$E\lambda d\lambda$ = energy radiated in wavelength

λ = wavelength, in cm

T = absolute temperature, in $^\circ\text{K}$

Planck's Law for any temperature may be represented by a curve giving the relationship between energy radiated and wavelength as shown in Figure 2. The area under the curve is total energy radiated and graphically represents the Stefan-Boltzmann Law.

In nature, objects are not perfect radiators throughout the entire spectrum and therefore will reflect or transmit as well as absorb some portion of incident radiation. The total amount of emitted radiation reaches a peak at certain wavelengths. This is shown through Wien's Displacement Law which states that the wavelength at which peak radiation power occurs equals a constant ($K = 2.89 \mu ^\circ\text{K}$) divided by the object's absolute temperature in degrees Kelvin ($\lambda = K/T$). This peak radiation power shifts toward the shorter wavelength as the temperature of the radiator is increased, as shown in Figure 2 (9). The well known variation in the color of hot objects with temperature is explained by these curves in Figure 2. At room temperature virtually all the radiation is in the infrared region of the electromagnetic spectrum, while as the temperature increases more and more of the radiation appears in the visible range. At first it is in the red end of the spectrum, with

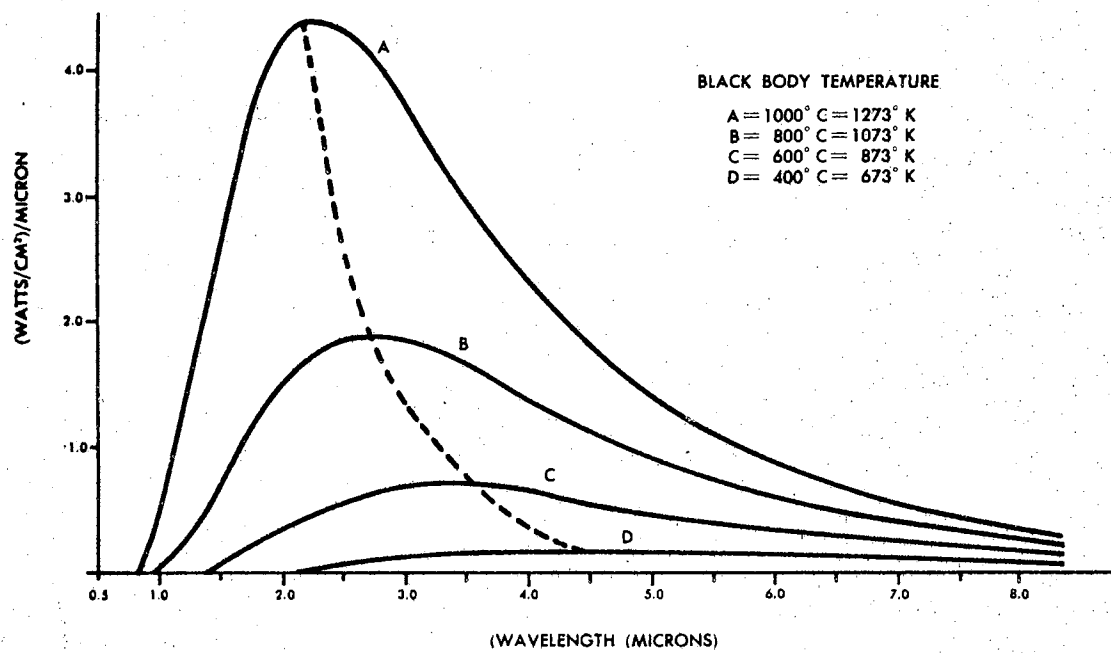


Figure 2. Blackbody Radiation Curves

the result that the object appears red. When its temperature is of the magnitude of the temperature of the sun's surface (6000°K) the radiation is distributed throughout the visible region to produce the color sensation which we call white. The earth has an average temperature of 300°K which is shown to have its peak radiation emittance power at $9.6\ \mu$ (Figure 2). This radiant energy is an excellent energy source and permits the detection of terrestrial objects by the use of emitted energy in the infrared spectrum band. Since the sun is an excellent energy source it permits the detection of terrestrial objects by reflected energy in visible and near infrared spectrum bands by use of panchromatic and infrared sensitive film (17).

Atmospheric Absorption and Scattering

Due to the passage of radiation through the atmosphere, it is necessary to consider the atmospheric constituents and their effects on radiation. These constituents, such as gasses, smoke and dust particles, water vapor, and water droplets, cause losses by scattering and/or reflecting and absorbing the radiation. Particles in the atmosphere of approximately $.25\ \mu$ in size greatly scatter radiation of ultraviolet and blue violet frequency with scattering decreasing toward the longer wavelengths in the visible portion of the spectrum. Likewise fog particles of approximately $4\ \mu$ radii will completely attenuate by scattering radiation in the near and much of the middle infrared region.

A typical transmission spectrum of the atmosphere, as shown in Figure 3, depicts the transparency of the atmosphere at different wavelengths along the spectrum. At certain wavelengths it is nearly opaque. Wavelength ranges for which the atmosphere is transparent to radiation

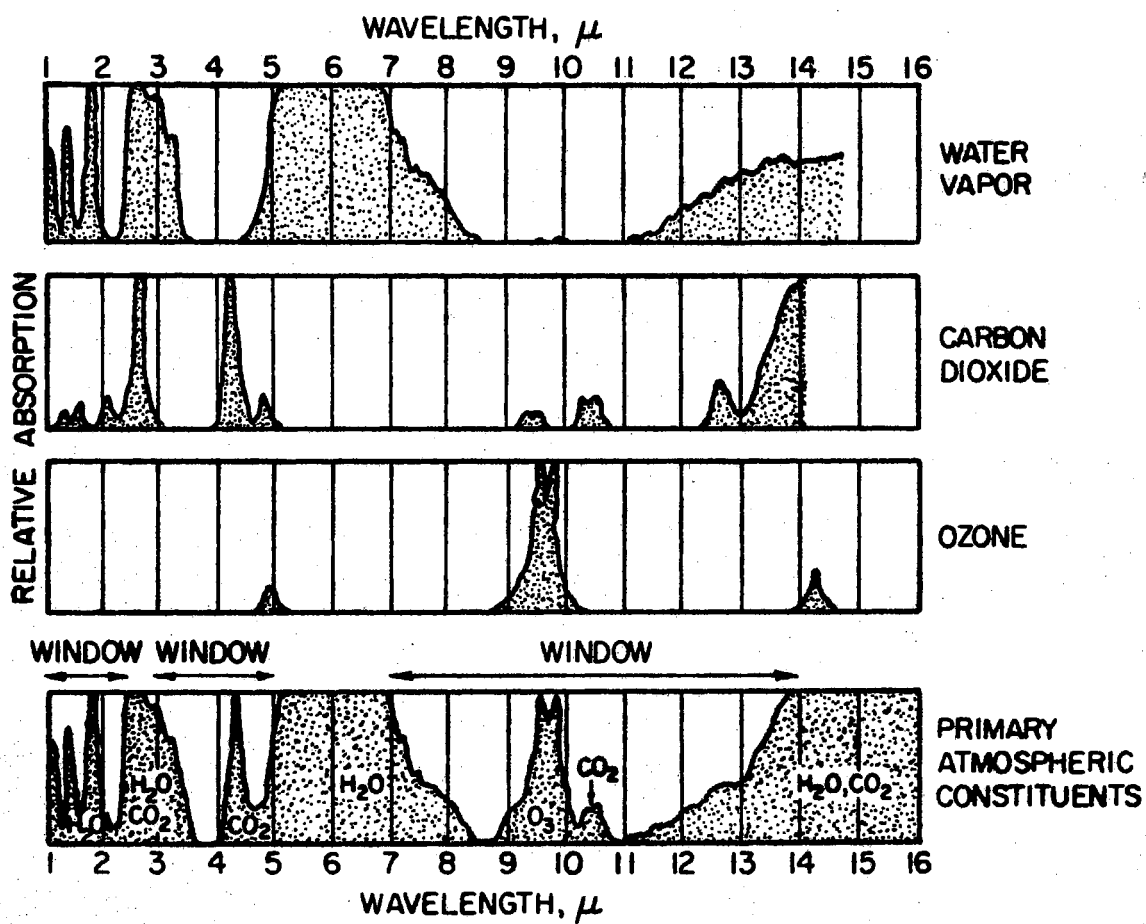


Figure 3.. Infrared Absorption Bands of Primary Atmospheric Constituents

are called "windows". "Windows", however, are not perfectly transparent, 80% transmission would more nearly represent their transparency.

The amount of atmospheric absorption and scattering is a function of pathlength, the amount of scattering and absorbing constituents within this pathlength, and the wavelength particle size relationship. A number of different atmospheric constituents attenuate an infrared beam as it passes through the atmosphere. Most of the absorption of the radiation in the beam is due to molecules of water vapor (H_2O), carbon dioxide (CO_2), and ozone (O_3). However, ozone is limited to a layer at an altitude of approximately 80,000 feet and for most purposes the effect of it on infrared transmission below this altitude can be neglected. The effect of minor atmospheric constituents such as methane (CH_4), nitrous oxide (N_2O) and carbon monoxide (CO) may be observed over long paths. Two oxygen absorption bands occur at 1.06 μ and 1.27 μ in the infrared portions of the spectrum. The absorption of the electromagnetic energy by water vapor virtually closes the atmosphere beyond the 25 micron wavelength to the start of the microwave region (11) (8).

Remote Sensing

Today's modern aircraft and earth satellites make it possible to survey the earth from a vantage point heretofore unobtainable. By this method of remote sensing the identification of an object or determination of some properties of the object can be done in two ways. The identification may be obtained on the basis of shape, size, color or other criteria; or it can be obtained by duplicating our visual senses by photography. In many cases the measurement we wish to make is one of quantity, i.e., the absolute value of the radiation; while on the other

hand, we may wish to measure qualitative aspects, i.e., the relative value of the radiation and thus acquire the information that this conveys in the light of our knowledge of some of the physical properties of the object.

There are two types of remote sensing, active and passive systems. The active system is one that actually illuminates the object under investigation with radiation of a particular wavelength and then samples the portion reflected back to the detecting device. Radar is the best example of an active system among the radiation sensors. Night photography, where an artificial illumination system is required, would be another example (17).

The passive system is one that samples emitted and reflected radiation from a source. Radiometers and scanners that measure emitted and reflected radiation in the infrared and microwave regions are examples of passive systems (17).

Due to the fact that matter which is struck by photons either transmits, reflects, emits, absorbs or scatters electromagnetic energy at different magnitudes is the principle by which different types of matter are identified by measuring these differences with some type of remote sensor (5).

In selecting the type of sensor to be used, one should consider these three aspects, (1) what atmospheric window is to be used, (2) in what part of the electromagnetic spectrum will best detection be possible, and (3) what properties are to be detected. For example, if the reflectance properties are to be detected by a sensor these are best disclosed in the atmospheric window between .4 to .9 microns by photography. The photography is usually obtained by an Itek nine lens multi-band camera shown in Figure 4.

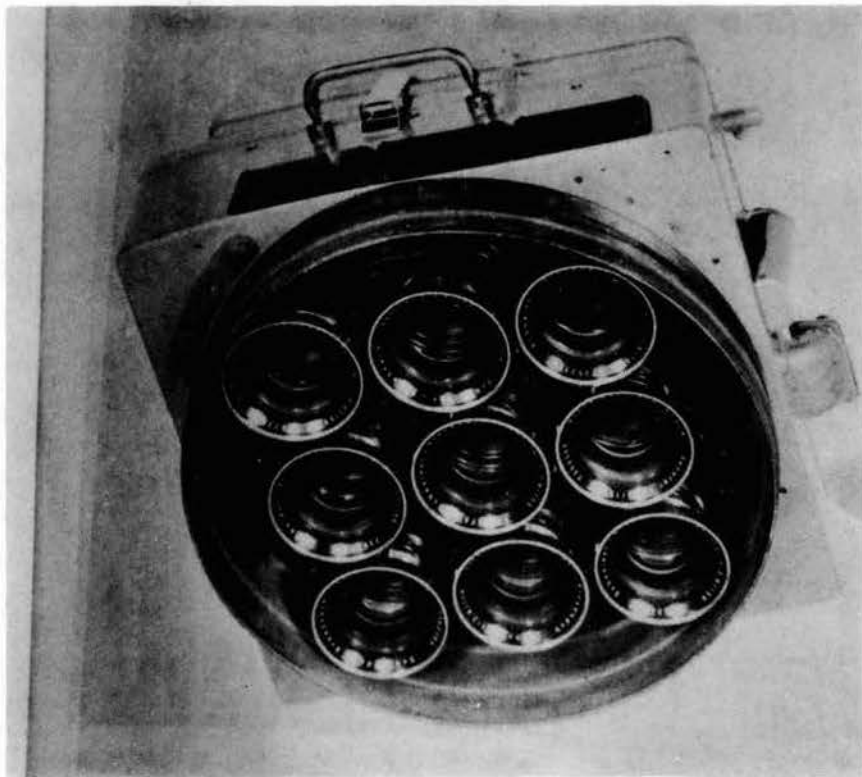


Figure 4. Itek Nine Lens Multiband Camera

The camera consists of nine matched lenses with different filters to take nine simultaneous exposures of the same identical ground area. By using different filters it is possible to filter out different portions of the spectrum to show which part of the spectrum will detect the particular property of the object which is of interest to the viewer (13).

If the emitted energy is the property which is to be investigated by a sensor, the atmospheric windows between 3.5 to 5.5 and 8 to 14 microns are used. Furthermore, since approximately 30% of the radiation emitted by terrestrial objects falls within the 8 to 14 micron range, it is a good range of the spectrum for detecting emitted energy. The measuring of emitted energy by sensors is usually done by use of an electronic instrument which measures emitted energy in degrees centigrade or Fahrenheit and is called a radiometer. See Figure 10.

Effects of Salinity on Plants

Plants grown in soils having excessive soluble salt concentrations in the root zone are affected in many ways. The increased osmotic pressure of saline solutions restricts uptake of water by plant roots which affects the growth of plants. Specific ion toxicity can also affect plant growth and appearance. Since the behavior of plants is a good index of their environment, plants are frequently used as indicators of conditions that occur below the soil surface. These are manifested in the plants' appearance and spectral reflectance from the leaf surfaces. Since the root systems of plants explore a rather large soil volume, a plant sample is more representative of the site conditions than is a single soil sample (14).

Because of a leaf's physiological structure light of different wavelengths is reflected from different portions of the leaf's cross-section (Figure 5). The chloroplasts which are used in photosynthesis affects the electromagnetic spectrum between .3 - .6 microns. The blue and red lights are largely absorbed by the chloroplasts while the green light is largely reflected. Chloroplasts, being transparent to the infrared portion of the spectrum, have little effect on infrared radiation striking it, however, the spongy mesophyll affects the infrared by highly transmitting and reflecting it (4) (15).

A leaf's reflectance changes as it matures because its spectral properties are closely associated with cell structure. Due to an increase in chlorophyll with maturity, the reflectance of leaves increases in the visible region of the spectrum with maturity. Reflectance also increases with maturity in the near infrared region. It is probable that the near infrared reflectance is a function of the cell shape and size as well as the amount of intercellular space. Initially, the mesophyll of the very young leaf consists primarily of spongy parenchyma with considerable air spaces which are favorable to the mechanism of internal reflection. Then as the leaf matures these cells enlarge, crowding together, reducing the intercellular space and increasing the reflectance. It would then appear that during the final maturing the cell structure and intercellular space relationship become favorable for increased reflectance. In order to establish this relationship, careful experiments will need to be done relating the reflectance to the leaf morphology (21).

The relative turgidity, which is a measure of field moisture content of plants expressed as a percentage of their turgid water content, does not have much effect on the reflectance of leaves in the visible

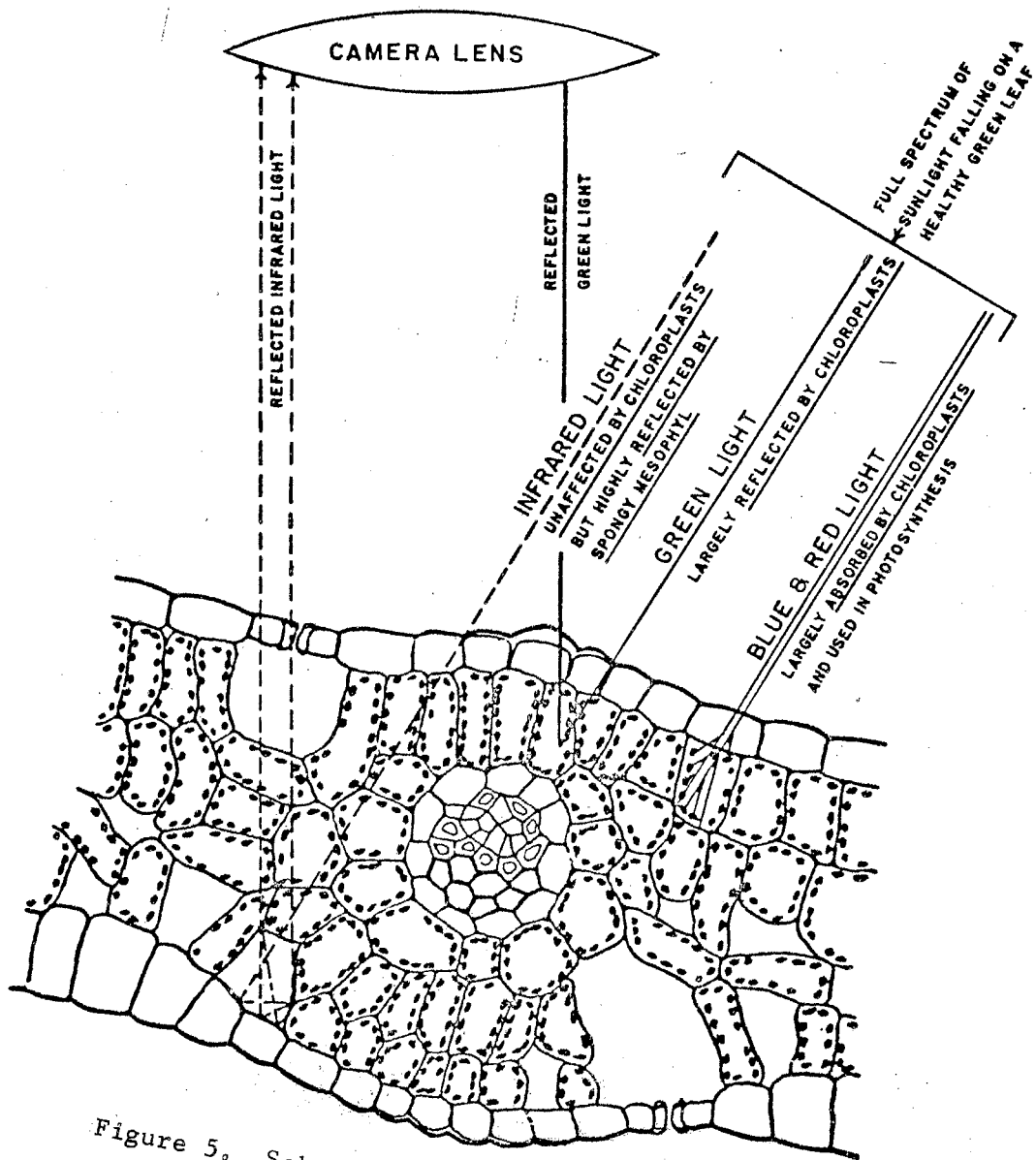


Figure 5. Schematic Drawing of Cross-Section of Healthy Leaf

range of the spectrum. While reflectance does increase as relative turgidity decreases, it must be less than 80% to affect reflectance.

Reflectance of leaves in the infrared region of the electromagnetic spectrum increases when leaves are stacked on each other, while no effect in magnitude of reflectance occurs in the visible region. This shows that the reflectance in the visible region is a surface phenomenon and that in the infrared portion it depends on thickness of leaves. Thus a spectrophotometer study of single leaves can be very misleading in terms of the magnitude of reflectance of the infrared region of vegetation in the field (16).

A study by Myers, et. al. (1963) predicted the expected tonal contrasts on the basis of single leaf thickness spectrophotometer studies. The study showed that the differences in reflectance between cotton leaves affected by salinity and non-affected leaves to be only 2 percent. On the other hand, infrared photographs taken of the area showed quite a contrast in the tone of the affected and non-affected plants. This is believed to be due to the reinforcement of reflectance by multiple leaf layers (16).

Since the materials of the leaf that affect spectral reflectance are common to all agricultural plants, all foilage exhibits the same general pattern of reflectance. But it is important to recognize that other factors such as crop geometry and background soil conditions affect the reflectance measured in the field.

Numerous factors influence the timing of the picture taking mission and the photo interpretation. Experience has shown that physiological drought caused by salinity is most apparent in the morning hours. In the afternoon hours leaves lose turgidity as a result of high ambient

temperature, an effect that makes part of the evidence of physiological drought predominant (15).

Detection of Saline Soils

Many areas of the United States are plagued with the problem of saline soils with a high water table. The detection of these saline soil areas is of special significance to agricultural scientists, agronomists and civil engineers. Remote sensing has been considered by some scientists as an effective approach to the detection and mapping of such areas.

The detection of salinity through the atmospheric window 8 - 14 microns is done through the temperature characteristics of the soil. Therefore, the emitted energy is a function not only of the physical characteristics of the surface but of the surface temperature as well. Since the surface temperature is dependent upon factors which are affecting the soil temperature at depth it is possible to detect such things as water table by use of these longer wavelengths.

The ability of soil to radiate thermal energy is dependent upon many factors, such as emissivity, thermal conductivity, tone, specific heat, and thermal diffusivity of the soil. All of these factors are raised as the moisture content rises in the soil due to water replacing air voids (2). Therefore water has a large effect on the rate at which thermal energy is emitted from the soil. Since different concentrations of salt in the soil change the rate of evaporation in the soil's moisture content, it is possible to detect its amount of salinity.

Due to the rate of differential heating and cooling of the soil, there are times during the day when maximum differences in emitted

energy will be observed. These differences depend upon the climate and type of soil. There is also a time of day when there will be no difference in emitted thermal energy (7). Thus, the best time for detection has to be found by trial and error.

CHAPTER III

MATERIALS AND LABORATORY PROCEDURES USED IN INVESTIGATION

General

The soils used in this investigation were silty clay and sand. Both types are found in the Salt Plains area. Since such large amounts of soils were needed during the investigation, the samples to be used were taken from the area surrounding Stillwater, Oklahoma and not from the Salt Plains.

The types of vegetation grown in the greenhouse during the investigation were wheat and salt grass. The choice of these types of vegetation was made on the basis of salt grass being a native grass which can be grown in soils having high saline content; and, wheat being the most prominent crop grown in the Salt Plains area.

Properties of Soils

Sand

The sand used in the investigation was river sand having a specific gravity for particles passing a No. 10 sieve of 2.69. A sieve analysis of the sand is shown in Figure 6. All sand used in the investigation was sieved through a No. 10 sieve. Table I shows the results of the chemical analysis of the sand.

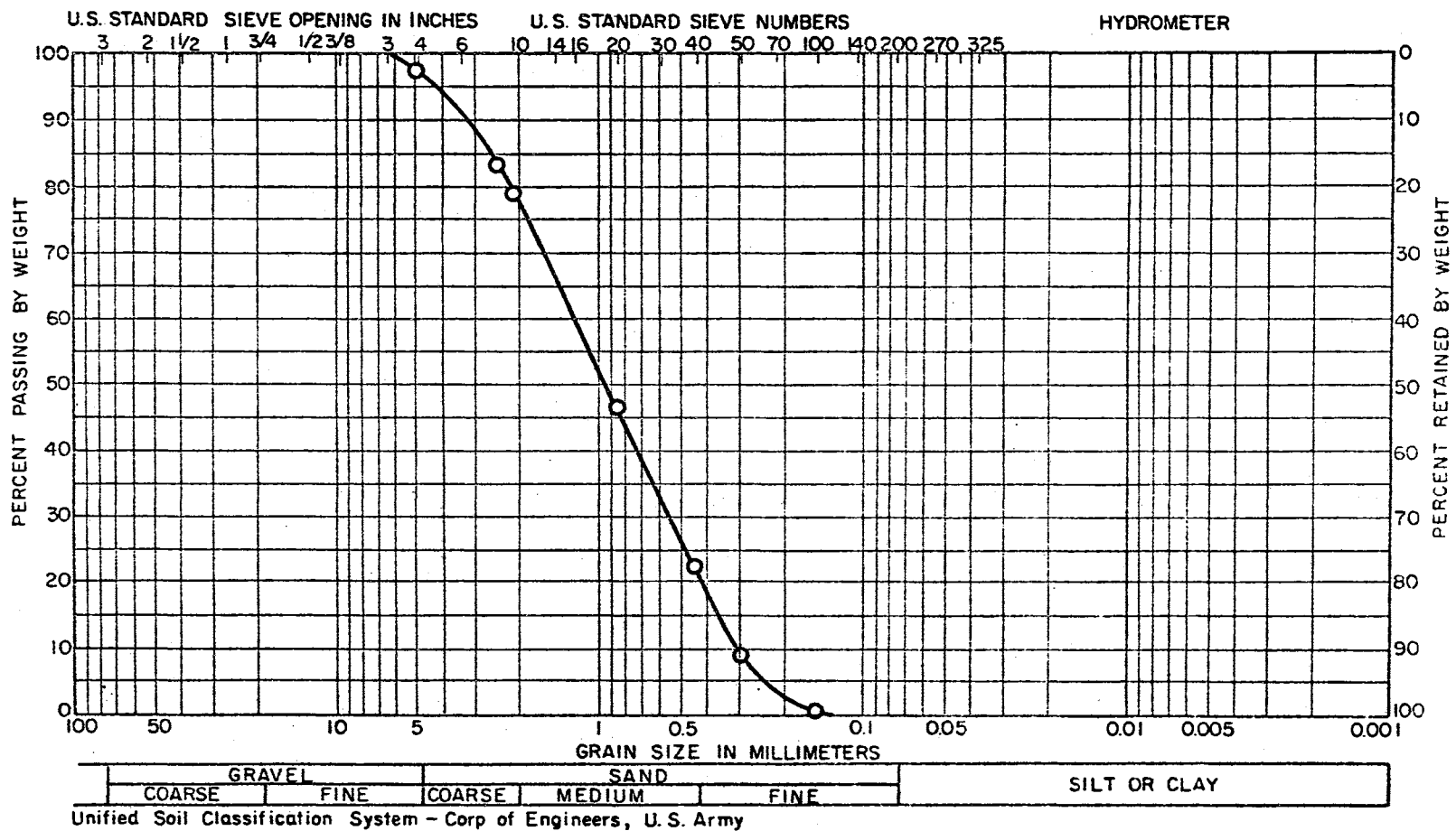


Figure 6. Grain Size Distribution Curve for Sand

TABLE I
CHEMICAL ANALYSIS OF SAND

Soil Test Results	Reading	Adjective Rating
Organic Matter (Nitrogen %)	0.07	Very low
Phosphorous #/A (lbs/acre)	9.40	Very low
Potash #/A (lbs/acre)	35.00	Very low
PH (Acidity or Alkalinity)	7.20	Neutral to very slightly alkalinity

Silty Clay

Standard sieve and hydrometer analyses on samples of the silty clay yielded the grain size distribution curve shown in Figure 7. The specific gravity, determined on samples passing a standard No. 16 sieve, was 2.77. All silty clay used in this investigation was sieved through a No. 16 sieve. The Atterberg Limits of the soil were:

Liquid Limit . . . 28.5

Plastic Limit . . 17.9

Table II shows the results of a chemical analysis of the silty clay.

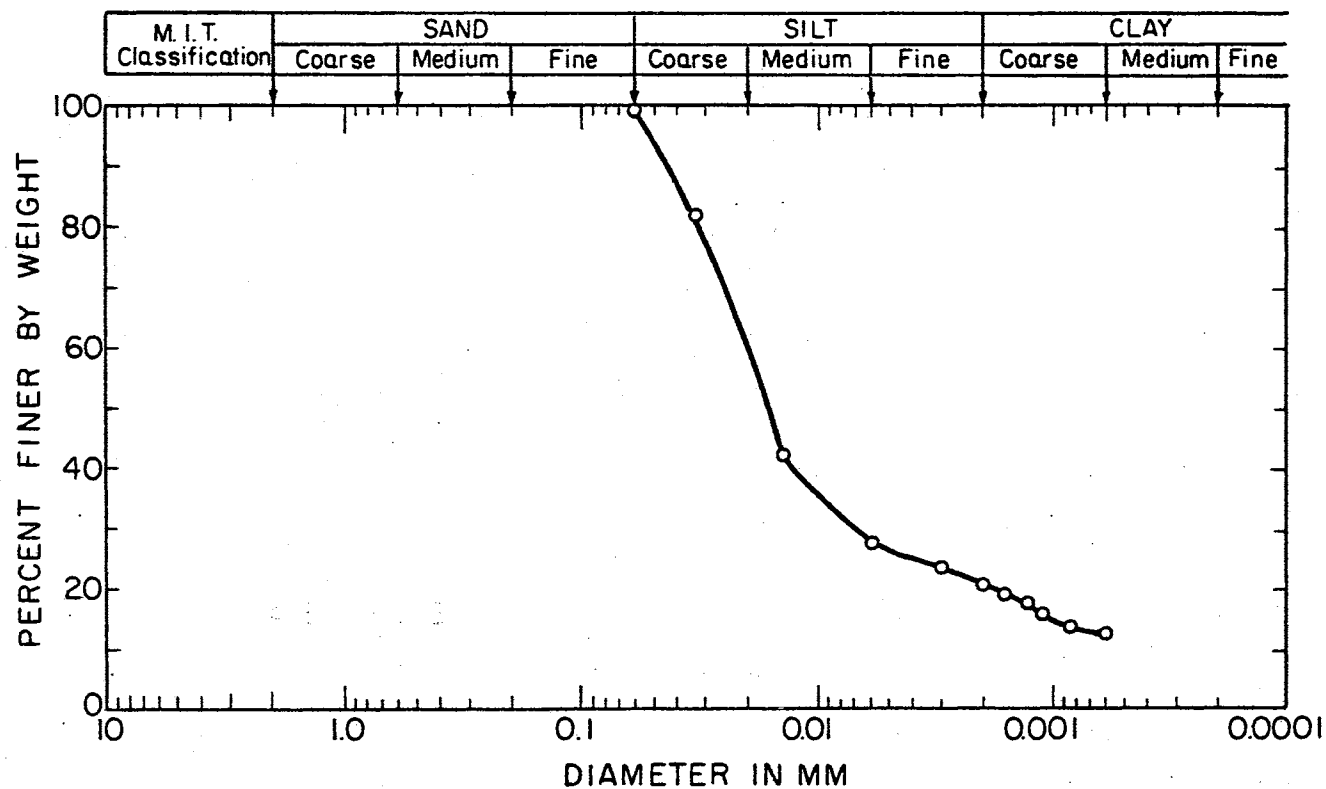


Figure 7. Grain Size Distribution Curve for Silty Clay

TABLE II
CHEMICAL ANALYSIS OF SILTY CLAY

Soil Test Results	Reading	Adjective Rating
Organic Matter (Nitrogen %)	1.35	Very low
Phosphorous #/A (lbs/acre)	105.60	Very low
Potash #/A (lbs/acre)	375.00	Very high
PH (Acidity or Alkalinity)	5.95	Moderately Acid

Water Holding Capacity

The ability of soils to hold water depends on the grain size and the amount of organic matter present in the soil. Water holding capacity of a soil is determined by placing a sample of soil in water until it is saturated and then draining it for 24 hours. During this period most of the excess of non-capillary water is drained away by gravity, and the soil attains the water holding capacity (22).

Procedure

Two-inch square brass containers with screened bottoms were used for soil samples. Filter paper was placed in the screened bottoms, and the containers were filled with soil to two-thirds capacity. The containers were then placed in a water bath one-half immersed. After remaining in the water bath for 24 hours, they were taken out and allowed to drain for another 24-hour period. Having drained, they were weighed and placed in an oven at 110° C. for 24 hours, then reweighed. The

percent of water holding capacity is $\left(\frac{X - Y}{Y - Z}\right) \times 100$, where,

weight of drained sample X grams
 weight of oven-dried sample Y grams
 weight of containers Z grams
 content of water (X - Y) grams

The water holding capacities of the silty clay, sand and sandy clay mixture were found to be 45, 18 and 20, respectively.

Relative Turgidity and Water Content of Leaves

Water content of the leaves is simply the amount of water by weight which is in the leaves. Relative turgidity is the amount of water by weight which the leaves can hold. The method used in determining the relative turgidity was to weigh three or four leaves and placing them in a dish of distilled water in a refrigerator at the temperature of 7° C. for 24 hours. Leaf cuts were then blotted dry with tissue paper, weighed immediately, and oven-dried at 85° C. for 24 hours. After this period, the leaf cuts were then reweighed (19).

The relative turgidity of leaves was expressed as a percentage and evaluated according to the equation:

$$\text{R.T. \%} = \left(\frac{X}{Y}\right) \times 100$$

The water content of leaves was expressed on a fresh weight basis and calculated as follows:

$$\text{W.C. \%} = \left(\frac{A - B}{A}\right) \times 100$$

where,

leaf tissue fresh weight A

leaf tissue oven-dried B

leaf tissue saturated weight C

water content of the fresh leaf tissue $X = \left(\frac{A - B}{A} \right)$

water content of the same leaf tissue at full saturation $Y = \left(\frac{C - B}{C} \right)$

CHAPTER IV

INVESTIGATIVE PROCEDURES FOR REFLECTANCE

General

Since vegetation is affected by chemicals present in the soil, it has been proved to be one of the best indicators of salinity in soil profiles. Therefore, a reflection study was made of vegetation native to the Salt Plains area to determine if these types of vegetation could be used as indicators of salinity in soils of that area.

The vegetation was grown in the greenhouse under controlled climate. Samples were contained in two-gallon buckets. By growing the vegetation under controlled conditions it was possible to eliminate some of the variables which could affect the growth of the plants. The salt grass was grown at five levels of salinity, that is 0, 20,000, 60,000, 80,000, 100,000 and 160,000 ppm and in two types of soil -- silty clay and sandy clay mixture. The wheat was also grown at five levels of salinity -- 0, 2,000, 4,000, 6,000, 8,000 and 10,000 ppm, but only in one type of soil, silty clay. The wheat was grown at much lower levels of salinity since the salts become toxic to wheat at lower levels than salt grass.

Determination of Salt Contents

The determination of the levels of salinity at which the vegetation should be grown was made by the osmotic values found to be present in

the soils of the Salt Plains area by Dr. Elsharkawi (6). The predominant salts found in the area were sodium chloride and calcium chloride. In calculating the amounts of both sodium chloride and calcium chloride which when dissolved in one litre of water give a solution of particular osmotic pressure, the equation followed by Lagerwerff and Holland (1960) was used (12).

The amount of salts to be used at different osmotic values is dependent upon the level of the sodium absorption ratio (SAR) at which the sodium becomes toxic to the plant. Dr. Elsharkawi found an SAR value of 1/8 to be low enough to avoid sodium toxicity in wheat and salt grass.

$$SAR = \frac{Na^+}{(Mg^{++} + Ca^{++})} \text{ m moles}^{1/2} \text{ litre}^{-1/2}$$

However since the amount of magnesium in the soils is negligible compared to that of calcium (usual in soils with high calcium content) the equation can be written as

$$SAR = \frac{Na^+}{(Ca^{++})} \text{ m moles}^{1/2} \text{ litre}^{-1/2}$$

The amount of sodium chloride and calcium chloride for a solution of one atmosphere at an SAR of 1/8 is

$$SAR = \frac{Na^+}{(Ca^{++})^{1/2}} = 1/8 \quad \therefore 8Na^{++} = (Ca^{++})^{1/2}$$

or

$$64(Na^+)^2 = Ca^{++} \text{ m moles litre}^{-1} \quad (1)$$

And since 32 me $CaCl_2$ per litre and 24 me $NaCl$ per litre constitute a solution of one atmosphere osmotic pressure

$$\therefore \frac{\text{me NaCl}}{24} + \frac{\text{me CaCl}_2}{32} = 1 \text{ atm} \quad (2)$$

Thus substituting for Ca in equation (2) its value in equation (1)

$$\text{NaCl} + 48(\text{NaCl})^2 = 24 \quad (3)$$

Equation (3) is a quadratic equation of the form

$$Ax^2 + Bx = C$$

where x can be determined from the solution of the equation

$$x = \frac{-B \pm (B^2 + 4AC)^{1/2}}{2A}$$

$$\therefore \text{NaCl} = \frac{-1 \pm [1 - 4 \times 48 \times (-24)]^{1/2}}{2 \times 48} = \frac{-1 \pm (4609)^{1/2}}{96}$$

$$= \frac{66.89}{96} \text{ me litre}^{-1}$$

or

$$\text{CaCl}_2 = 64 \times \left(\frac{66.89}{96}\right)^2 \text{ me litre}^{-1}$$

$$\text{CaCl}_2 = 64 \times \left(\frac{66.89}{96}\right)^2 \times 55.49 = 1723.83 \text{ mg litre}^{-1}$$

The amounts of sodium chloride and calcium chloride to be dissolved in one litre to give osmotic pressures from 0 to 170 atmospheres were calculated and are shown in Table III and Figure 8.

TABLE III
CALCULATED AMOUNTS OF NaCl AND CaCl₂ PER LITRE
WHICH GIVE DIFFERENT VALUES OF OSMOTIC
PRESSURE IN SOLUTION AT FIXED
SAR OF 1/8

Osmotic Pressure in Atm.	NaCl(g)	CaCl ₂ (g)	Total Salt in ppm
0.5	0.02865	0.8525	880
1.0	0.0407	1.723	1764
1.5	0.0501	2.610	2660
2.0	0.0570	3.480	3537
2.5	0.0645	4.340	4405
3.0	0.0712	5.250	5321
3.5	0.0768	6.118	6194
4.0	0.0821	6.997	7079
4.5	0.0871	7.882	7969
5.0	0.0918	8.760	8852
10.0	0.1300	17.58	17710
15.0	0.1592	26.35	26509
20.0	0.1842	35.28	35464
25.0	0.2060	43.81	44016
30.0	0.2257	52.98	53206
40.0	0.2607	70.69	70951
50.0	0.2916	87.35	87612
60.0	0.3195	99.75	100069
90.0	0.3914	159.31	159701
110.0	0.436	198.75	199186
130.0	0.4708	250.24	250711

Method of Growing Vegetation

In preparing the soil in which the vegetation was to be grown, the silty clay and sandy clay mixture was dried, sieved and placed in buckets. Buckets for the wheat and one-half of the salt grass were filled with thirty pounds of silty clay. The remaining half of salt grass buckets

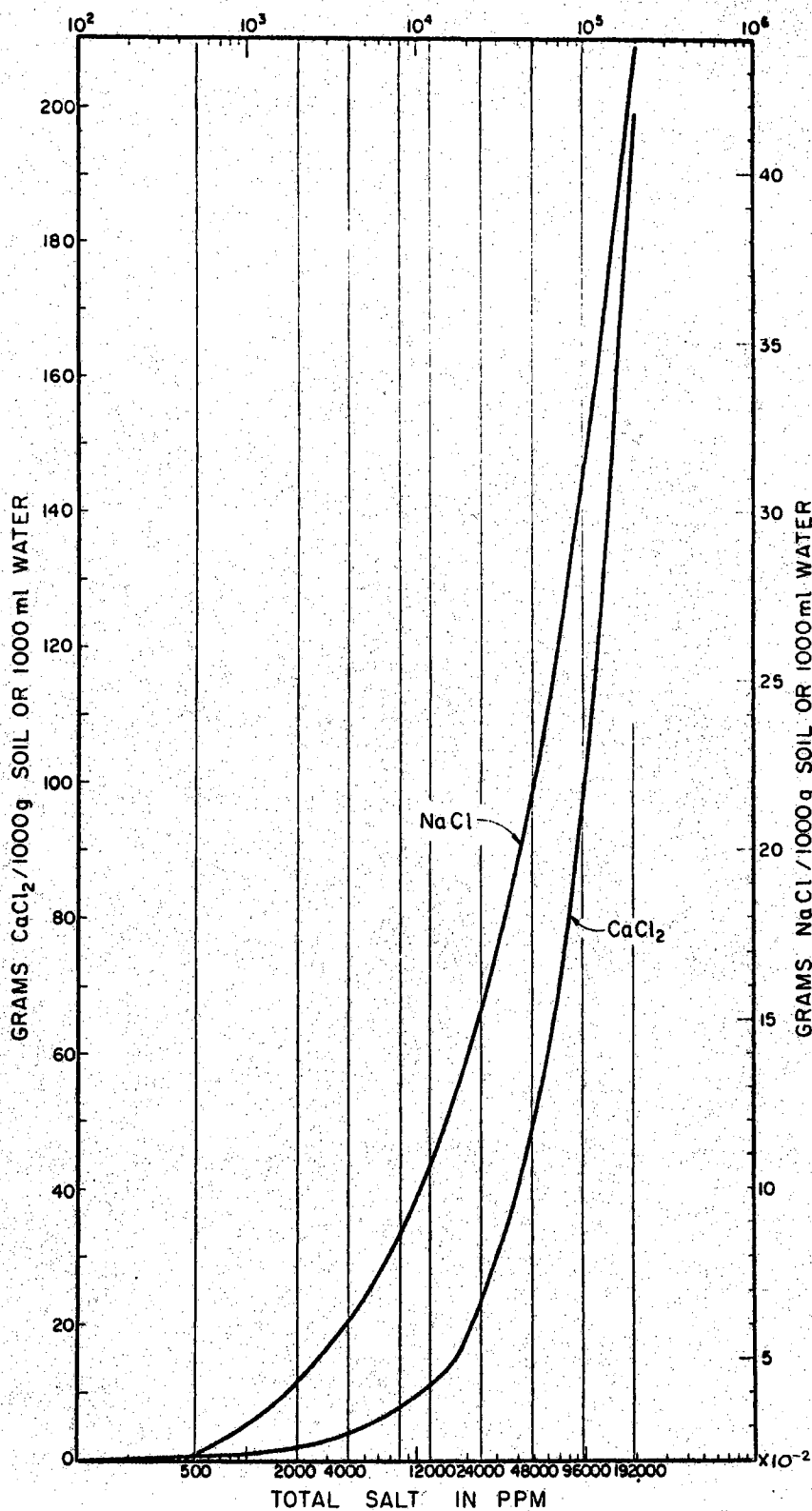


Figure 8. Amounts of NaCl and CaCl₂ that Give Concentrations up to 192,000 ppm at Fixed SAR of 1/8

were filled with thirty-five pounds of sandy clay mixture. Salts for the different levels of salinity were determined and are given in Tables IV, V, and VI.

The amounts of salts to be added to the soil were based on ppm in solution. Therefore the amount of water which can be held by the soil, or its water holding capacity, under saturated conditions had to be determined. For example, the amount of water required to saturate thirty pounds of silty clay at loose state is

$$30 \text{ lbs} \times 453 \text{ gm/lb} = 13,590 \text{ grams}$$

$$13,590 \text{ gm} \times .45 = 6,115 \text{ gm} \approx 6.1 \text{ litres}$$

where water holding capacity of silty clay = 45%. Therefore the total salts to be added to thirty pounds of silty clay is

$$\begin{aligned} \text{total CaCl}_2 &= \text{CaCl}_2 \text{ (gm/litres for particular ppm} \\ &\quad \text{solution)} \times \text{(number of litres to saturate} \\ &\quad \text{thirty pounds of silty clay)} \end{aligned}$$

$$\begin{aligned} \text{total NaCl} &= \text{NaCl (gm/litres for particular ppm} \\ &\quad \text{solution)} \times \text{(number of litres to saturate} \\ &\quad \text{thirty pounds of silty clay)} \end{aligned}$$

Salts were added and mixed thoroughly to the soils in the amounts given in the tables.

The salt grass and wheat were then planted in the buckets and the soil was saturated. The salt grass was planted from roots taken from the Salt Plains area and the wheat was planted with seeds. The soil was kept saturated throughout the growing period of the vegetation. The

TABLE IV

AMOUNT OF SALTS FOR TREATMENT OF WHEAT
GROWN IN 30 LBS. OF SILTY CLAY

ppm	CaCl ₂ gm/litre	NaCl gm/litre	litres	CaCl ₂ gm	NaCl gm
2000	1.96	.043	6.1	11.9	.26
4000	3.94	.0615	6.1	24.03	.37
6000	5.73	.074	6.1	34.95	.45
8000	7.72	.086	6.1	47.09	.52
10000	9.90	.097	6.1	60.40	.59

TABLE V

AMOUNT OF SALTS FOR TREATMENT OF SALT GRASS
GROWN IN 30 LBS. OF SILTY CLAY

ppm	CaCl ₂ gm/litre	NaCl gm/litre	litres	CaCl ₂ gm	NaCl gm
20000	19.85	.138	6.1	121.08	.84
60000	59.70	.237	6.1	364.17	1.44
80000	79.70	.277	6.1	486.00	1.68
100000	99.70	.307	6.1	608.10	1.87
160000	159.30	.391	6.1	971.00	2.38

TABLE VI
AMOUNT OF SALTS FOR TREATMENT OF SALT GRASS
GROWN IN 35 LBS. OF SANDY CLAY

ppm	CaCl ₂ gm/litre	NaCl gm/litre	litres	CaCl ₂ gm	NaCl gm
20000	19.85	.138	3.2	63.52	.44
60000	59.70	.237	3.2	191.04	.76
80000	79.70	.277	3.2	255.00	.88
100000	99.70	.307	3.2	319.00	.98
160000	159.30	.391	3.2	509.76	1.25

vegetation was irrigated through one-inch plastic tubes which were punctured with 1/8 holes all around them. These tubes were placed down through the soils to the bottom of the buckets in order to prevent the salts from being leached out of the top of the soil when water was added. The soil was irrigated everyday to keep it completely saturated throughout the growing period.

Procedure for Measuring Reflectance

In order to measure reflectance of the vegetation, it was grown for one month allowing a suitable growth of leaves and time for the salts to affect the plants. At the end of this growing period, leaves which appeared to be healthy to the visible eye, were harvested for reflectance investigation. These leaves were then placed on 2-inch by 1.5-inch

metal strips which had been prepared with double adhesive tape to hold the leaves in place. Care was taken in placing the samples on the metal strips, making a uniform cover of leaves with no overlapping or gapping between the leaves.

The reflectance was measured on a DU-2 Spectrophotometer with reflectance attachment connected to chart recorder shown in Figure 9. Four samples of each level of salinity were tested in this investigation. To guard against the leaves losing their freshness, only one sample was prepared at a time. Only twenty minutes elapsed from the time the leaves for one sample were harvested and reflectance measured from 200 to 1000 millimicrons. The four samples of each level of salinity were run on the same chart so the high and low values could be determined and plotted on graphs given in Chapter VI.

The reflectance measurements cannot be read directly from the chart; they must be calculated. The reflectance is equal to the ratio of the reference sample to the vegetation sample, using a piece of magnesium carbonate as the reference sample. The relative turgidity for the different levels of salinity were also determined and are given in Appendix A.

Description of DU-2 Spectrophotometer

The spectrophotometer is an instrument which measures the radiant energy reflected or transmitted by a specimen. It then compares these measurements to those of radiant energy reflected or transmitted by another substance, such as magnesium carbonate, taken as a reference standard. The following is a list and brief description of the essential spectrophotometer components:

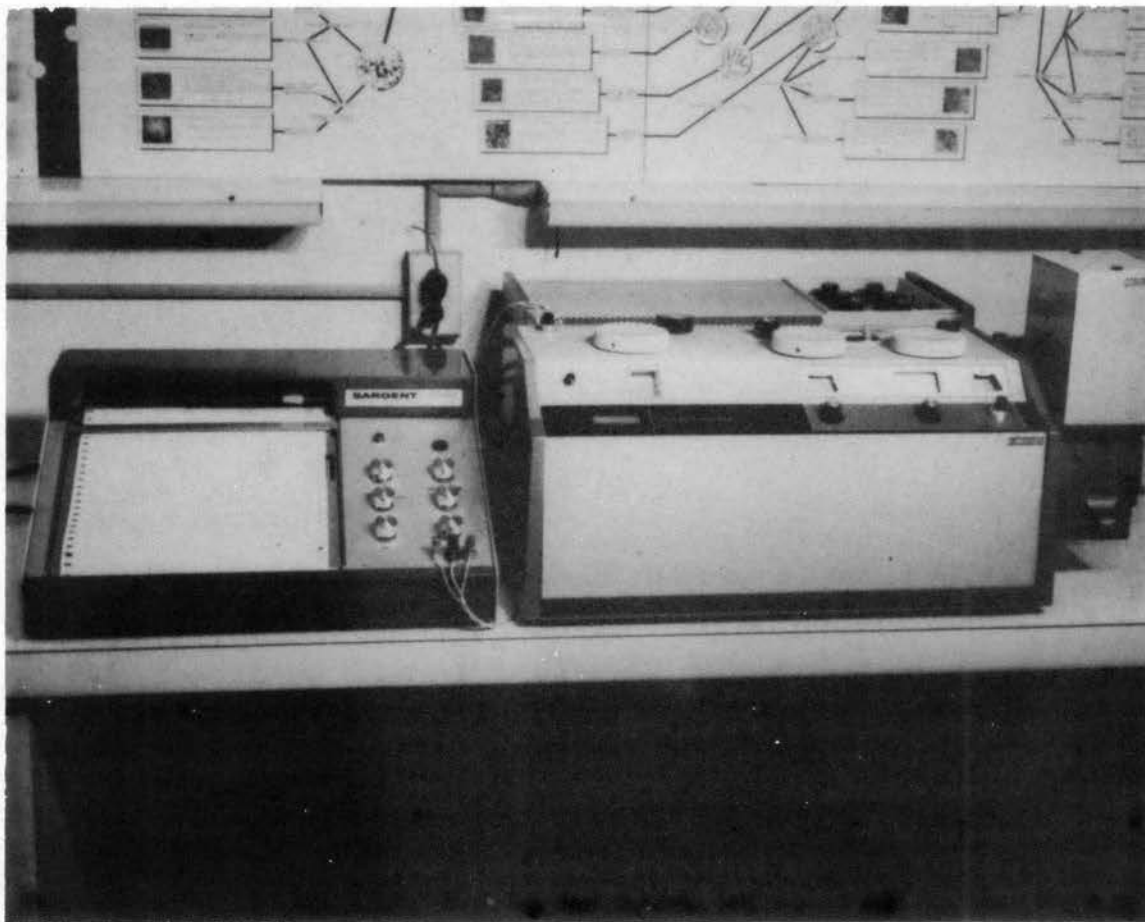


Figure 9. Spectrophotometer Attached to a Chart Recorder

- (1) Radiation sources - usually hydrogen arc for ultraviolet range, the incandescent tungsten filament for the visible range, and nernst glower for the infrared range.
- (2) Intensity control - a variable slit controlling the intensity of radiant energy falling onto the specimen and the standard.
- (3) Wavelength control (Monochromator) - a grating or prism which isolates a narrow spectral band from a continuous source.
- (4) Sample holder - container for the sample and the reference standard.
- (5) Receptor - a detector, such as a photocell, which measures the radiant energy reflected by the sample and the reference.
- (6) Indicator - a recorder or simple galvanometer.

Reflectance data is obtained with a spectrophotometer by directing radiant energy of various portions of the electromagnetic spectrum to strike a specimen and then determine what portion of energy is reflected by the specimen. The radiant energy is provided by a hydrogen lamp in the 190 to 350 millimicrons wavelength region, while a tungsten lamp provides the energy in the wavelength range extending from 350 to 1000 millimicrons (9).

CHAPTER V

INVESTIGATIVE PROCEDURE FOR EMITTANCE OF SOIL

General

Soils with different chemical compositions would be expected to emit radiant energy at different magnitudes depending upon their compositions. Since the emittance of a substance is best measured at the higher wavelengths of the electromagnetic spectrum, a PRT-5 Radiometer was used in this portion of the investigation.

The measuring of emissivity of soils consisted of two tests. Both tests were run on silty clay and on sand. The first test determined the effect of salinity on soils as moisture is evaporated by solar energy. The second test, consisting of two parts, determined the effect of changing levels of water table on emissivity of soils (a) with non-saline water, and (b) with saline water.

Description of Radiometer

The PRT-5 Radiometer measures the emissivity of a substance between 8 to 14 micron wavelengths. It is a completely portable, non-contact, direct-reading, temperature-measuring instrument which consists of an optical unit and electronic unit (Figure 10). The optical unit continuously compares the amount of energy emitted by the target with that emitted by an internal, controlled, reference environment. The

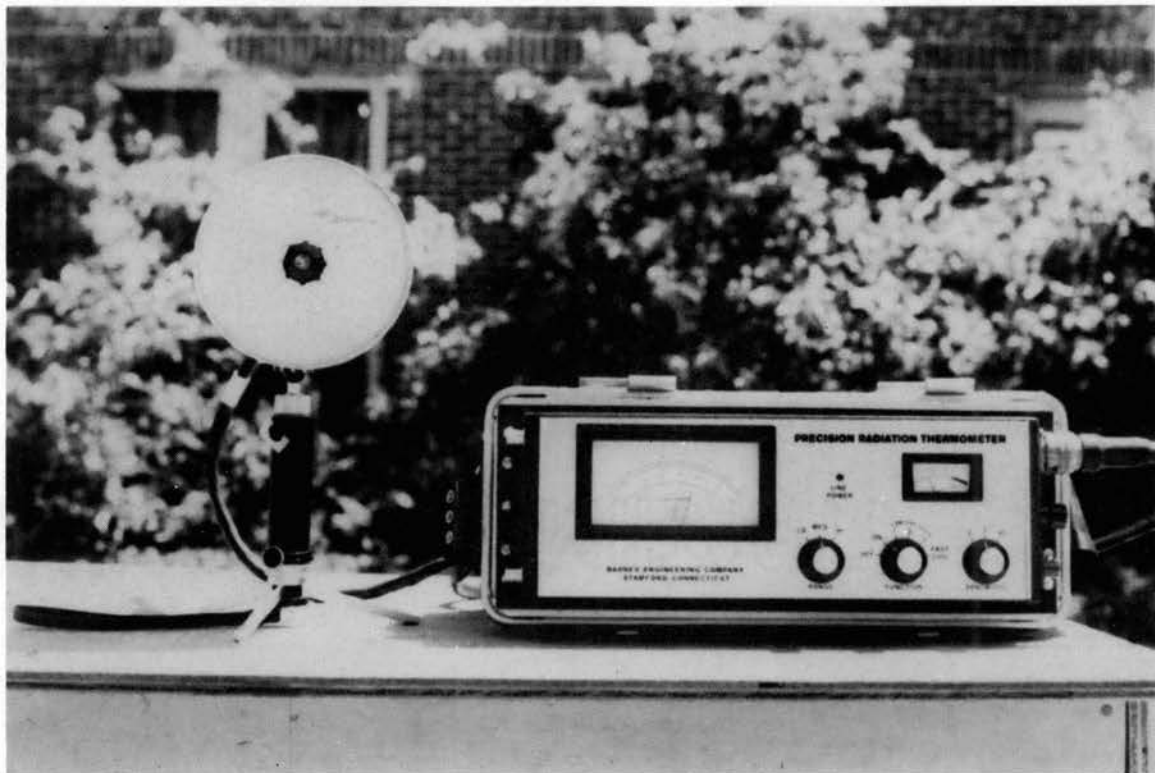


Figure 10. PRT-5 Radiometer

electronic unit processes this comparison into a voltage which is directly related to the energy difference between the target and the reference (1).

The PRT-5 gives precise readings of equivalent blackbody temperature from -20°C . to $+75^{\circ}\text{C}$. Readings are accurate down to $.5^{\circ}\text{C}$. The optical unit has a field view of 2° and a spectral filter which limits the measurement to the 8 to 14 micron atmospheric window.

Preparation of Panels

Panels, one-foot square by one-half foot deep, were made from 3/8-inch plyboard to contain the soils for the first test. The panels were sealed on the inside with epoxy to make sure the plyboard would not soak up any moisture from the soil. Soils used for testing were dried, sieved and placed in thirty-gallon galvanized barrels for storage. The salts to be used in this investigation were calculated, using the same procedure as described in Chapter IV, for the levels of salinity of 0, 20,000, 40,000, 60,000, 80,000, 100,000, 130,000 and 160,000 ppm in solution. These amounts of salts were then weighed and dissolved in the proper amounts of water and stored in five-gallon plastic containers. The panels were first weighed empty, then eight of them were filled with 25 pounds of silty clay in each panel. Another eight panels were filled with 30 pounds each of sand. The different levels of salt solutions were then added to each panel of soil and mixed into the soils by hand. The panels were then reweighed and placed outdoors on tables (Figure 11). The weights of the panels were recorded in order to determine the amount of moisture in the soils.

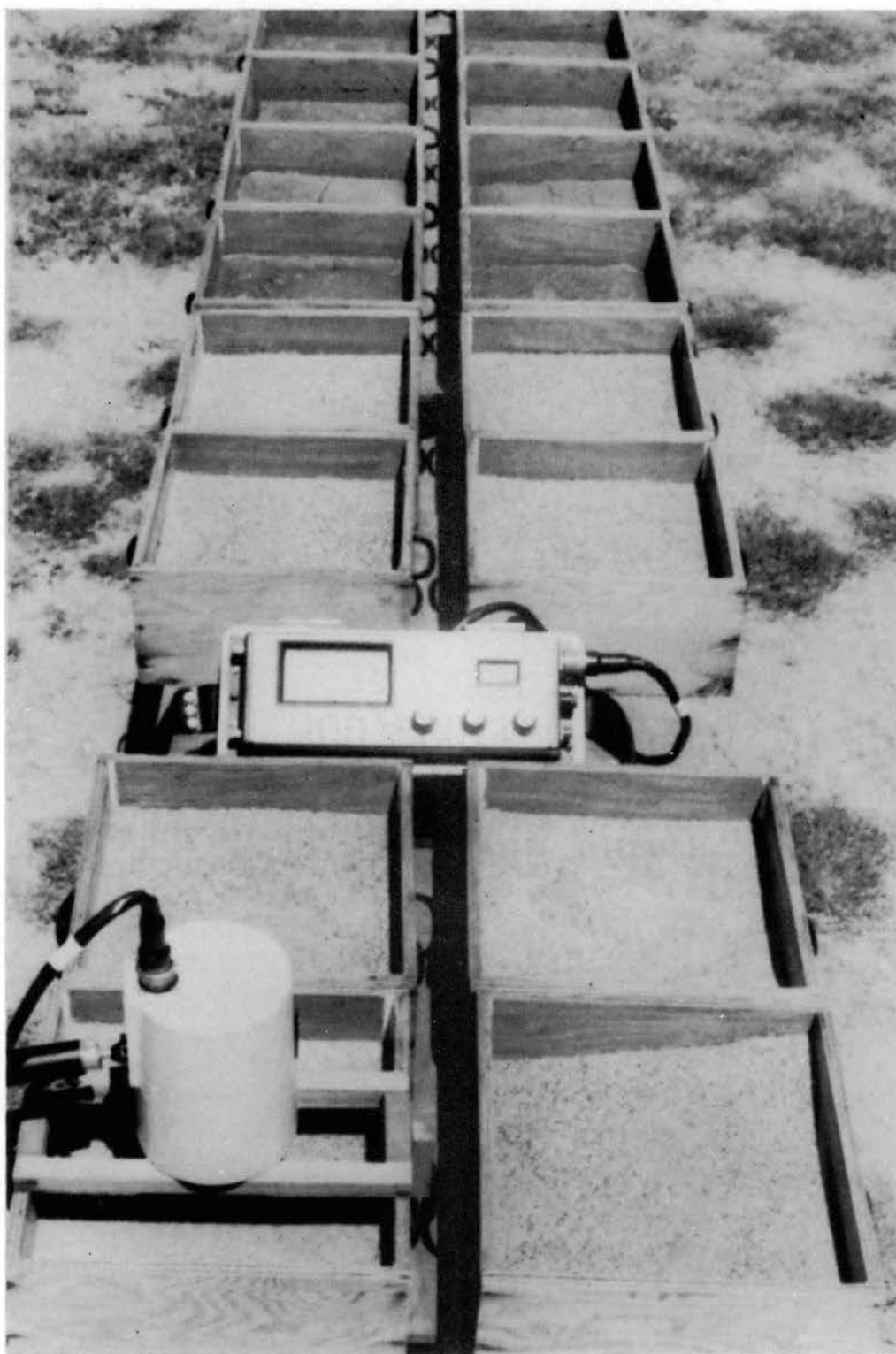


Figure 11. Soil Panels

Procedure for Measuring Emittance

The panels were left outdoors overnight before readings were begun allowing the soils temperature to come into equilibrium with the ambient temperature. Readings were taken with the radiometer for at least one week with eight readings taken during every 24-hour period. The ambient temperature and meteorological changes were also noted and recorded at every time of reading. Each day at 4:00 p.m. the panels were reweighed in order to determine the change in moisture. Care was taken to keep the optical unit of the radiometer at the same distance from the soils for each reading.

The PRT-5 Radiometer gives readings in equivalent blackbody temperature. From the graph shown in Figure 12 the emitted energy was determined and is given in the results of Chapter VI.

Description of Water Table Apparatus

Since the surface temperature of soils is affected by the temperature of soils beneath them, the changing water table was thought to affect the emissivity of the soil. Therefore, an apparatus was designed and built to simulate changing water table in a laboratory test. The apparatus consisted of a 10-inch square inner box, 4 ft.-2 in. high, to hold the soil, and a 17½-inch square outer box, 4-feet high, to hold the water (Figure 13).

The outer box had a 20-inch square base plate affixed to the bottom with built-up strips 10-inches square for the inner box to rest in. The inner box was made in 1-foot sections for ease of cleaning and tearing down after each experiment was completed. The inner box had 12 quarter-inch holes, three on each side, at every one foot level to allow passage

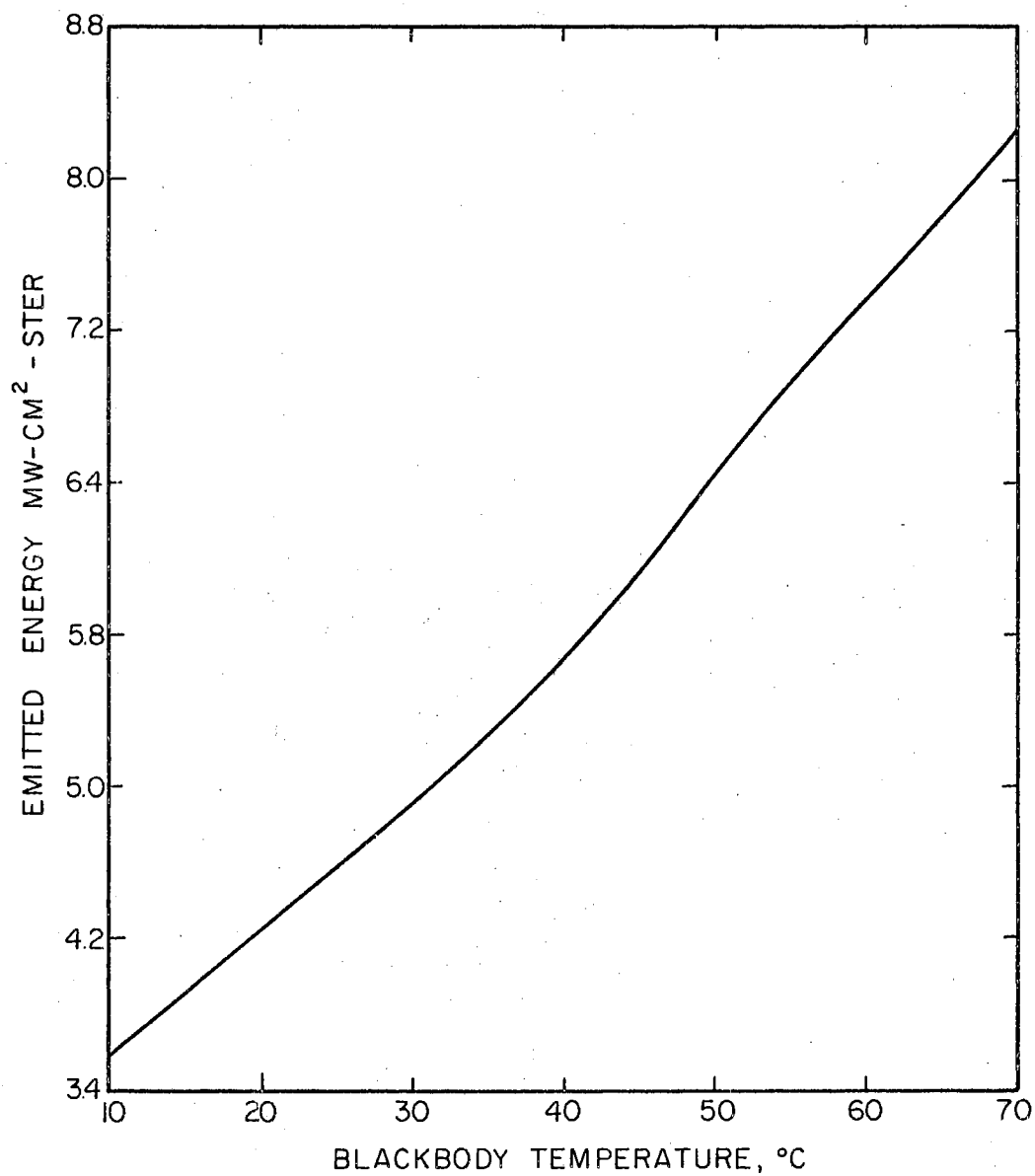


Figure 12. Effective Emitted Energy for
Standard PRT-5 Radiometer

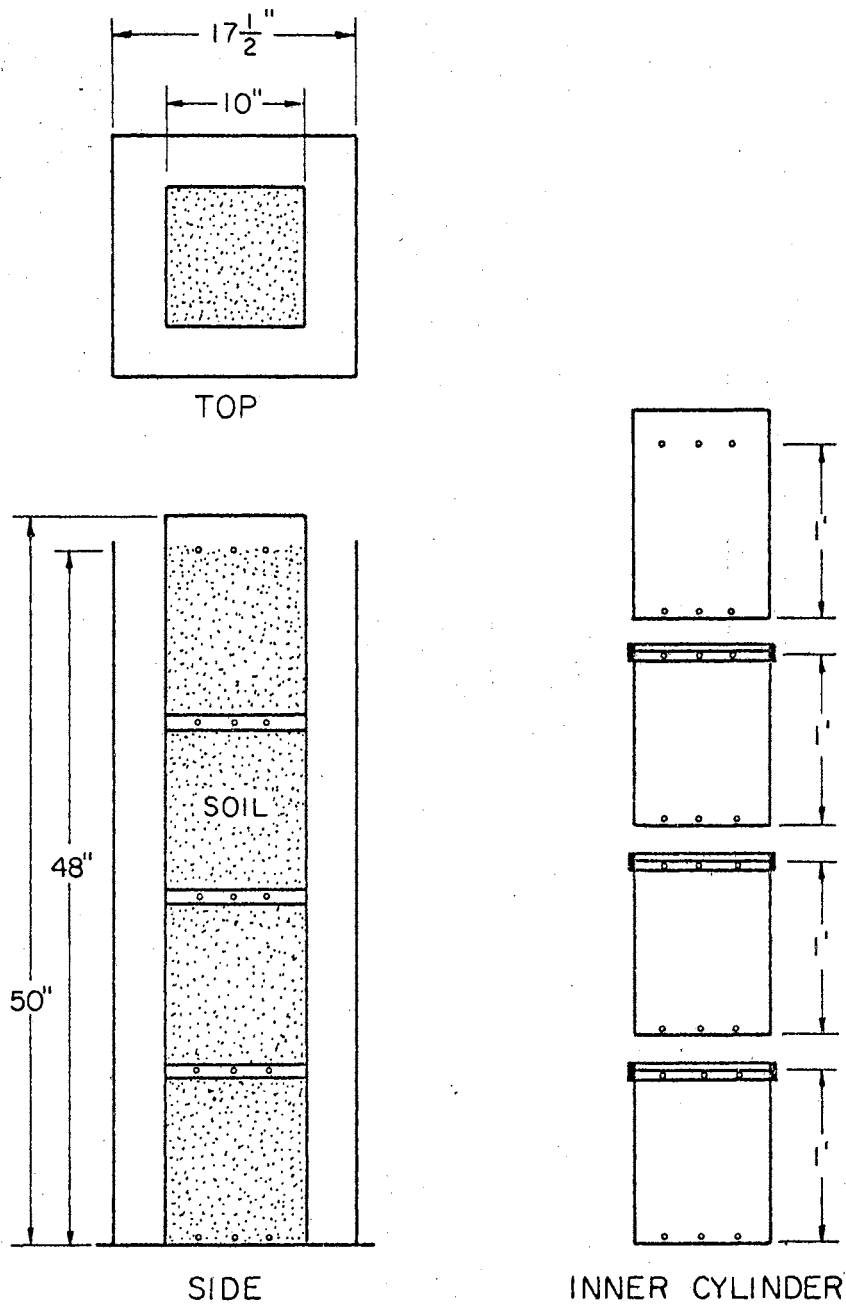


Figure 13. Water Table Apparatus

of the water into the soil as the water level changed. The apparatus was constructed from Plexiglas for ease of making and having a water tight system.

Preparation of Test for Effect of Water Table

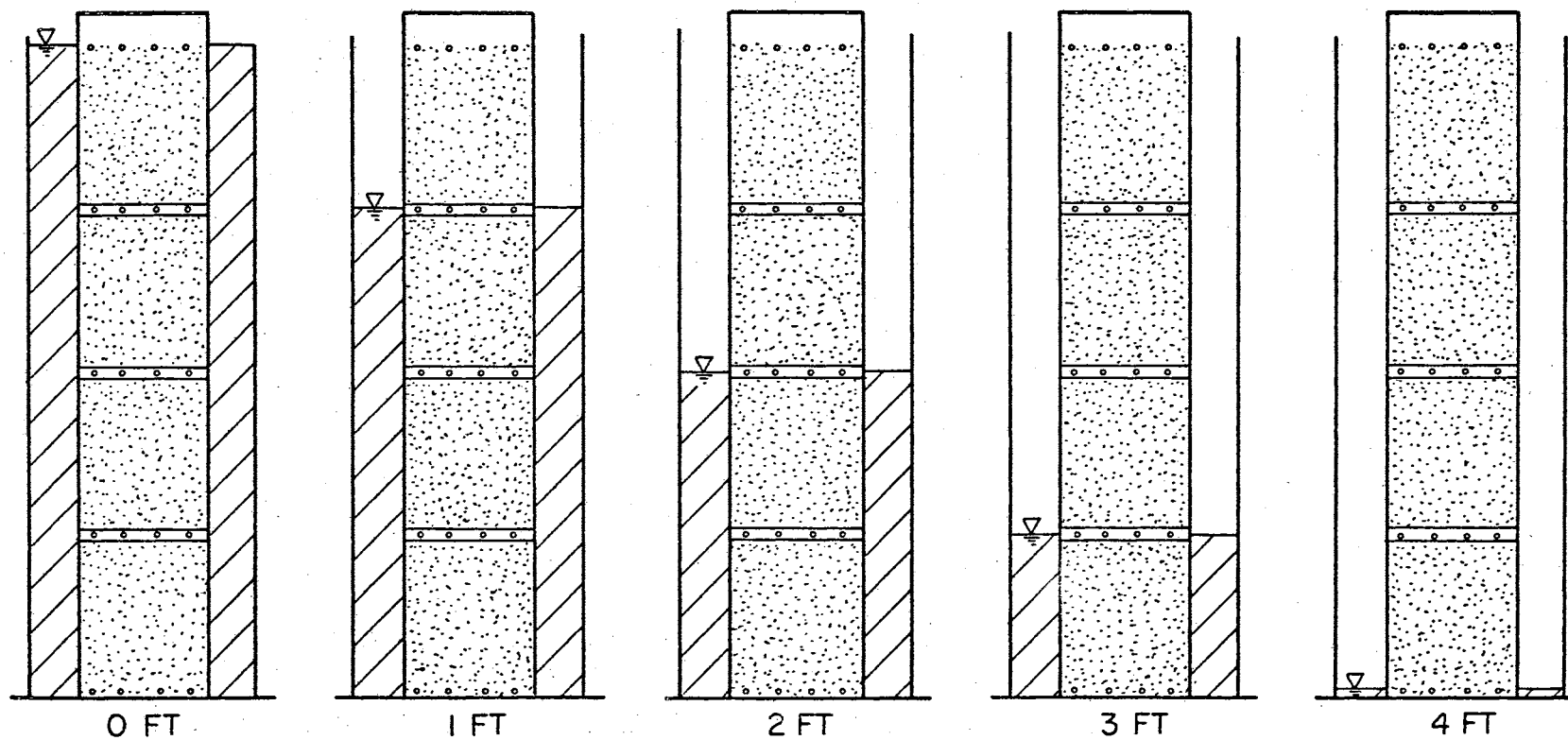
A. Non-Saline Water

The soils were again dried, sieved and stored in 30-gallon galvanized barrels. The boxes were placed outdoors and the inner boxes were filled with soil. When placing the soil in the inner box it was tamped every six-inch lift in order for all five boxes to have approximately the same density. This was done to eliminate the variable of density of soil on its emissivity. The outer boxes were then filled with tap water to the 0 ft., 1 ft., 2 ft., 3 ft., and 4 ft. levels (Figure 14).

This same procedure was followed with both sand and silty clay soil. The sand was dried and used again in part B of the investigation. The silty clay was dried, pulverized and sieved again to be reused in part B.

B. Saline Water

Salt solutions for 0, 20,000, 60,000, 100,000 and 160,000 ppm were prepared and stored in five-gallon containers. Forty litres of each level of salinity were prepared at a time. The soil was placed in the inner boxes in the same manner as was done in part A. The salt solutions were then put in the outer boxes. Only enough salt solution was put in all five boxes to have the depth to water table at the 4 ft. level.



WATER TABLE

Figure 14. Water Table Levels for Non-Saline Water Table Test

Procedure for Measuring Emittance

The boxes were left outdoors overnight before readings were begun to allow the soil temperature to come into equilibrium with the ambient temperature and all the capillary rise to occur. The readings were taken in the same manner as was used in measuring the emittance of soils in the panels (Figure 15).

When evaporation of the water brought the water level below the desired level, water was added to the outer box. The water table for the non-saline water test was kept at the same level as shown in Figure 14 throughout the test. The water table for the saline water test was raised one foot every two days. When raising the water level, the saline solutions were added at 12:00 midnight so soils could become saturated before readings were taken again at 6:00 a.m. The boxes were covered with plastic sheets when it rained to keep the soils from becoming wetter than would be natural from the levels of the water table and diluting the saline water.

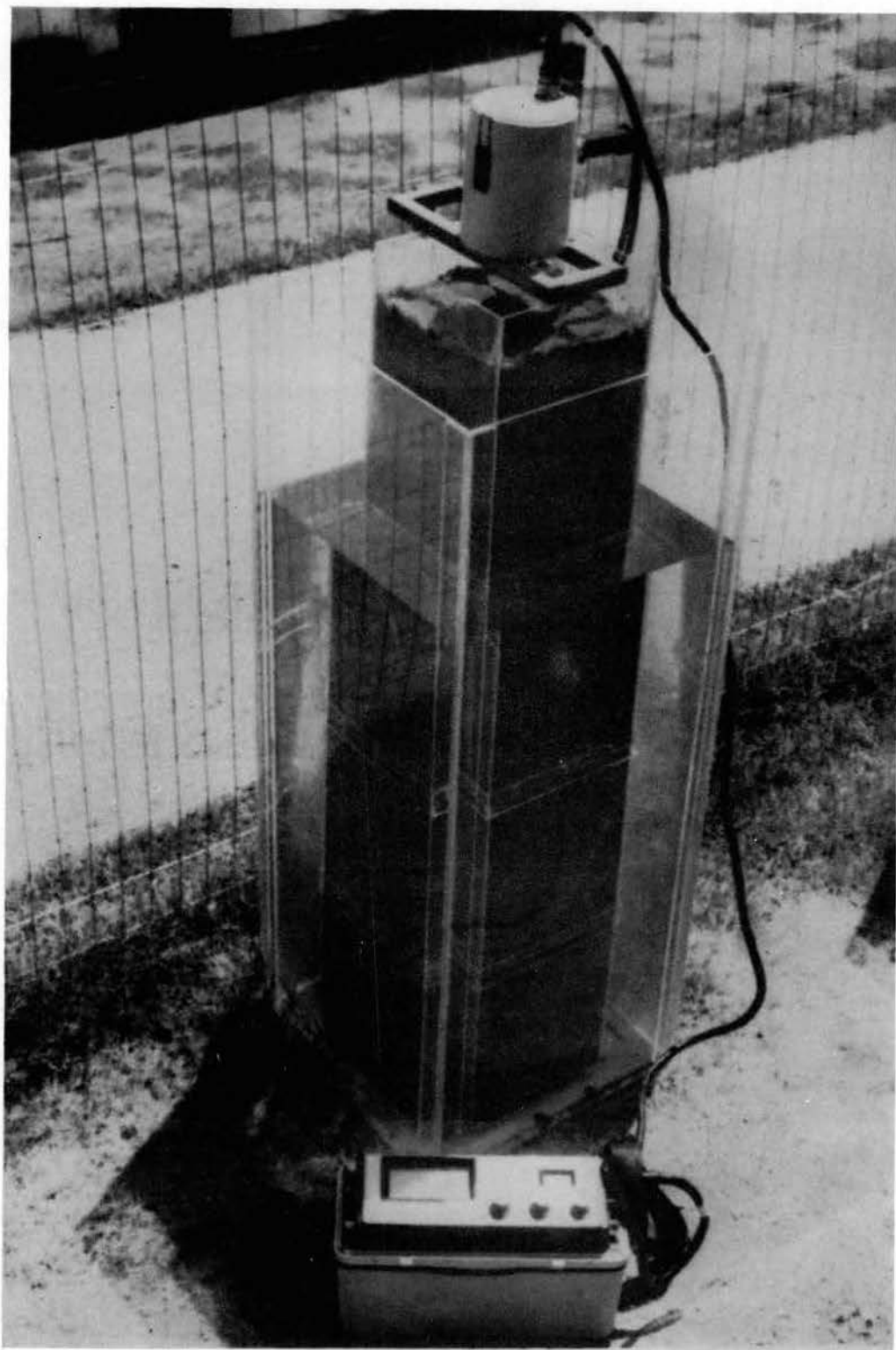


Figure 15. Water Table Apparatus

CHAPTER VI

DISCUSSION AND RESULTS

Effect of Salinity on Reflectance of Vegetation

For testing the effects of salinity on reflectance of vegetation, the wheat and salt grass were grown in soils containing varied salt concentrations. These test results are shown in Figures 16 through 19 and Appendix A. Since there are certain materials in both the wheat and salt grass leaves which affect spectral reflectance in the same manner, they have the same general pattern of reflectance throughout the electromagnetic spectrum tested in this study (15). Both types of vegetation had their lowest reflectance at the end of the ultraviolet portion of the spectrum, that is at 350 millimicrons. This is the violet and blue light portion of the spectrum which is largely absorbed by chloroplasts and used in the photosynthesis process. The chloroplasts of the leaves were unaffected by the salts to the point that the reflectance for any one wavelength did not change for the different levels of salinity.

Between 200 and 650 millimicrons the reflectance for both types of vegetation remained approximately the same, with only a slight variance at 550 millimicrons. At this point, the reflectance for both types had increased, but the variance due to the effect of different levels of salinity was greater for the wheat than that for the salt grass. This

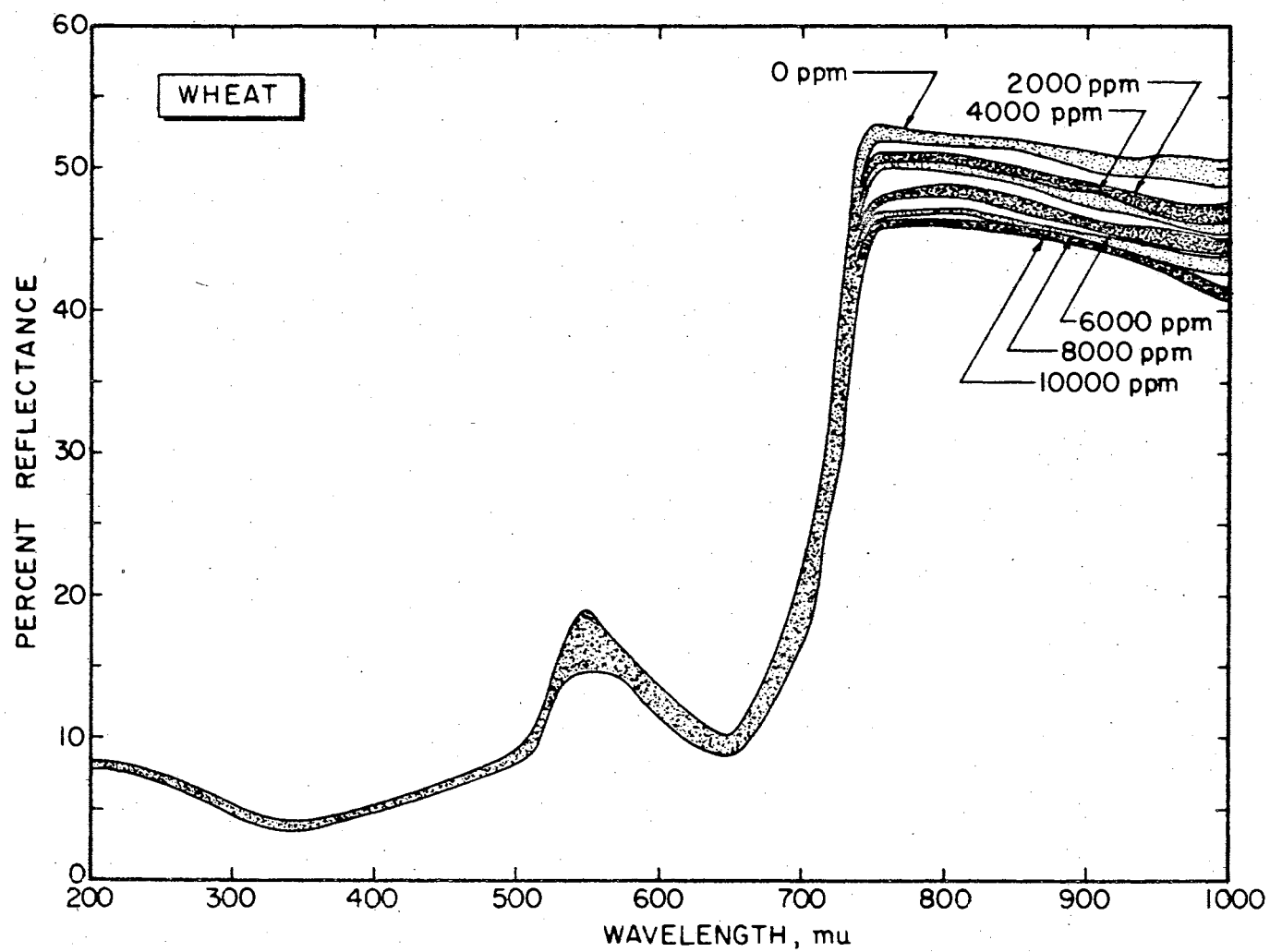


Figure 16. Plot of High and Low Reflectance of Wheat Grown in Varied Salt Concentrations

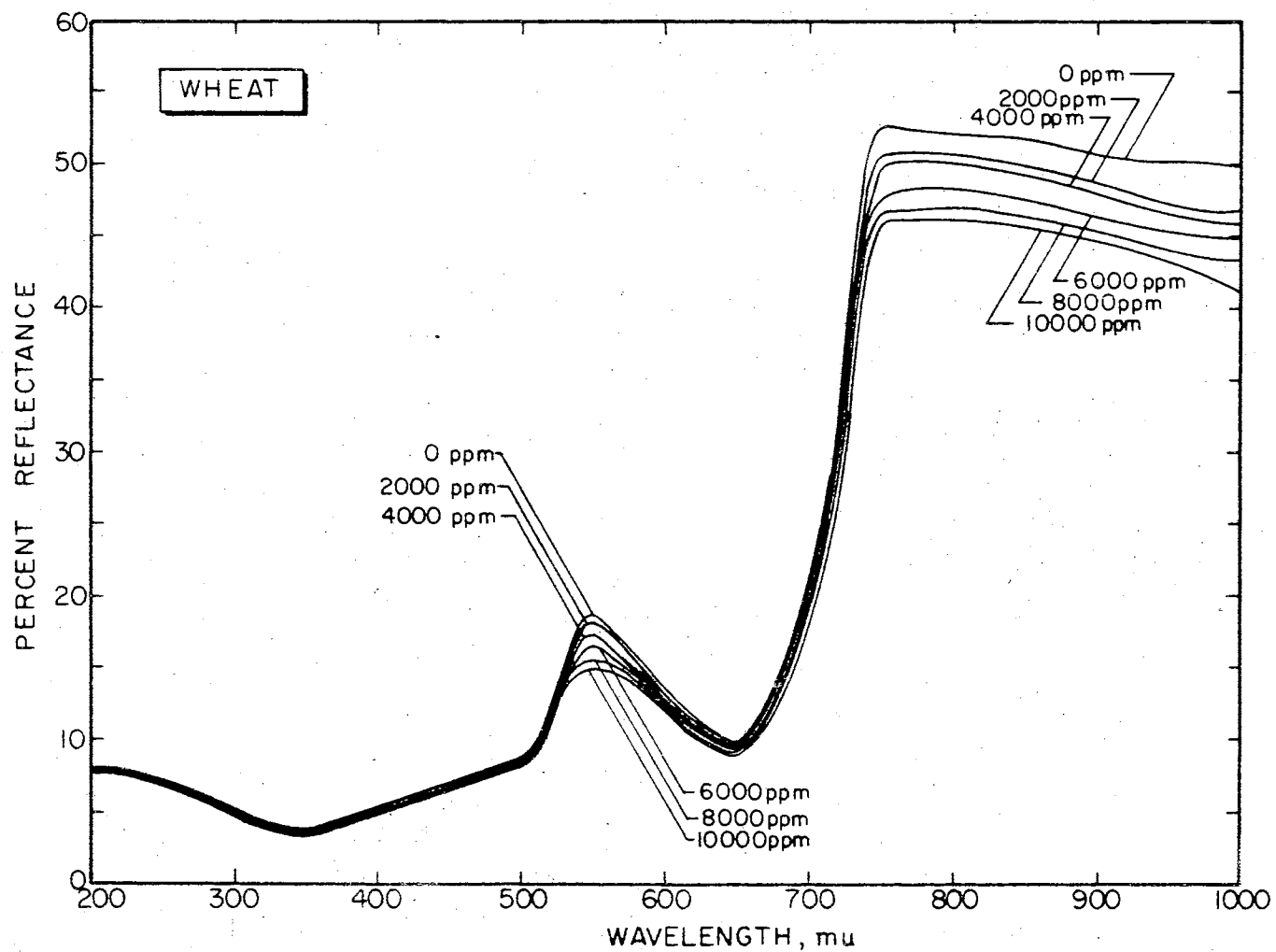


Figure 17. Average of High and Low Reflectance for Wheat

increase in reflectance occurred in the green light portion of the spectrum where it is largely reflected by chloroplasts.

At 750 millimicrons the near infrared portion of the spectrum begins, at which point the infrared light is highly reflected by the spongy mesophyll of a leaf's structure. For this reason, between 650 and 750 millimicrons both the salt grass and wheat reflectances sharply increased.

From 720 to 1000 millimicrons the reflectance stayed fairly constant, between 40% and 50%. It was within this region that the maximum difference for the varied levels of salinity was attained for both the salt grass and wheat. This is believed to have occurred when the spongy mesophyll of the leaf structures began collapsing due to the salts in the soils causing the leaves to reflect less near infrared energy (18). It could then be seen that the greater the salt concentrations, the greater the collapse of the mesophyll and more absorbent the leaves became to the near infrared light.

The relative turgidity for the leaves grown at different levels of salinity was determined for both types of vegetation. These results are given in Appendix A. The variance is from 90% to 98.5%. Since it would have to be less than 80% to affect the reflectance of the leaves, relative turgidity was not a factor in determining the reflectance of the vegetation used in this investigation (21).

As can be seen from Figures 18 and 19, the vegetation did not grow in soils containing high salt concentrations. The salt grass would not grow in silty clay having salt concentrations above 60,000 ppm, nor in sand having salt concentrations above 100,000 ppm. At these levels the salts became toxic to the plants.

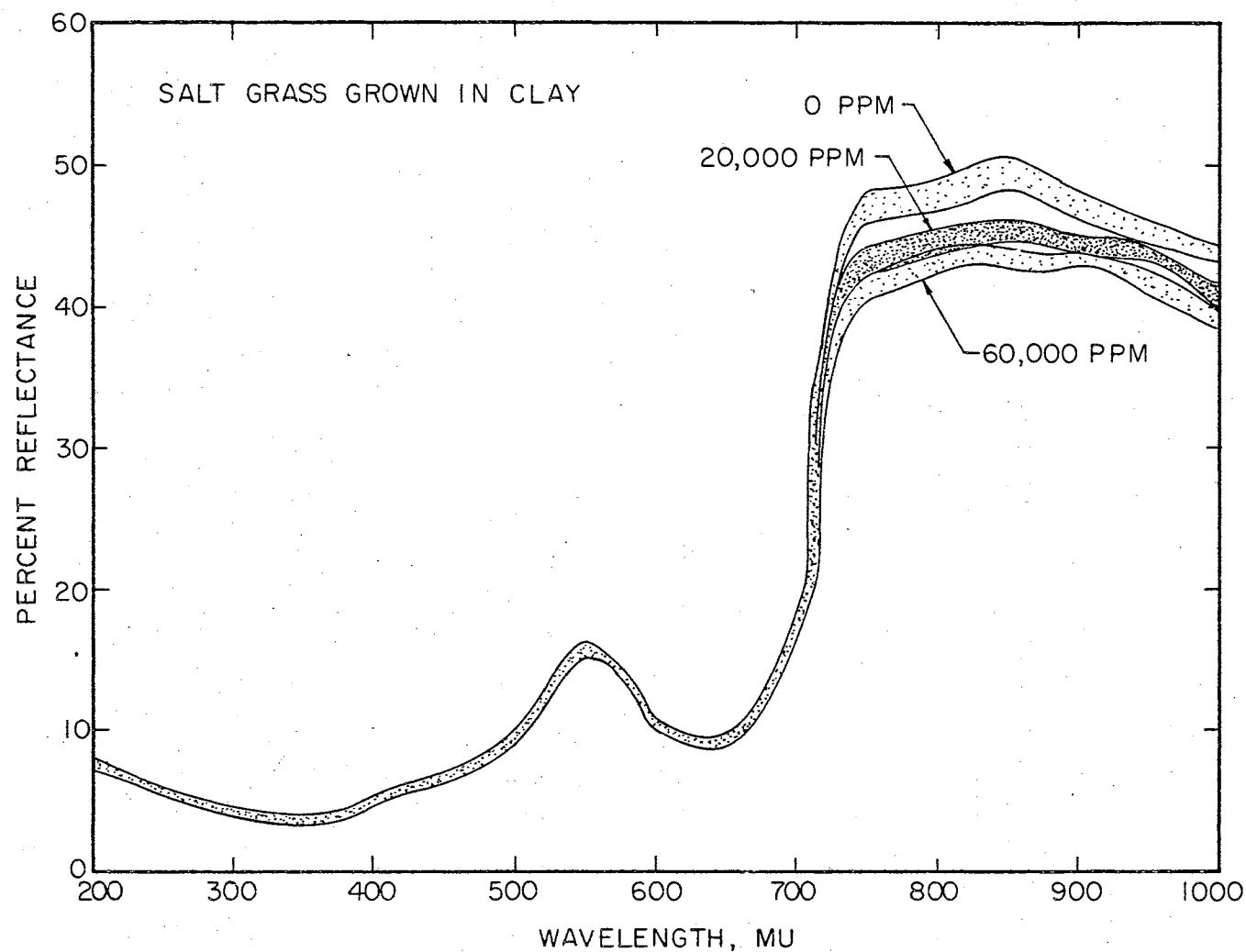


Figure 18. Plot of High and Low Reflectance of Salt Grass Grown in Silty Clay at Varied Salt Concentrations

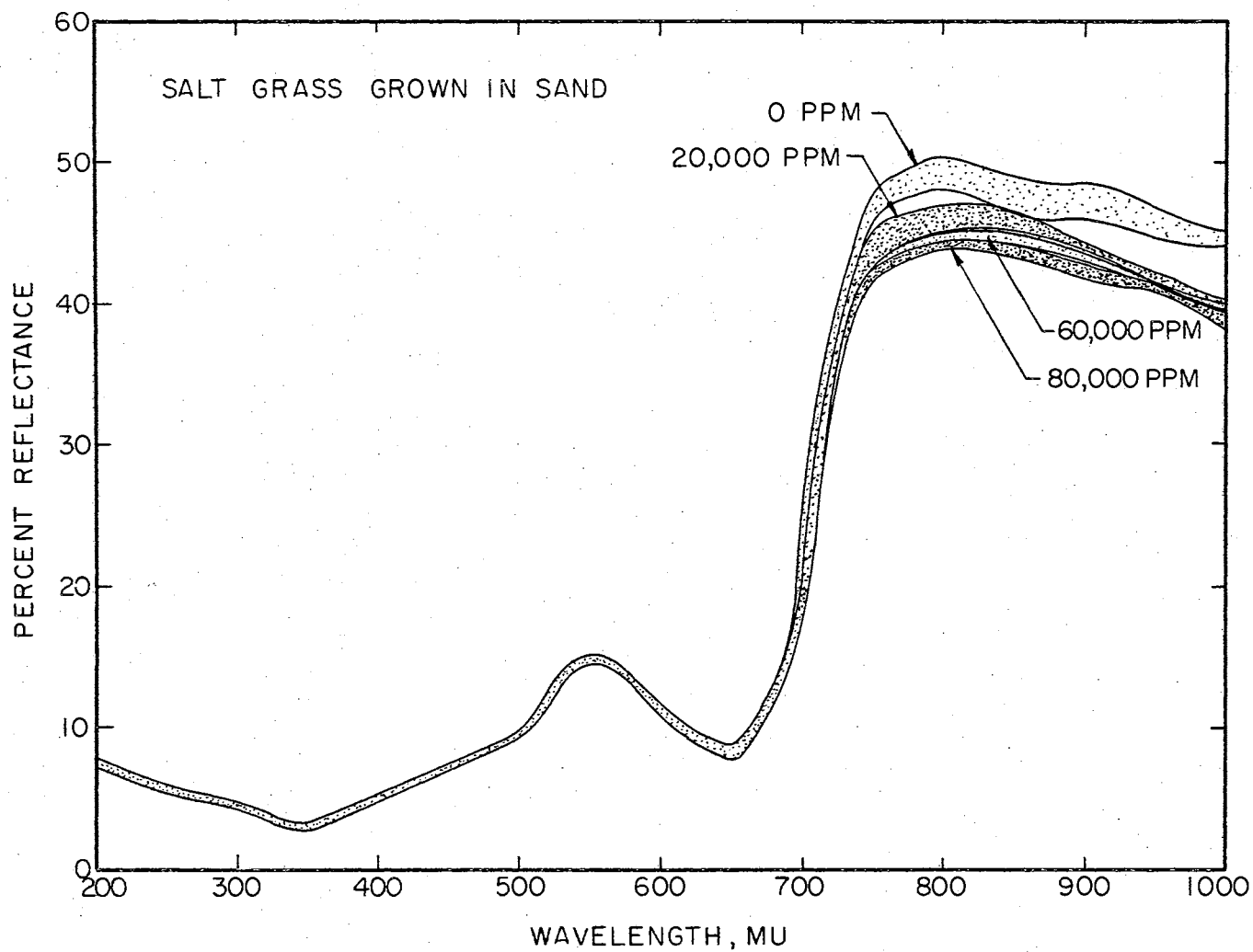


Figure 19. Plot of High and Low Reflectance of Salt Grass Grown in Sand at Varied Salt Concentrations

Infrared photographs were taken of samples of wheat leaves with a Kodak Retina IV camera, using an Ektachrome IR film with No. 12 and No. 25 wratten filters. The Ektachrome IR film is sensitive to light beginning with the visible light up to 860 millimicron wavelengths. The wratten filters block out all of the electromagnetic energy for certain portions of the electromagnetic spectrum; at the same time, they are transparent to the rest of the spectrum and electromagnetic energy passes through the filters unaffected. The portion of the spectrum which is filtered out depends upon the type of filter used. The No. 25 filter filters out all light throughout the spectrum up to 590 millimicrons, the No. 12 filter filters out all light up to 500 millimicrons, except for a small window between 300 and 320 millimicrons. Therefore, with these particular film-filter combinations, the photographs measured the magnitude of reflectance of the wheat primarily between 500 to 860 millimicron wavelengths with the No. 12 filter and between 590 to 860 millimicrons with the No. 25 filter.

Photographs were taken indoors with two 300 watt tungsten lamps as energy sources. These photographs were taken with different focal lengths and exposure times with best results obtained at a focal length of f/16 and exposure time of 1/60 for the No. 12 filter, and at a focal length of f/16 and exposure time 1/125 for the No. 25 filter. The photographs depicted a difference in color for the varied levels of salinity at which the wheat was grown. Since Ektachrome IR film is a false color film, the leaves photographed were varying shades of red. The wheat grown in the higher salt concentrations had the lightest shade of red, and the wheat grown in lesser salt concentrations had darker shades of red.

Emitted Energy of Soils with Variance in Moisture and Salinity

Since the temperature of soil is greatly affected by the amount of moisture it contains, and salts retard the evaporation of that moisture, a test was run to determine these effects on the emitted energy of sand and silty clay. The test was run from September 20 through 27, 1968. The data for the test are given in Appendix B. Figures 20 through 31 show plots for September 21-26. This same test was run in May 1969, and gave the same results as shown in these data.

The general trend for both silty clay and sand was for the highest emitted energy of the different levels of salinity to be in the 160,000 ppm salt concentrations and then decrease accordingly down to 0 ppm. Although some readings were erratic, this trend was predominant between 12:00 to 16:00 hours. The soils having higher salt concentrations emitted more radiant energy since the salts retarded evaporation of moisture in the soil. Therefore, the lesser salt concentrations allowed more evaporation of moisture causing the soils to be slightly cooled by the process (3). This was further verified by the moisture content in the soils which was determined each day of the investigation. (See Appendix B.) Since the soils were air dried they had small amounts of moisture in them before the salt solutions were added. The moisture content for Sept. 20 is the amount of moisture added to the soil with the salt solutions. However, the actual moisture content is greater than that shown, due to the small amount of moisture retained by soils when air dried and existent in them when testing began. The moisture content decreased about 26% for the 0 ppm in the silty clay and only about 9% for the 160,000 ppm. A similar occurrence took place in the sand with the

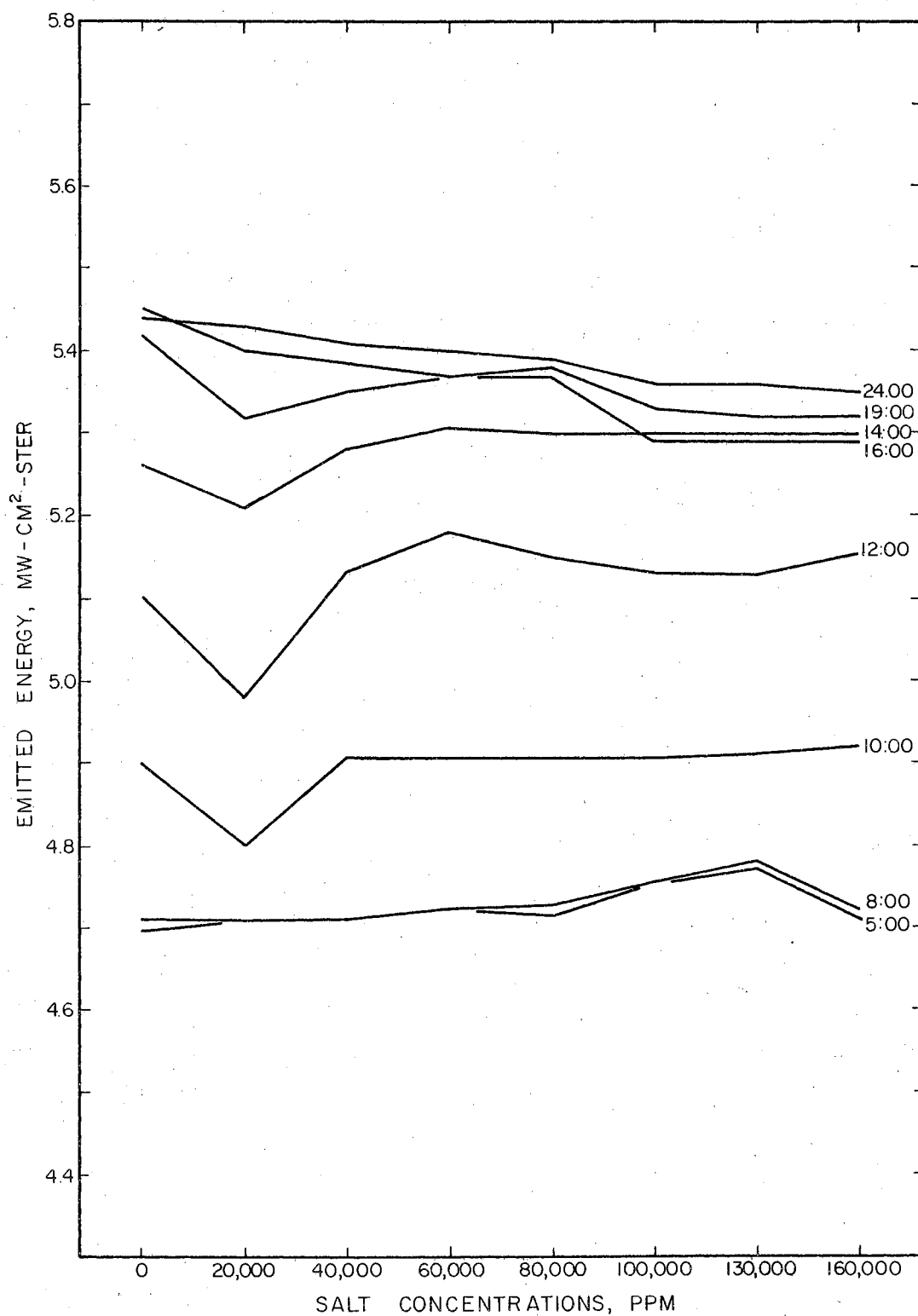


Figure 20. Emitted Energy of Sand with Variance in Moisture and Salinity for September 21

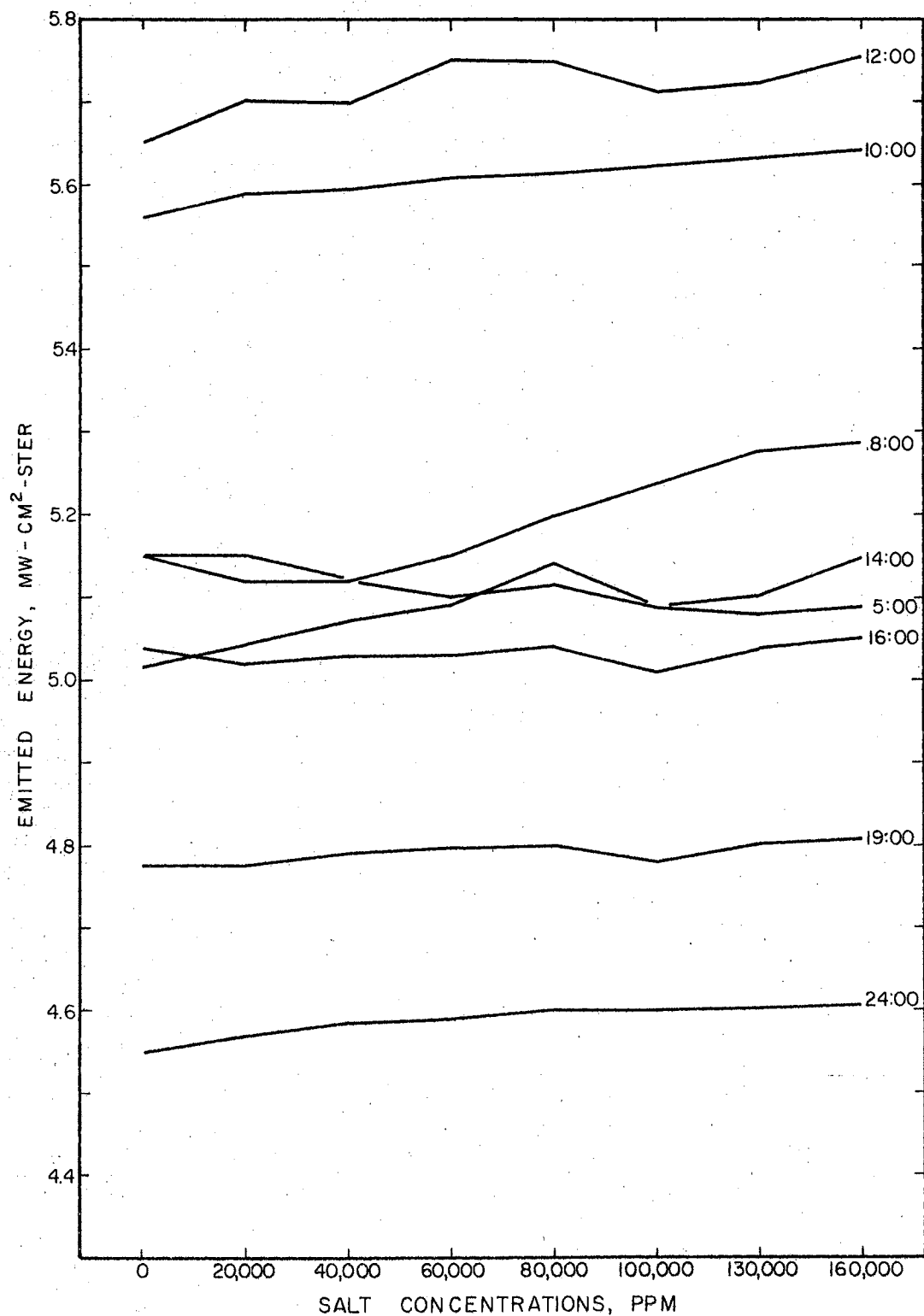


Figure 21. Emitted Energy of Sand with Variance in Moisture and Salinity for September 22

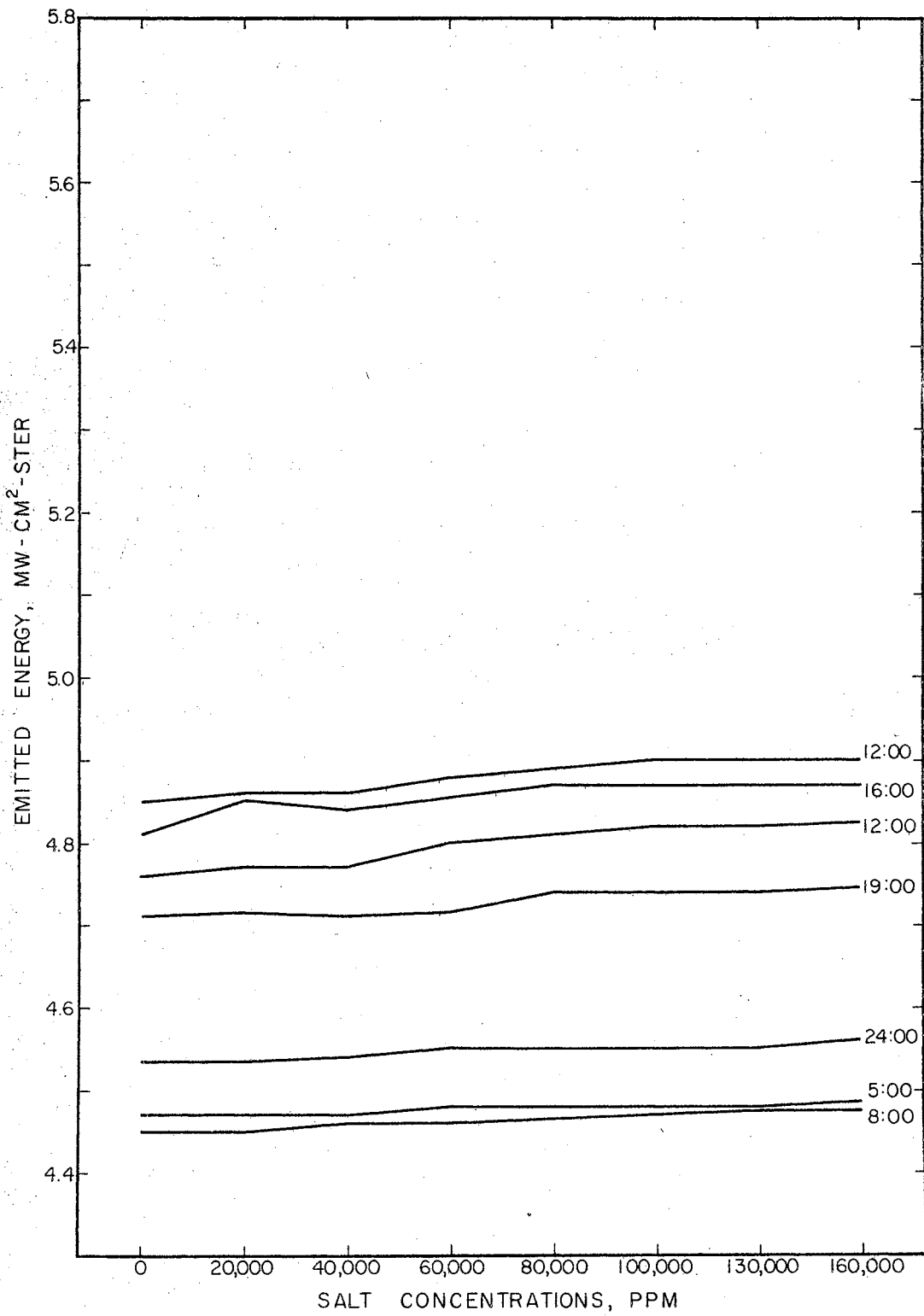


Figure 22. Emitted Energy of Sand with Variance in Moisture and Salinity for September 23

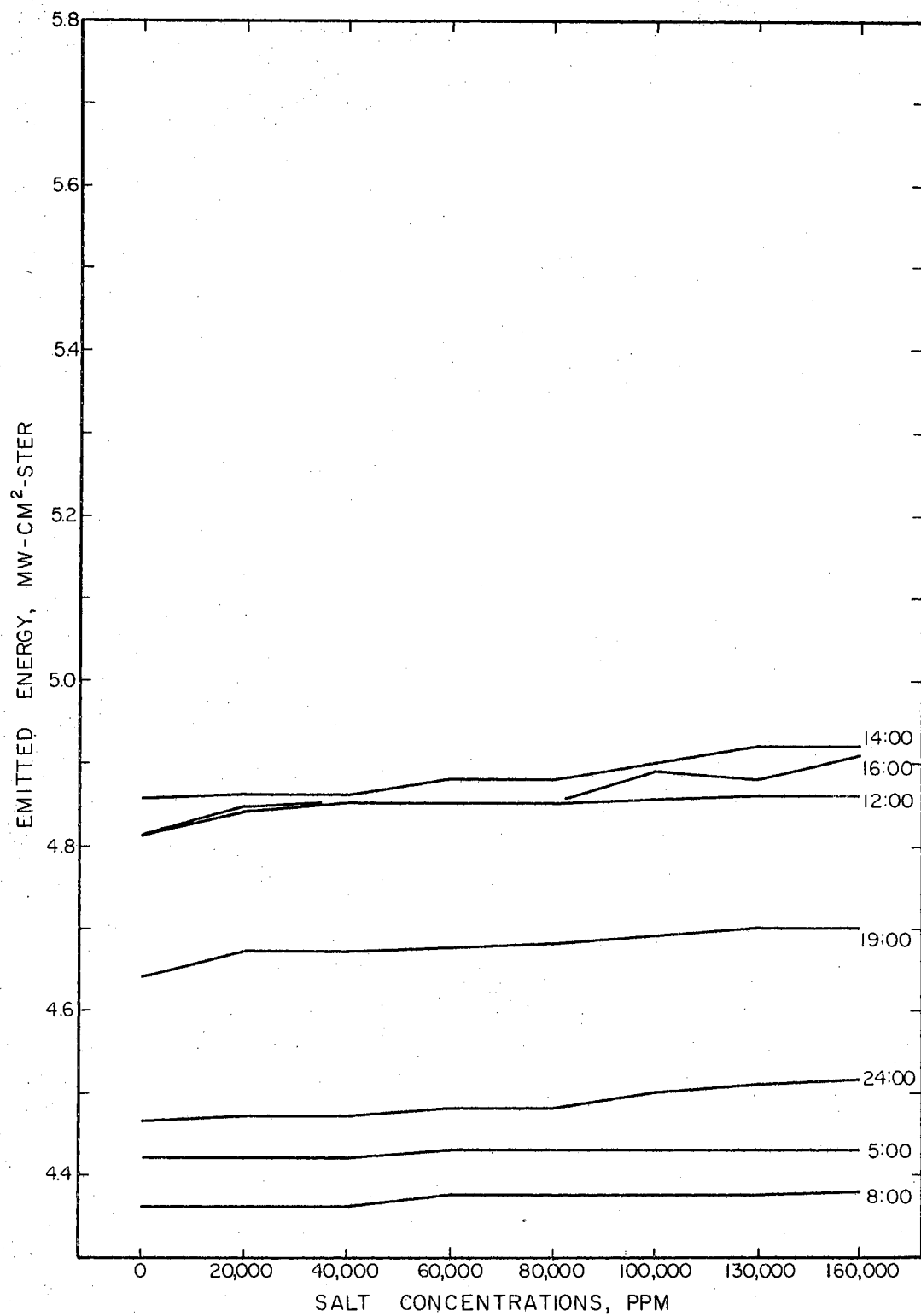


Figure 23. Emitted Energy of Sand with Variance in Moisture and Salinity for September 24

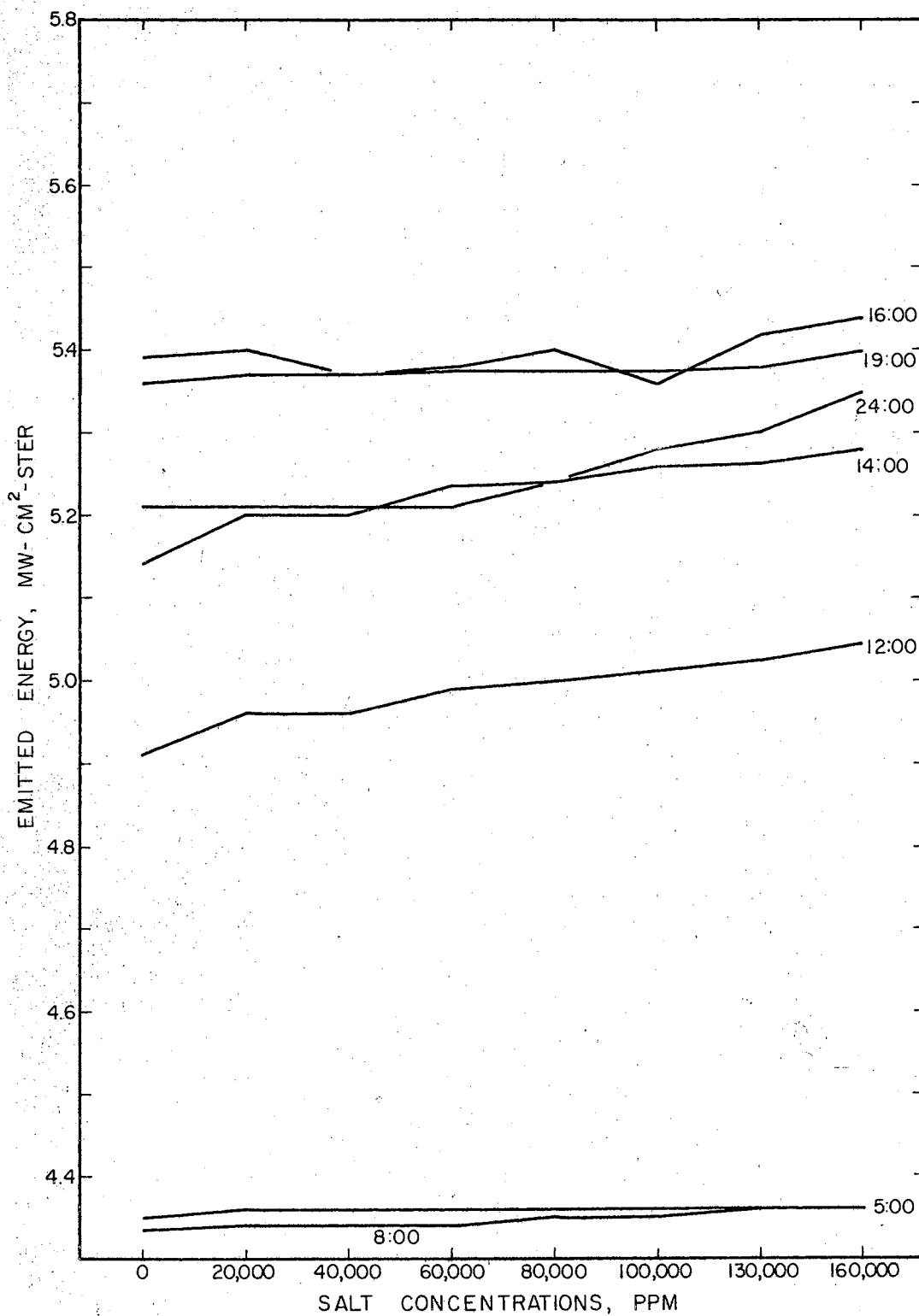


Figure 24. Emitted Energy of Sand with Variance in Moisture and Salinity for September 25

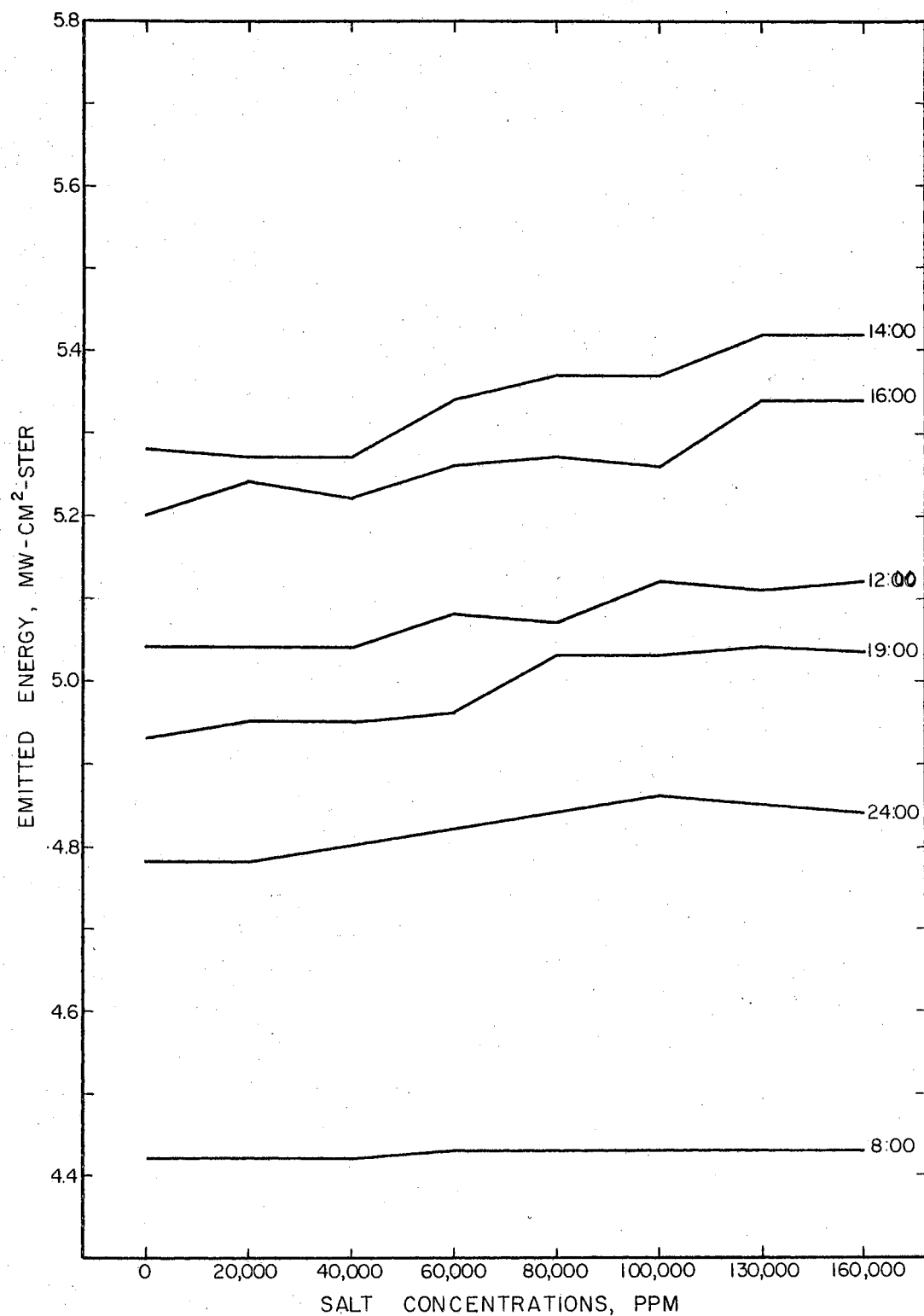


Figure 25. Emitted Energy of Sand with Variance in Moisture and Salinity for September 26

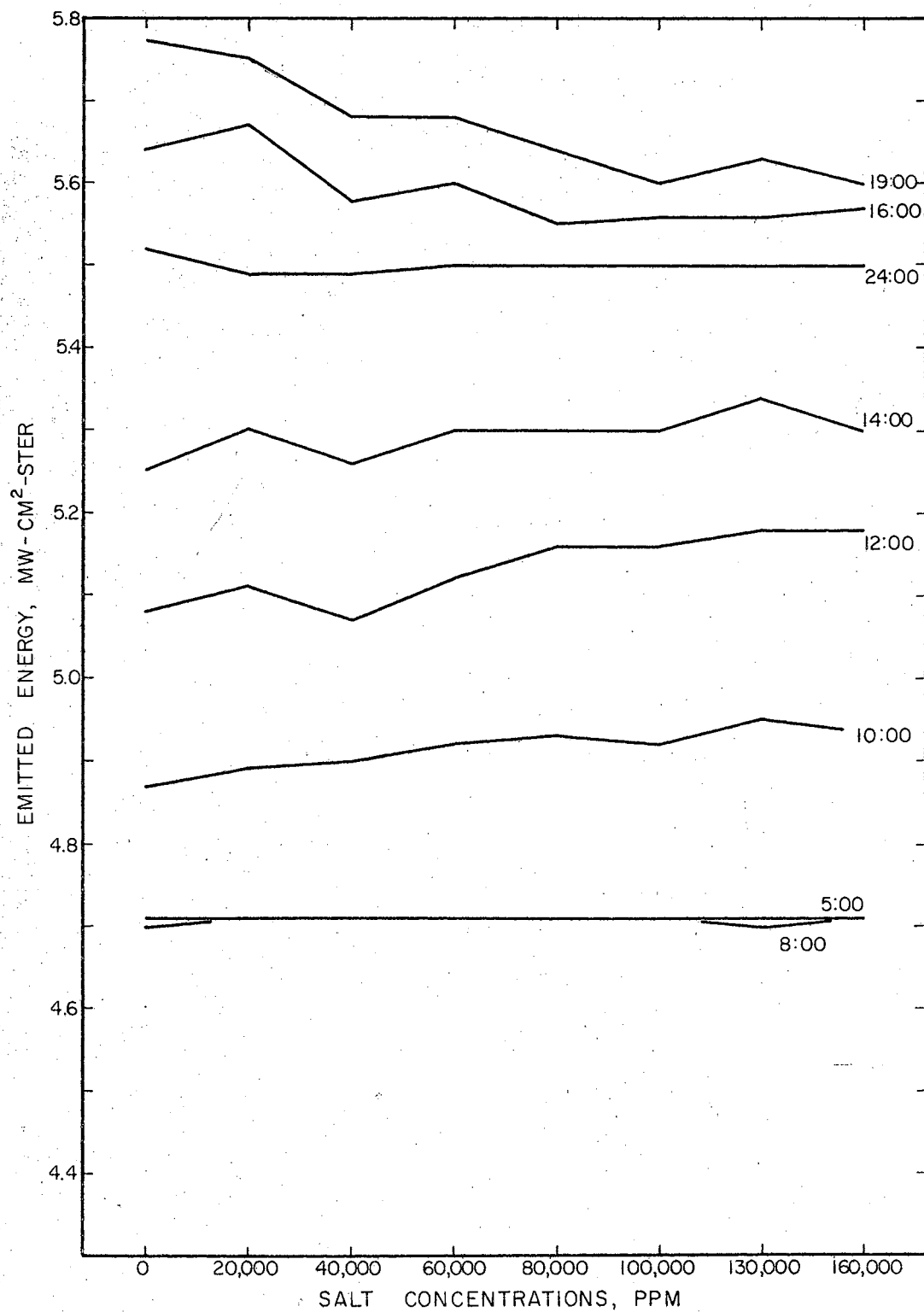


Figure 26. Emitted Energy of Silty Clay with Variance in Moisture and Salinity for September 21

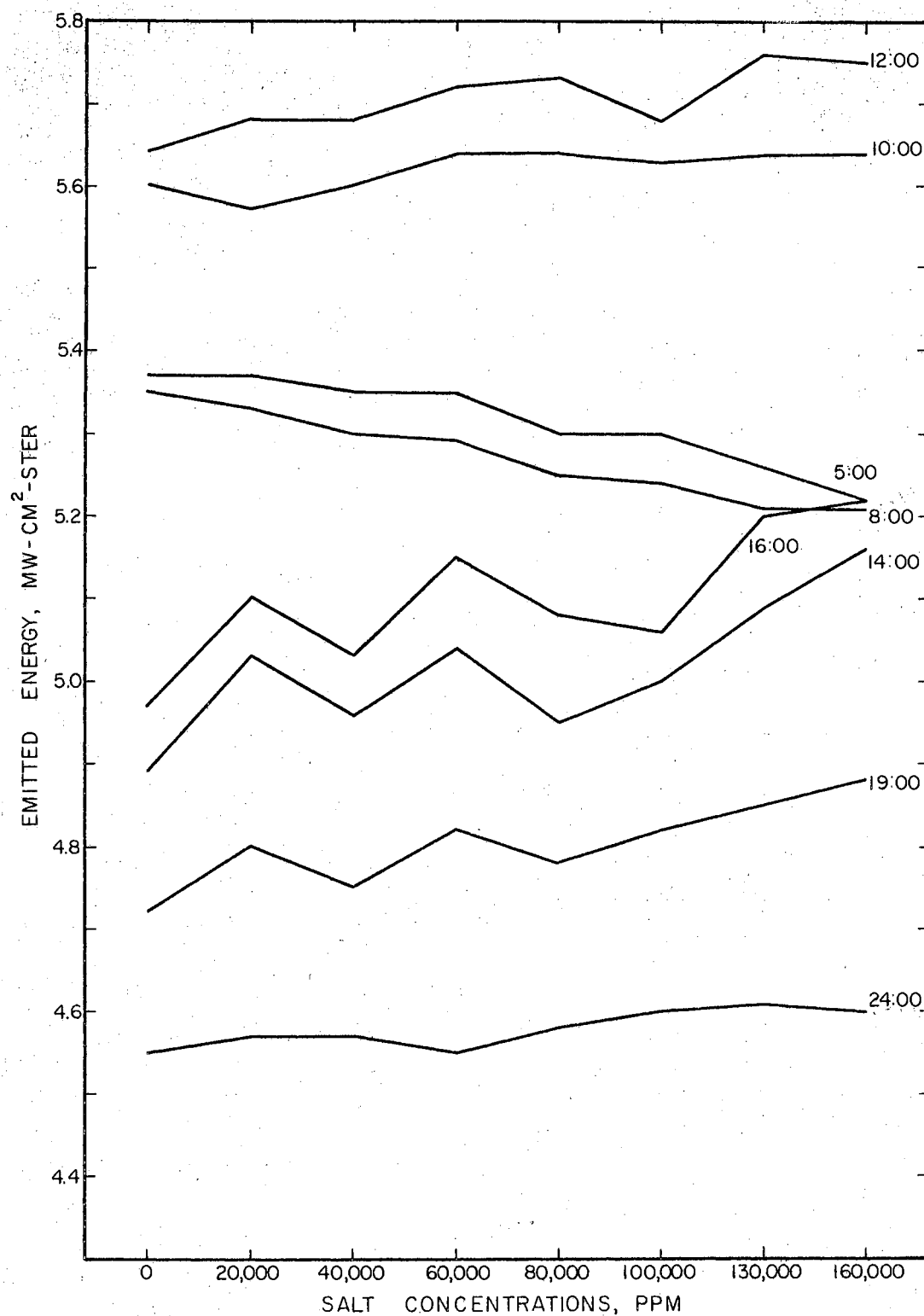


Figure 27. Emitted Energy of Silty Clay with Variance in Moisture and Salinity for September 22

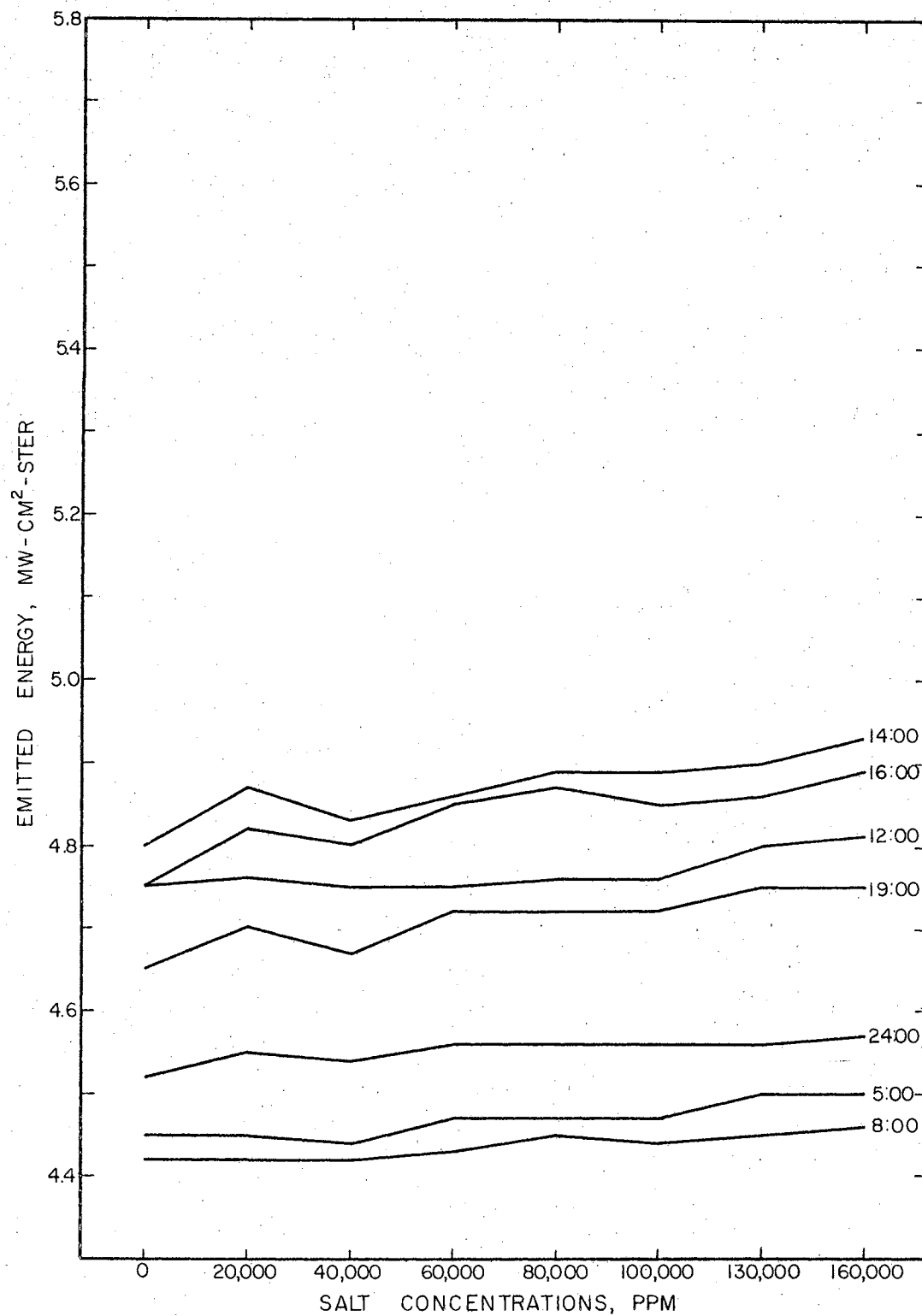


Figure 28. Emitted Energy of Silty Clay with Variance in Moisture and Salinity for September 23

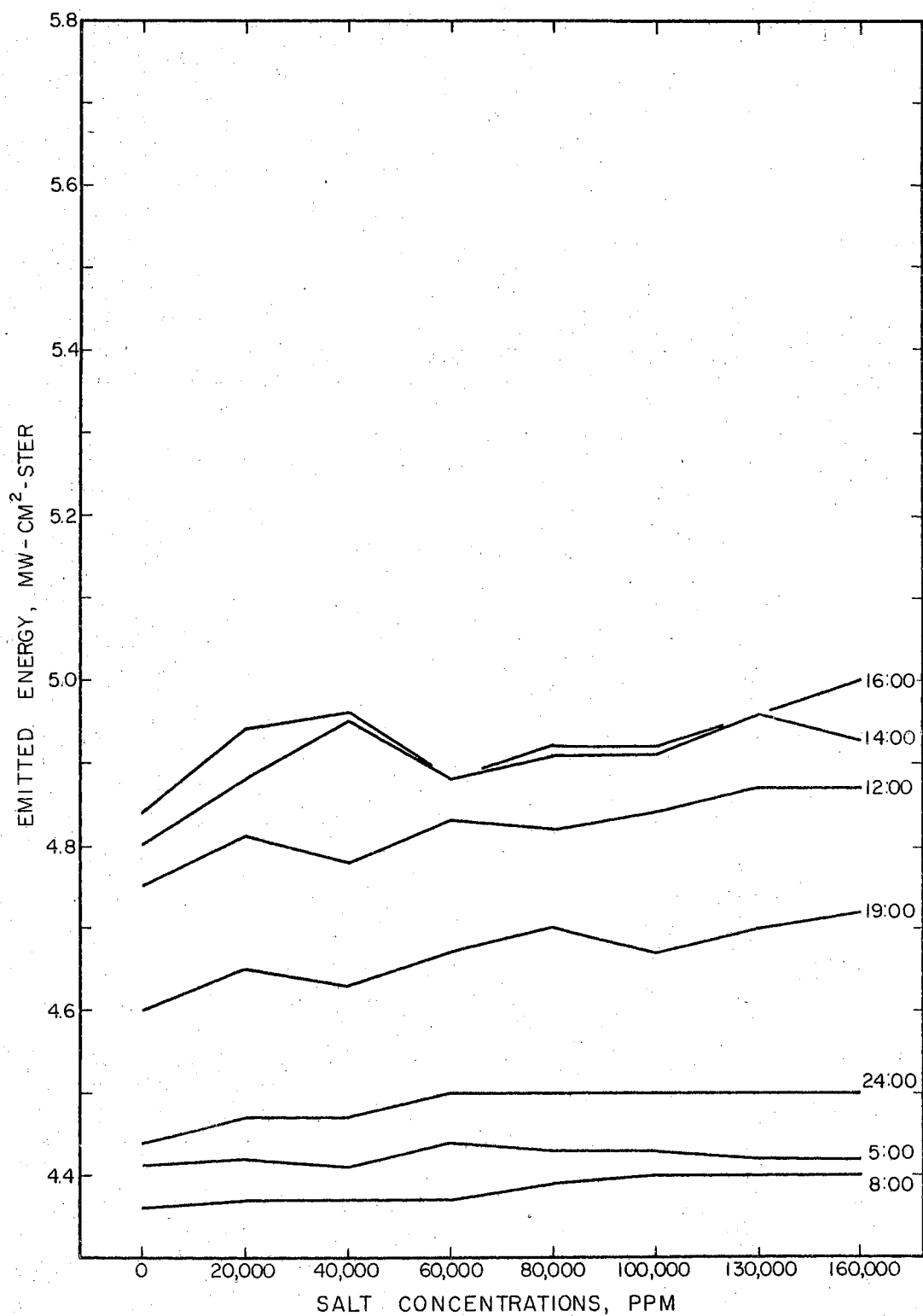


Figure 29. Emitted Energy of Silty Clay with Variance in Moisture and Salinity for September 24

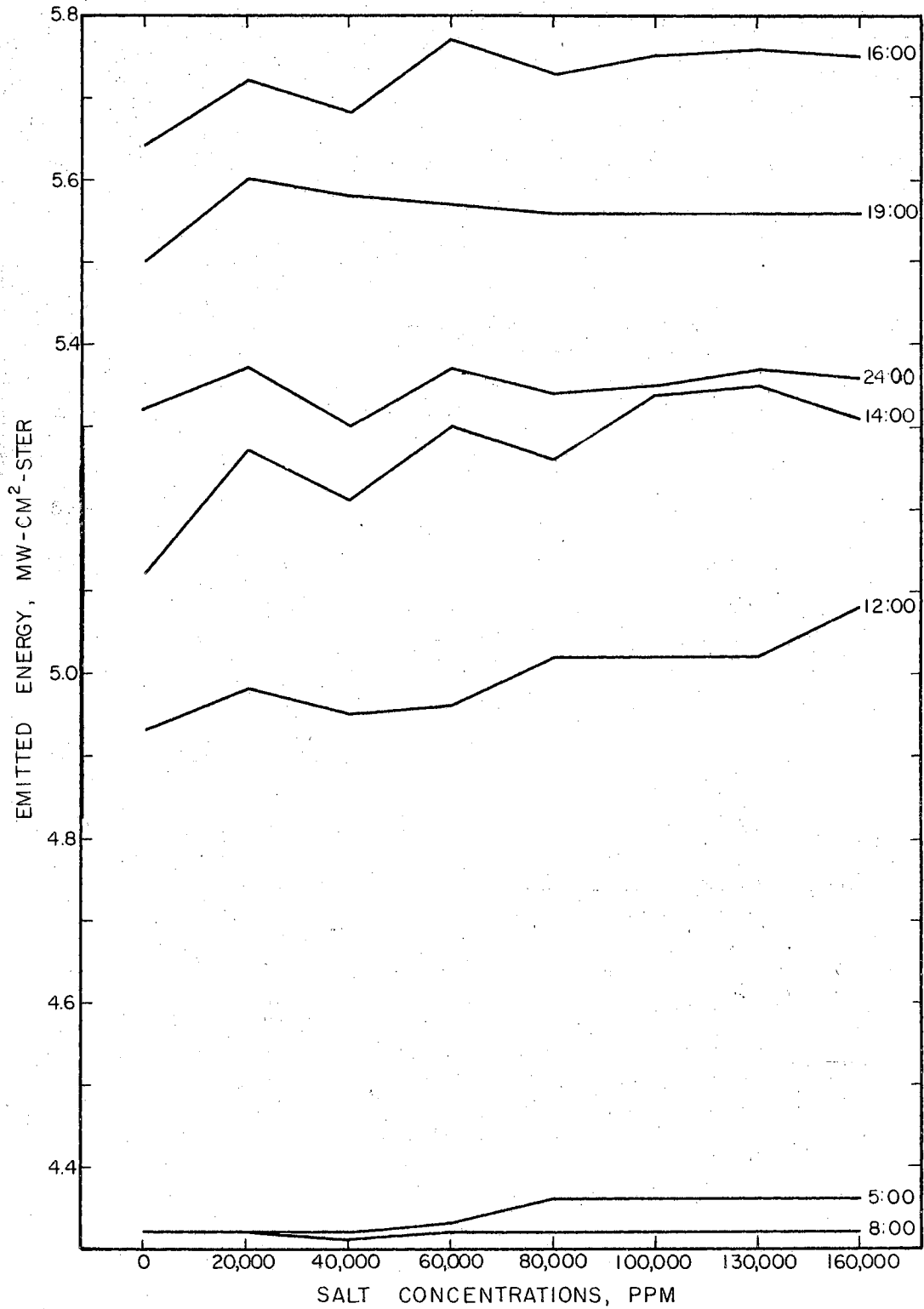


Figure 30. Emitted Energy of Silty Clay with Variance in Moisture and Salinity for September 25

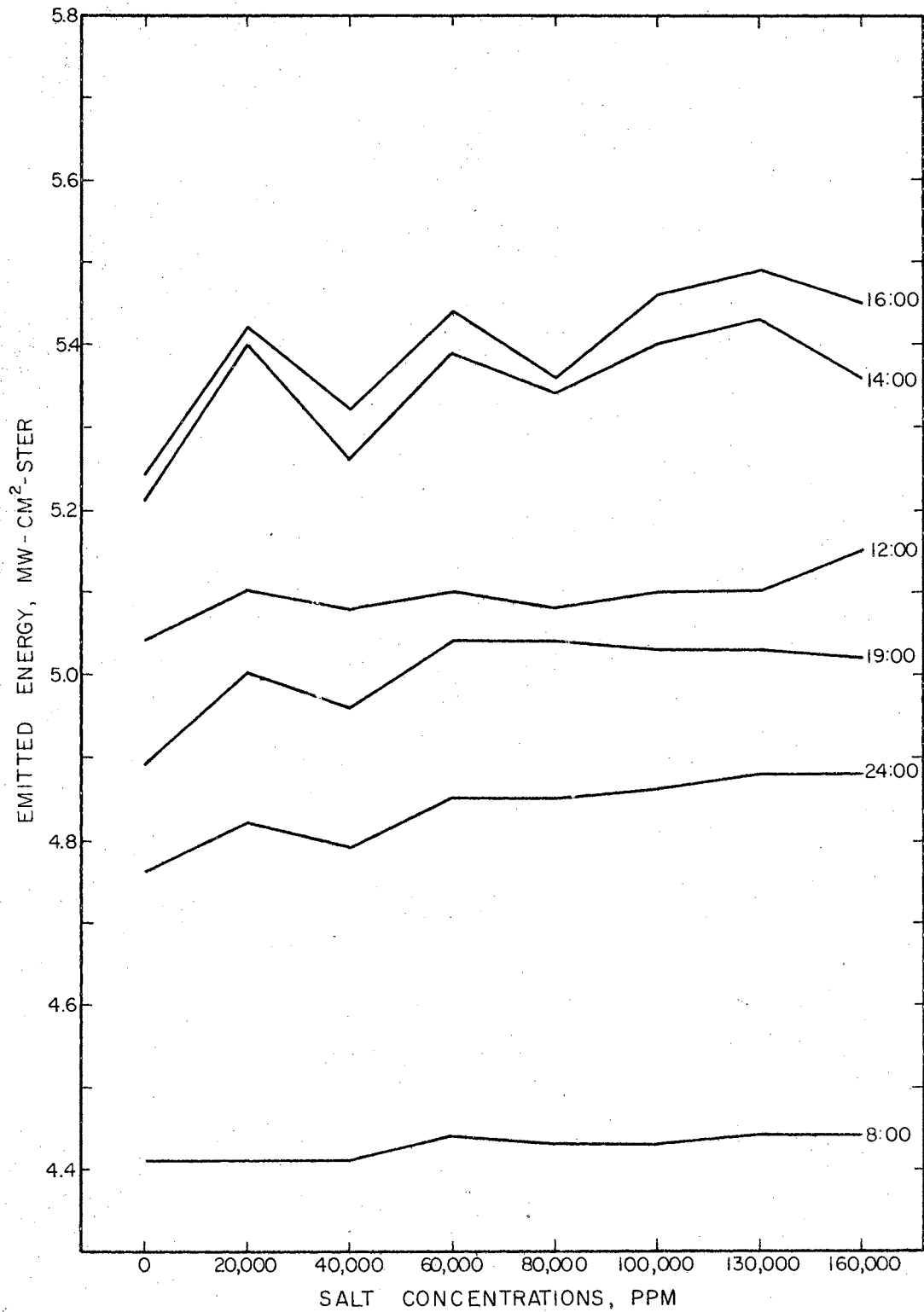


Figure 31. Emitted Energy of Silty Clay with Variance in Moisture and Salinity for September 26

0 ppm moisture content decreasing 9.5% and 5.5% for the 160,000 ppm. As can be seen the silty clay had a great deal more variance in the amount of moisture it contained at its different levels of salinity than that contained by the sand. For this reason, the emitted energy of the silty clay was much more erratic than that given off by the sand.

This trend of emittance occurred throughout the experiment except on Sept. 21. On this particular day the highest emittances were given off by the soils containing the lowest salt concentrations. This occurred for both soil types at 16:00, 19:00 and 24:00 hours. (See Figures 20 and 26.) There is no explanation for this occurrence either theoretically or experimentally.

Emitted Energy of Soils for Changing

Non-Saline Ground Water Table

The emitted energy was taken from the surfaces of both silty clay and sand in different levels of non-saline water. This test was to determine the effect of water table on the emitted energy of the soils. Figures 32 through 36 show the plots of emitted energy taken of the silty clay from May 9 through 13. Figures 37 through 42 show the plots of emitted energy taken of the sand from May 19 through 24. Appendix C denotes the emitted energy and meteorological conditions taken for these plots.

The test results indicated that both silty clay and sand have the same pattern of emittance over a 24-hour period. Their peak emittances occurred between 2:00 and 4:00 in the afternoon, and their lowest emittances occurred about 6:00 a.m., shortly before sunrise. The maximum variance in the emittances of the different levels of the water table

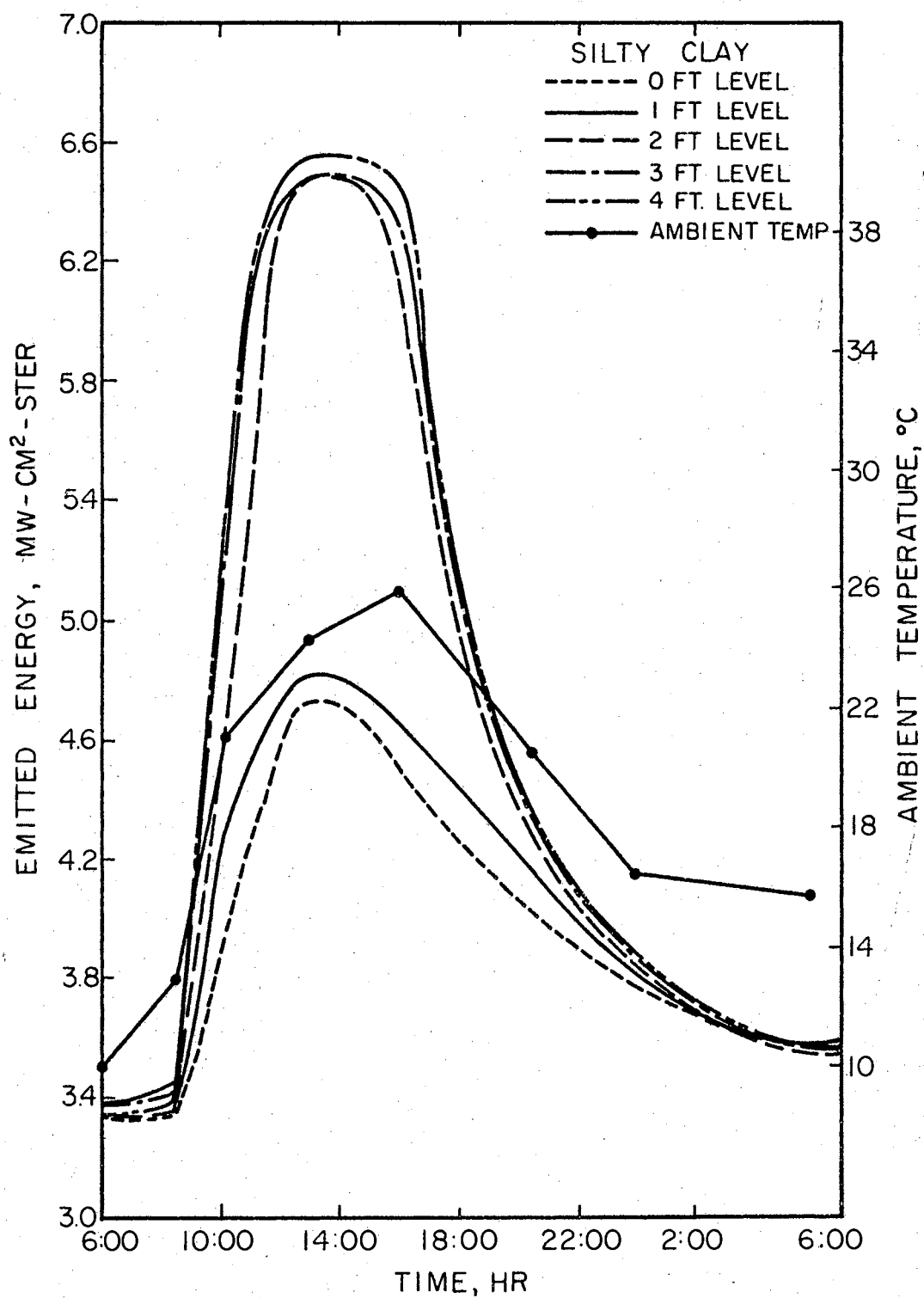


Figure 32. Emitted Energy of Silty Clay for Changing Non-Saline Water Table on May 9

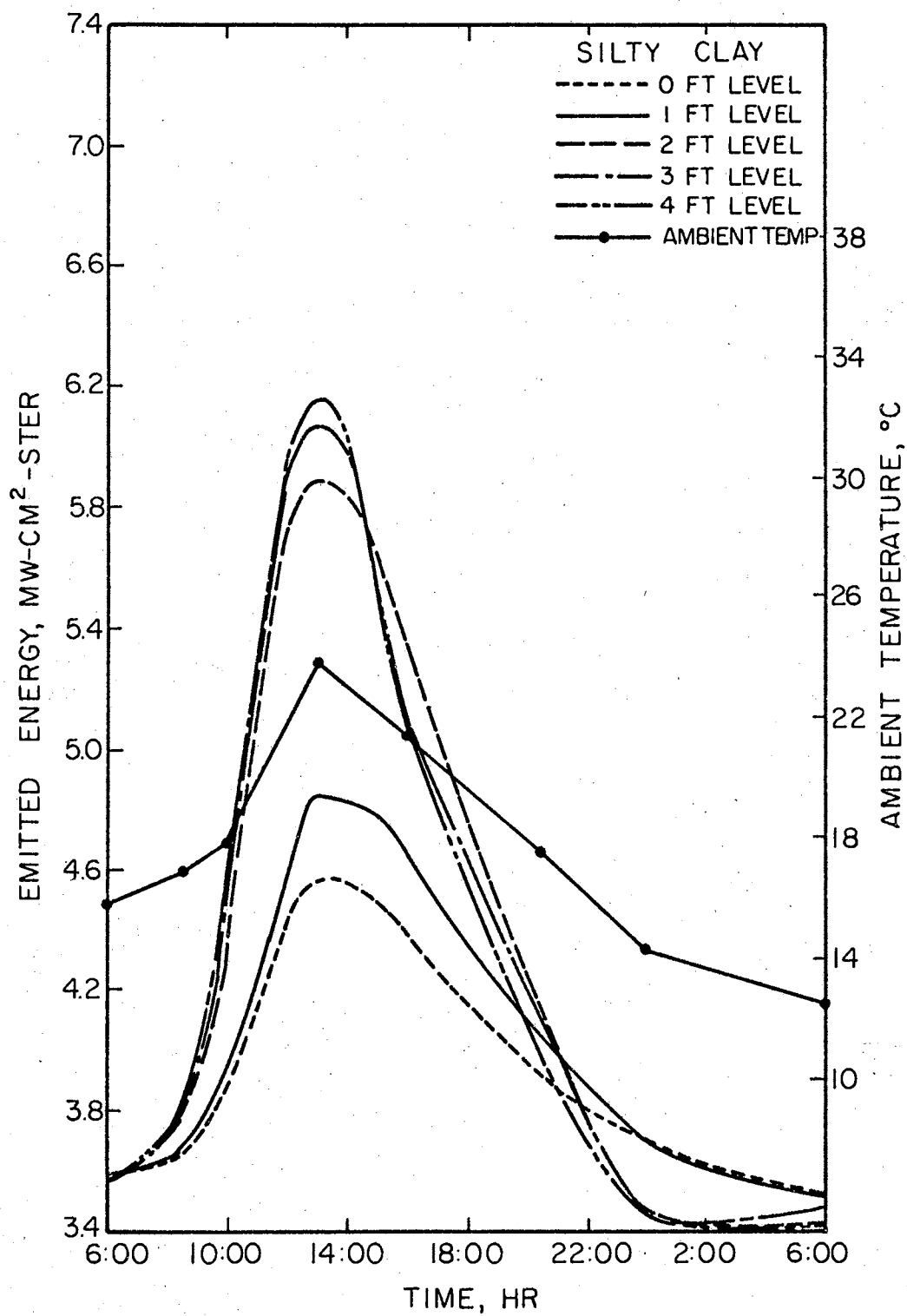


Figure 33. Emitted Energy of Silty Clay for Changing Non-Saline Water Table on May 10

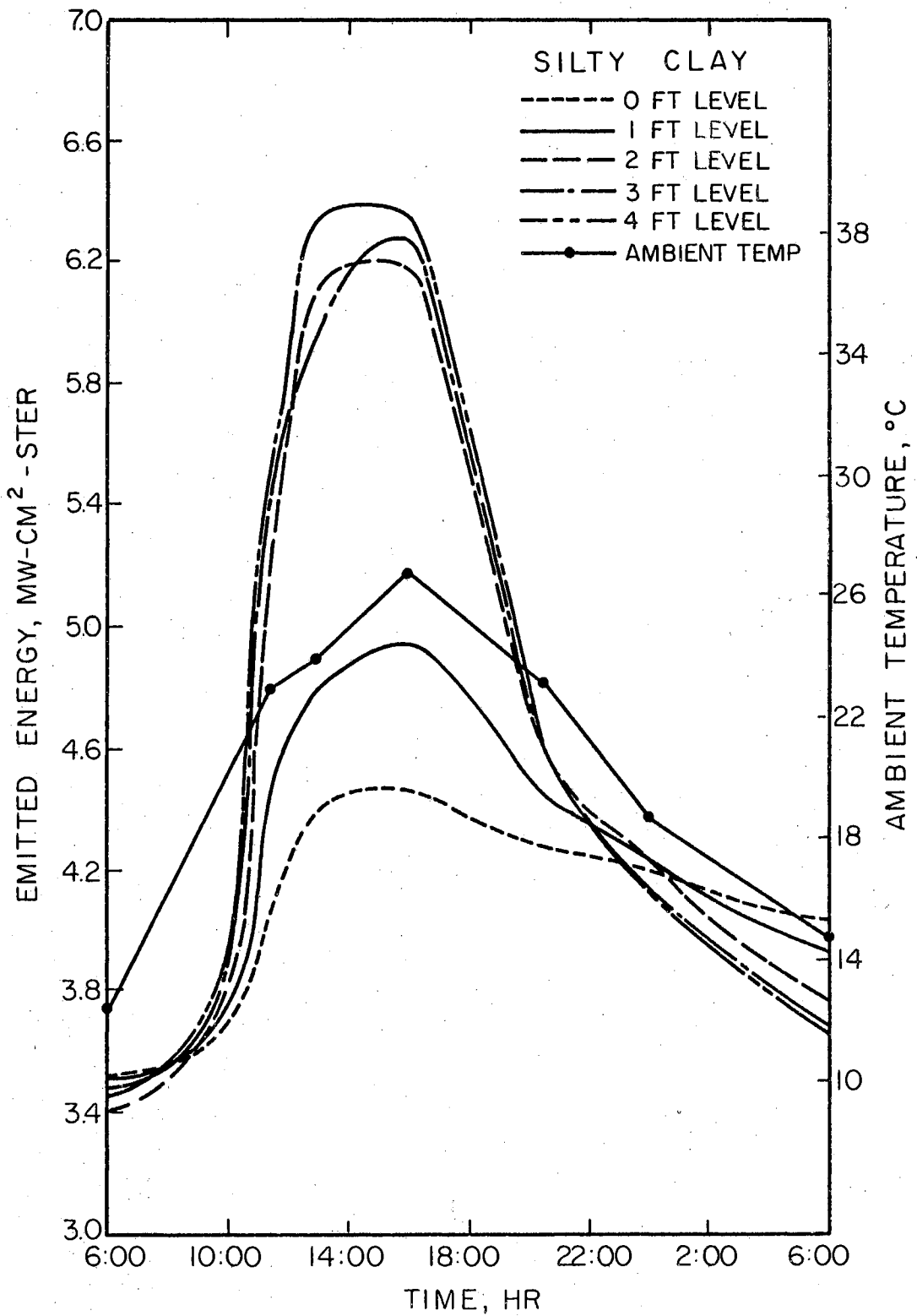


Figure 34. Emitted Energy of Silty Clay for Changing Non-Saline Water Table on May 11

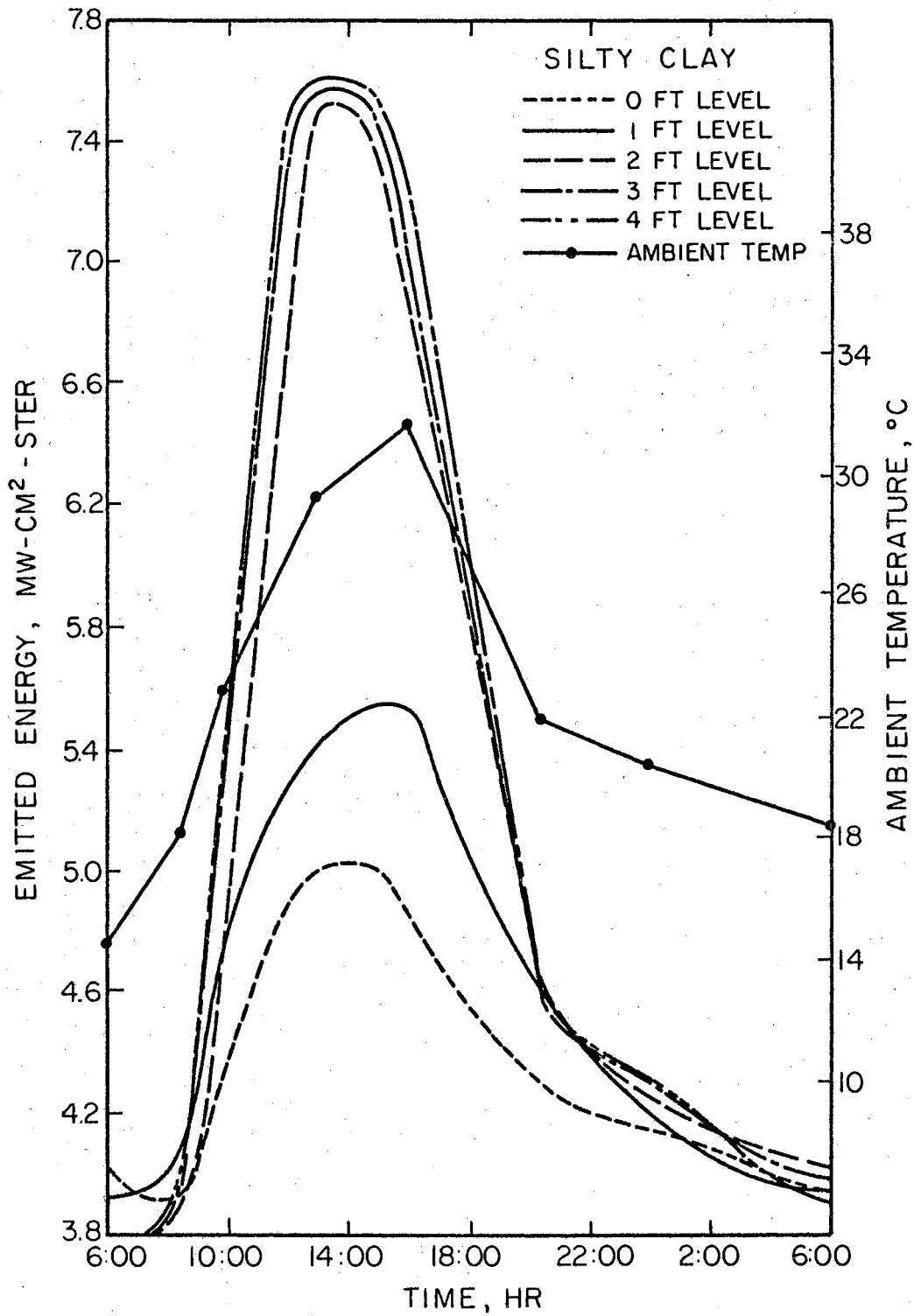


Figure 35. Emitted Energy of Silty Clay for Changing Non-Saline Water Table on May 12

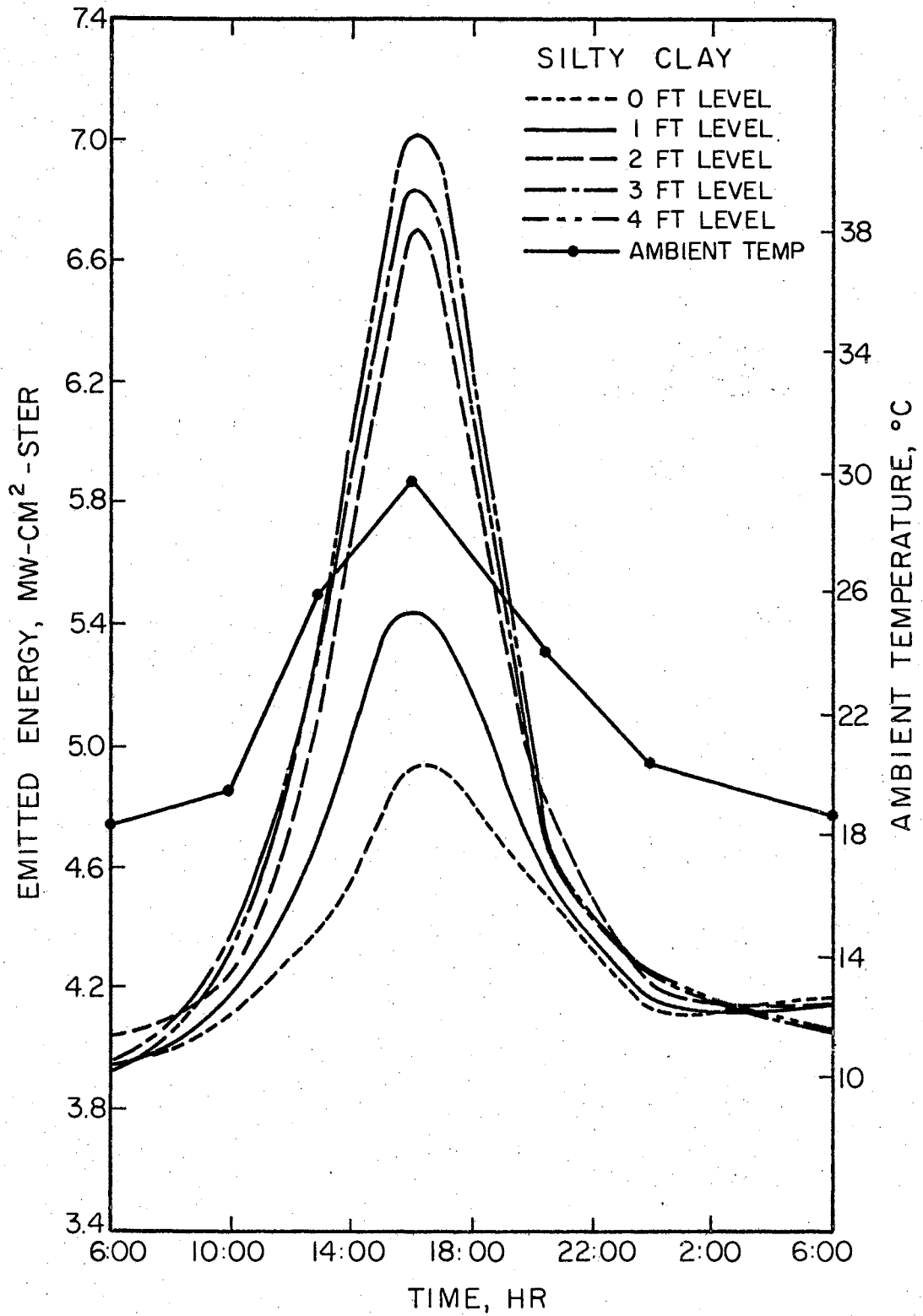


Figure 36. Emitted Energy of Silty Clay for Changing Non-Saline Water Table on May 13

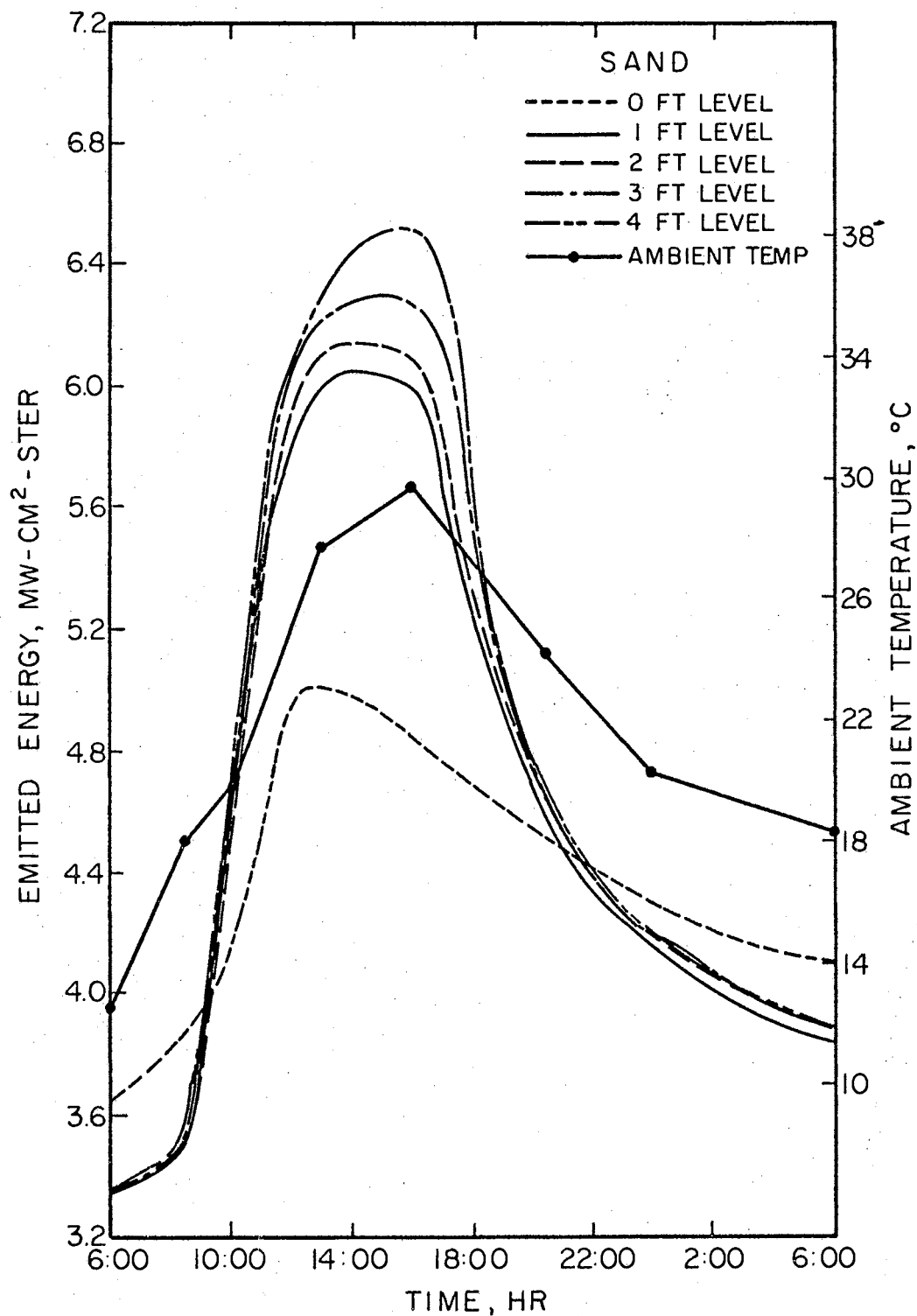


Figure 37. Emitted Energy of Sand for Changing Non-Saline Water Table on May 19

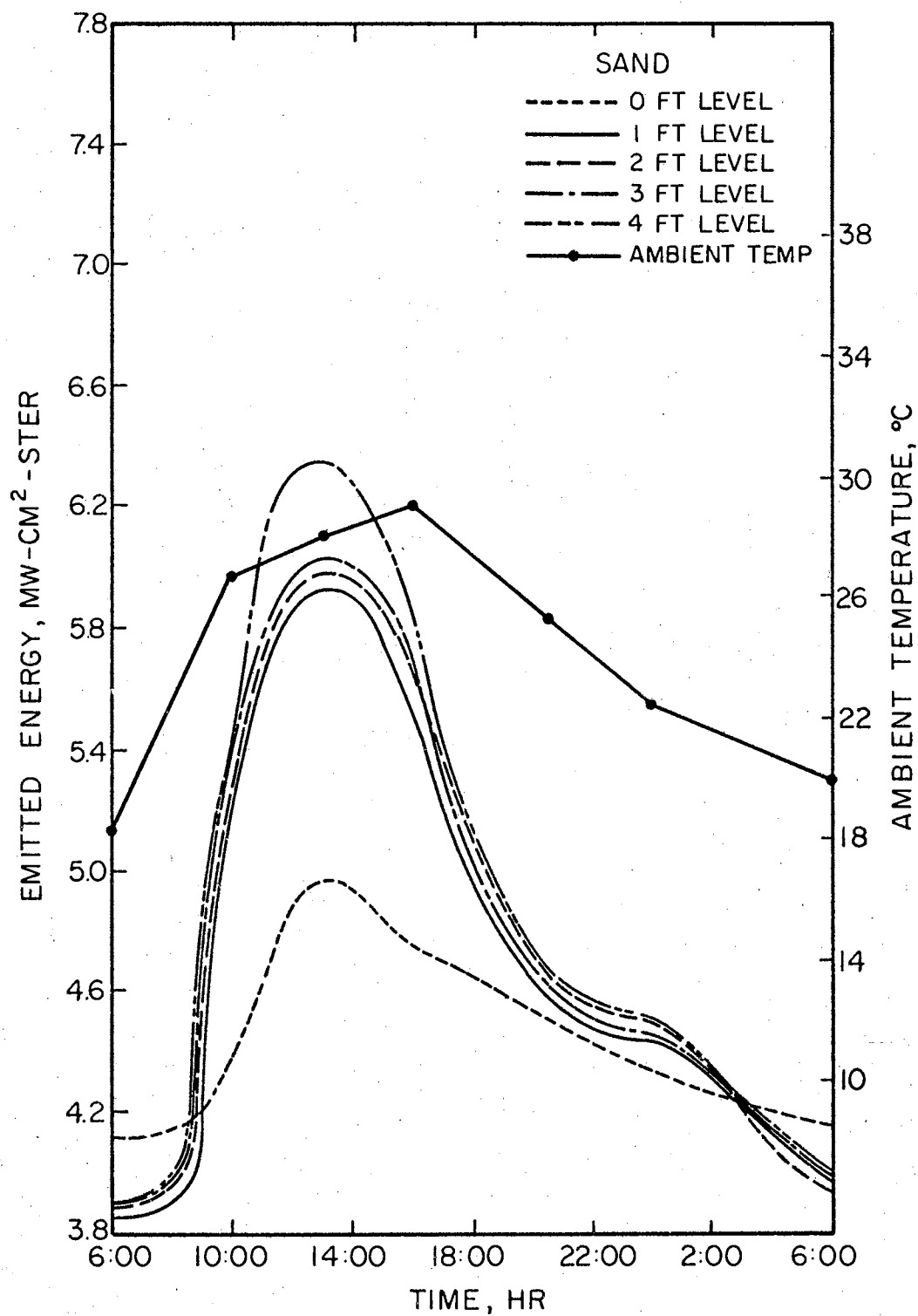


Figure 38. Emitted Energy of Sand for Changing Non-Saline Water Table on May 20

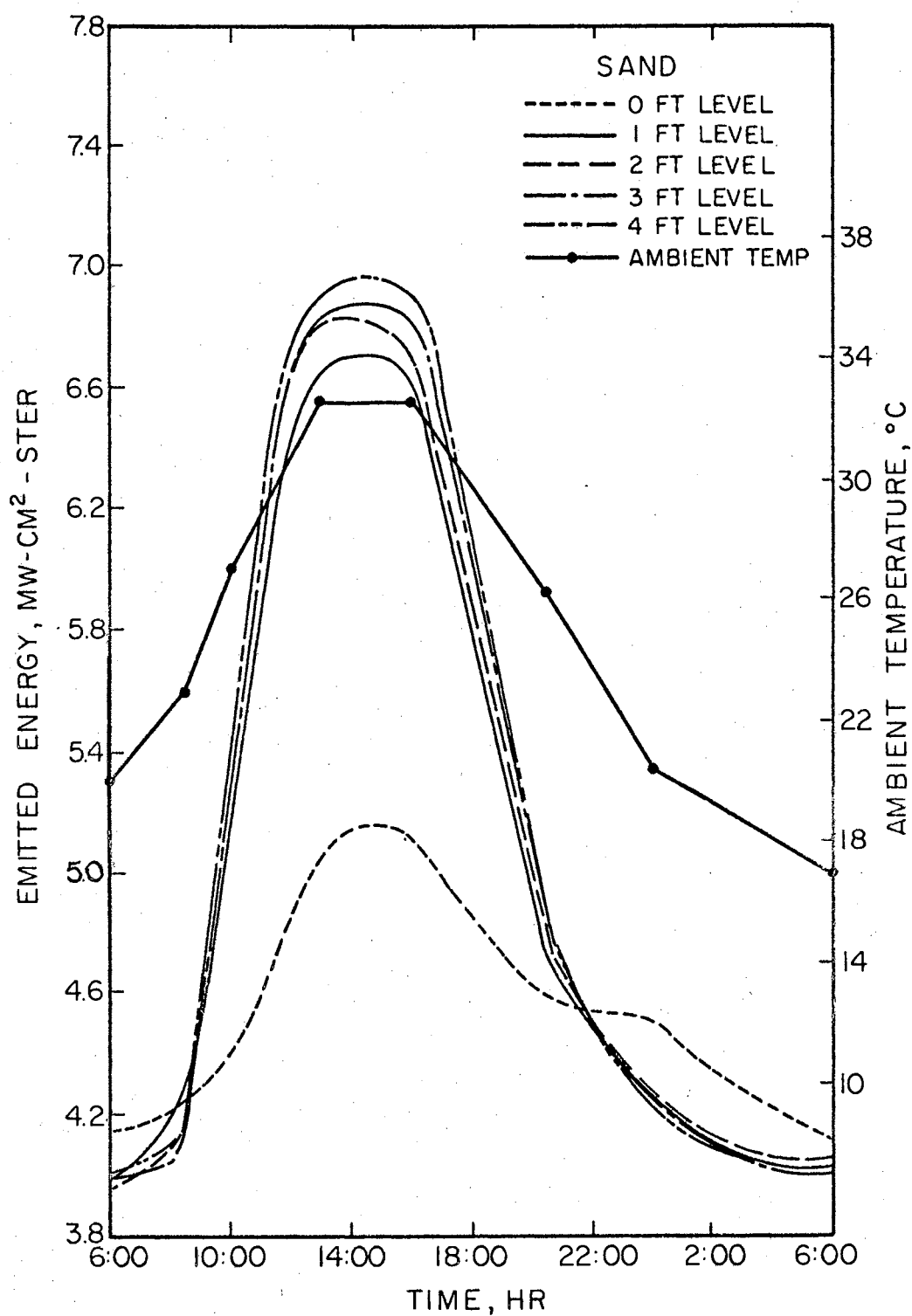


Figure 39. Emitted Energy of Sand for Changing Non-Saline Water Table on May 21

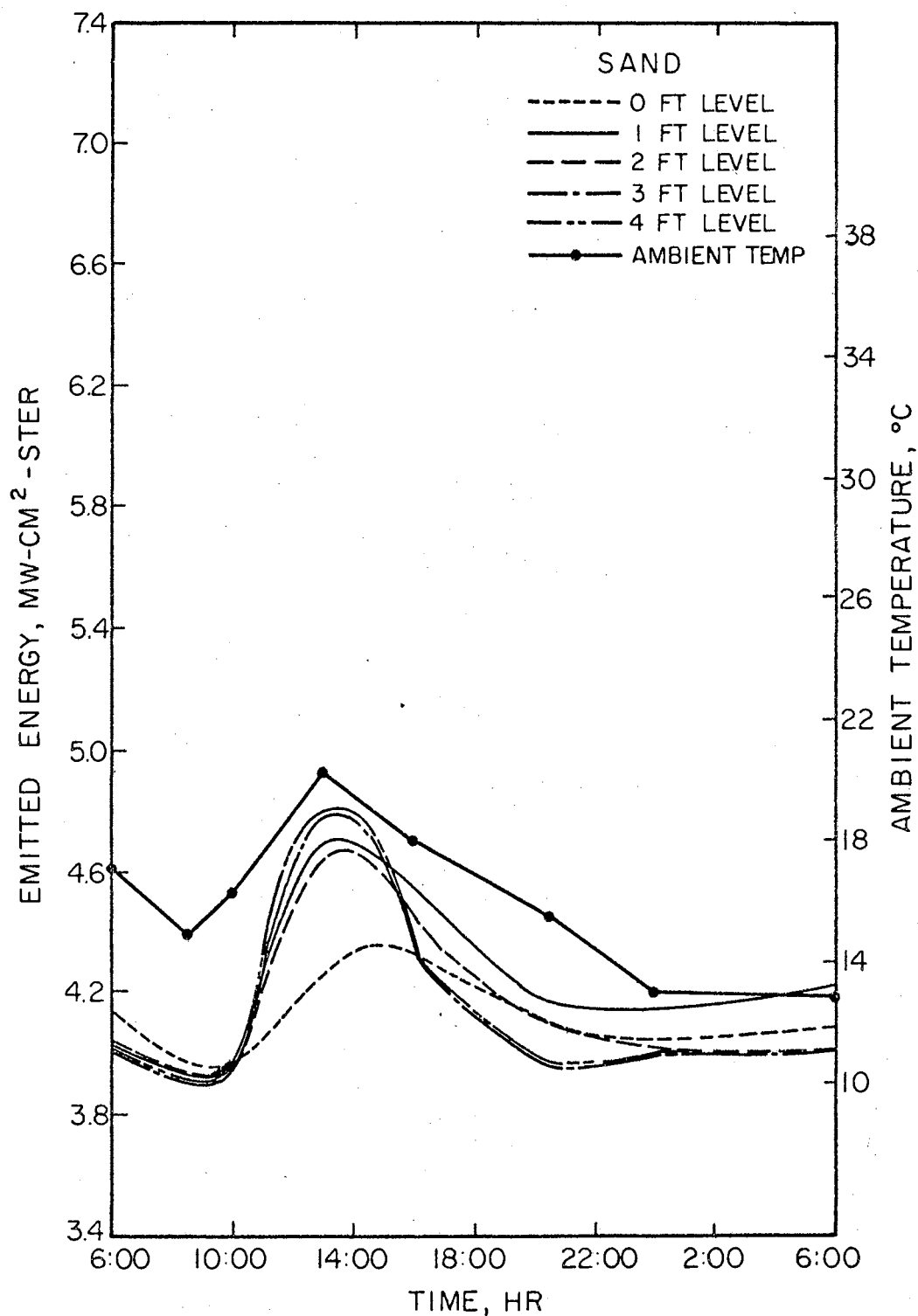


Figure 40. Emitted Energy of Sand for Changing Non-Saline Water Table on May 22

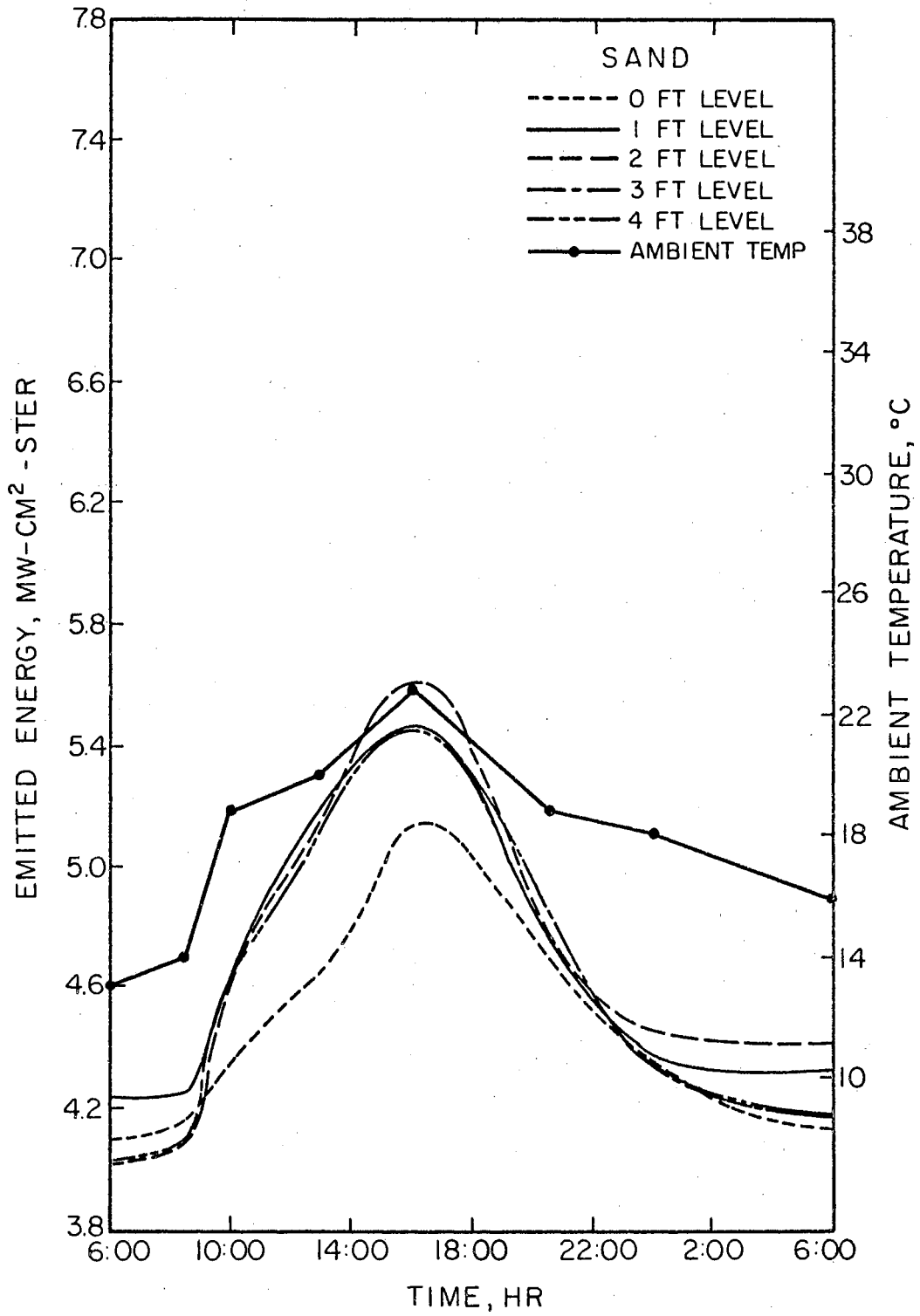


Figure 41. Emitted Energy of Sand for Changing Non-Saline Water Table on May 23

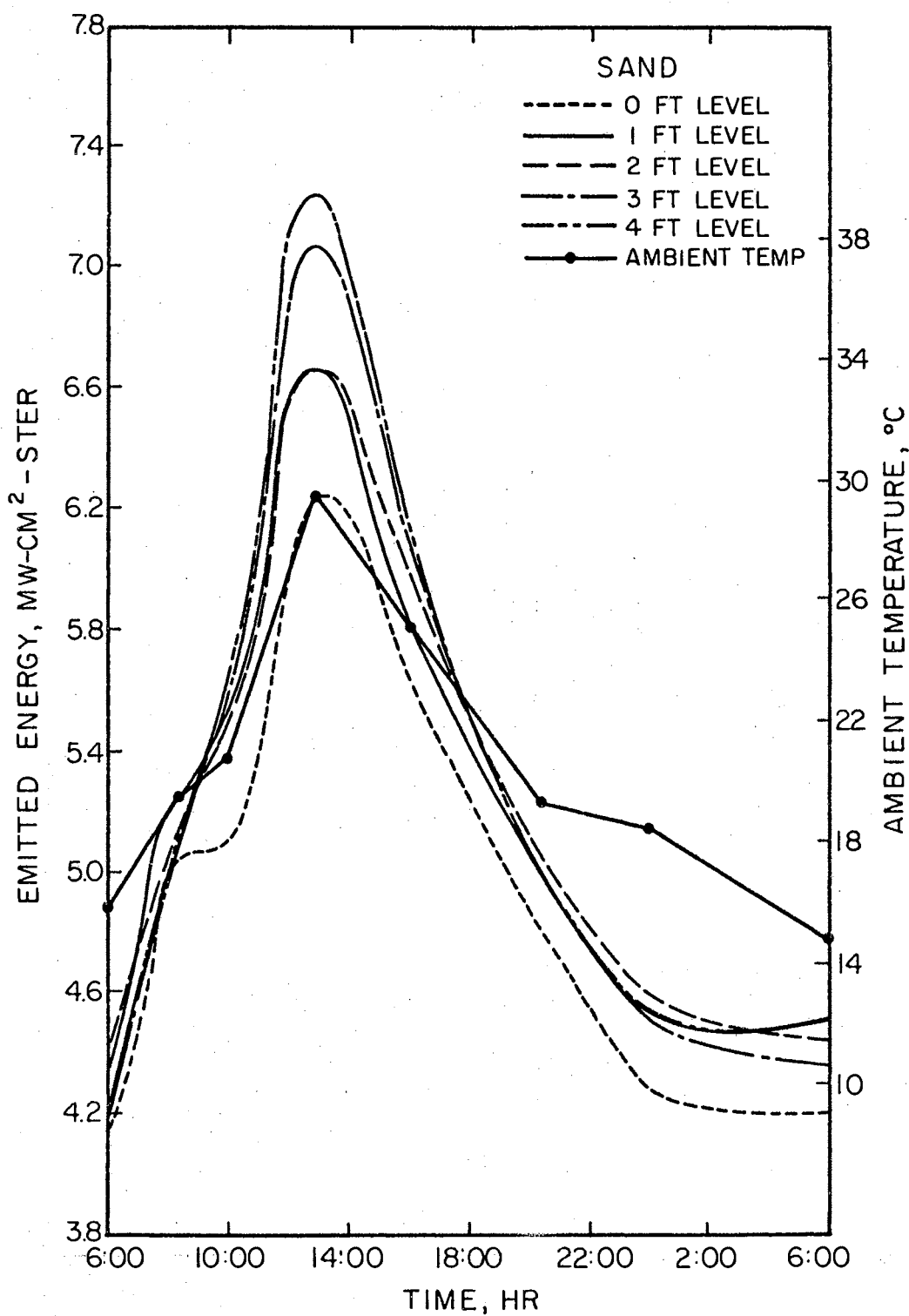


Figure 42. Emitted Energy of Sand for Changing Non-Saline Water Table on May 24

also took place between 2:00 and 4:00 p.m. for both soil types, with the 4 ft. level having the highest emittance and 0 ft. level (water level at surface of soil) having the least emittance of radiant energy. There was an inversion in the amounts of emittances given off by the soils between 12:00 midnight and 6:00 a.m., with the highest energy being emitted by the 0 ft. level (water level at surface of soil) and the least energy being emitted by the 4 ft. level of both soil types.

These variations in emittances between the different levels of the water table are due to the diverse thermal characteristics of air and water. Water takes the place of air in soil voids as the soil becomes moist. Since water has a specific heat of 1.0 and air has only .000306, water has great influence on the ability of the soil to be heated and cooled (2) (3). Thus, the emitted energy varies according to the moisture content of the soil. For this reason, the maximum variance in emittance between the different levels of the water table occurs during the hottest portion of the day, since more solar energy is needed to raise the temperature of water than is needed to raise that of the soil. This fact further explains the inversion of emittances which occurred at night. The more moisture contained in the soil results in a higher heat capacity due to the higher specific heat of water. Since heat capacity is the ability of a material to hold heat, the wetter the soil is, the longer it will hold heat. Furthermore, the drier soils cool faster than the wet soils, therefore the emitted energy is less (3). This resulted in the inversion of emittances at night.

Emitted Energy of Soils for Changing

Saline Ground Water Table

Test results for the changing saline water table are given in Appendix D for silty clay and sand. These results are also shown in Figures 43 through 53 for sand and Figures 54 through 63 for silty clay. The emitted energy for silty clay was taken from June 13 through June 23. The emitted energy for the sand was taken from May 30 through June 9. Test results for both the silty clay and sand show that there was a slight difference in emitted energy from the soils having variance in concentrations of saline water table.

The sand had the greater variance of emitted energy from the changing levels of its saline water table due to its higher thermal conductivity. The highest emittance for the sand occurred between 2:00 and 4:00 in the afternoon with the 0 ppm solution giving the highest emitted energy. This took place at the 4 ft., 3 ft., 2 ft. and 1 ft. levels of water table, as shown in Figures 43 through 50. The reason for this is that the higher the salt concentrations added to the soils, the higher the specific heat of these soils becomes, therefore the less its temperature can be raised with the same amount of solar energy. Since these water table levels were below the surface, evaporation was lessened and did not affect the surface temperature.

Figures 52 and 53 are plots of the emitted energy for sand at the 0 ft. level. As can be seen, there was an inversion of the emittance at this level, compared to the 4 ft., 3 ft., 2 ft., and 1 ft. levels, with the highest emitted energy being given off by the soils containing the highest salt concentrations. These same results were found in the

soils tested in the panels as well, and are due to the same reason. At the 0 ft. level the surface was wet and evaporation of the moisture had a cooling effect on the soil. The soils containing the least amount of salt concentrations had the greatest amount of evaporation, therefore, emitted less energy. This inversion of emitted energy also occurred on the second day of the test at the 1 ft. water table level, as shown in Figure 51. This occurrence was also due to evaporation of water from the soil. A possible reason this did not occur on the first day the water table was at the 1 ft. level could be the difference in temperature on the two days. The temperature on June 6 (Figure 50) was not as high as that on June 7 (Figure 51), therefore, on June 6 enough evaporation may not have taken place to cause an inversion of emittance.

The 4 ft. and 3 ft. water table levels for the silty clay had the same trend of emittance as that of the sand at the 4 ft., 3 ft., 2 ft. and 1 ft. levels, with the soil containing 0 ppm giving off the highest emittance. (Figures 53 through 56) This is also due to the difference in specific heat for the varied salt concentrations. The readings for the silty clay were more erratic than those for the sand and there was not as great a difference between the varied salt concentrations as in the sand. An inversion of the emitted energy by the silty clay occurred at the 2 ft., 1 ft. and 0 ft. water table levels. Therefore it could be seen that the evaporation of the moisture in the soil at these water levels was controlling the temperature instead of the difference in the specific heat, as at the 4 ft. and 3 ft. water levels. The reason for this inversion of emittance for silty clay at a lower water table than that of the sand is the difference in capillary rise in these soils. The silty clay was moist within six inches of the surface when the water

table was at the 2 ft. level, while the water table for the sand had to be at the 1 ft. level to moisten the soil within six inches of the surface. Therefore, the amount of emitted energy given off by both soil types is controlled by the difference in specific heat for the different salt concentrations until the water table is raised to a point where the capillary rise is high enough for evaporation of moisture from the soil to cause a cooling effect on the soil. At this point an inversion of emittance occurs. In this test, inversions of emittances occurred at the 2 ft. level and 1 ft. level for silty clay and sand, respectively.

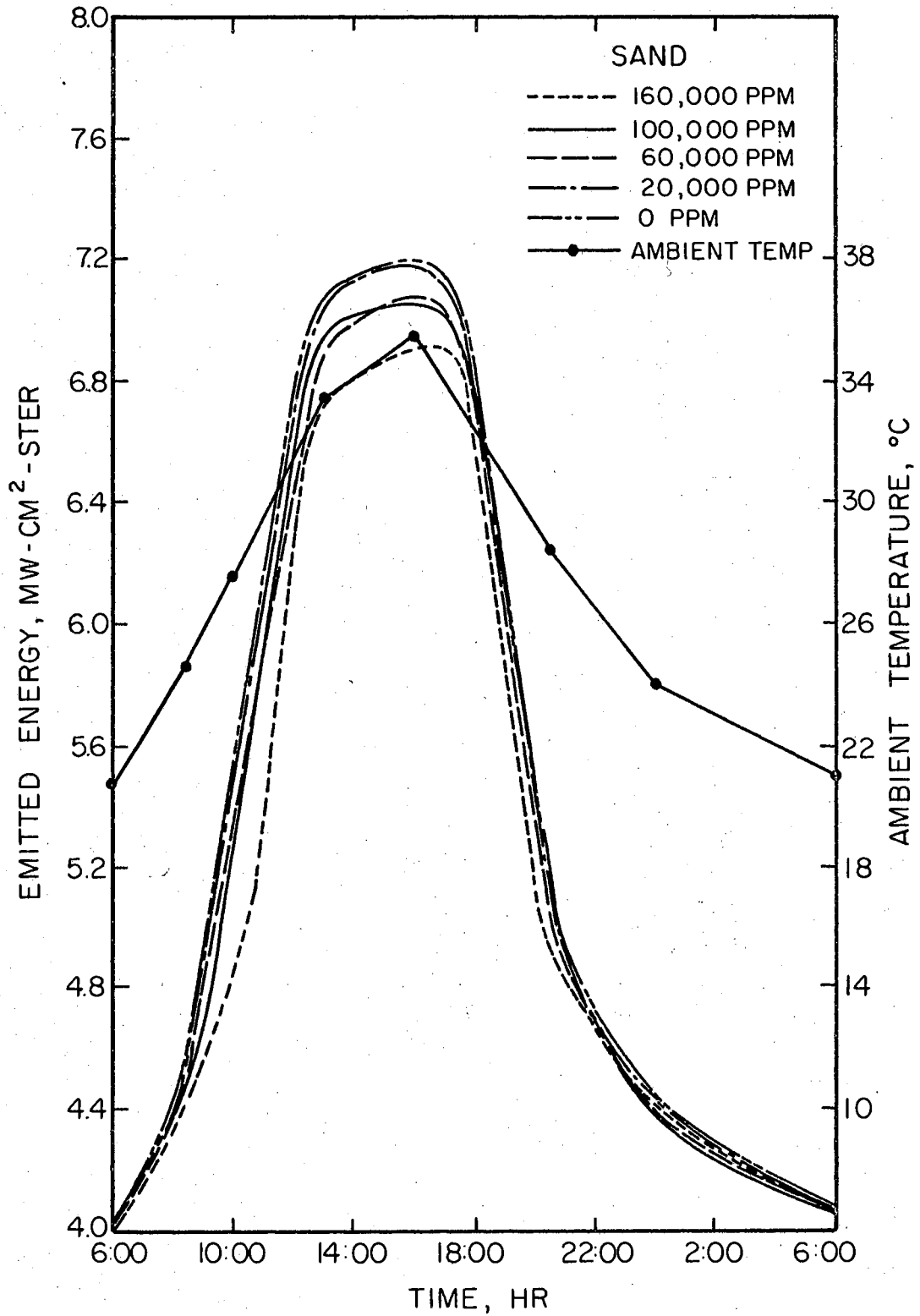


Figure 43. Emitted Energy for May 30 for Saline Water Table at 4 ft. in Sand

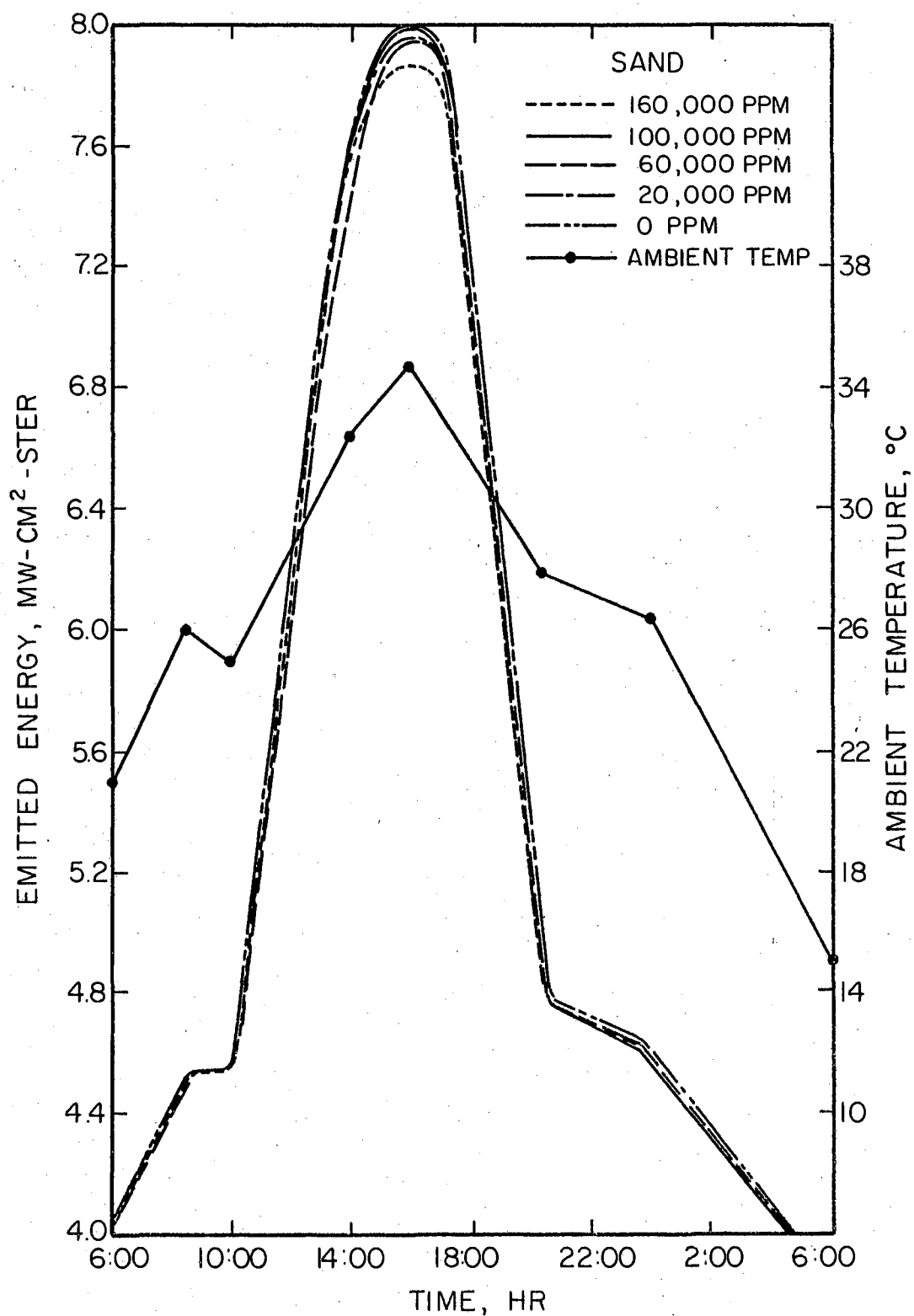


Figure 44. Emitted Energy for May 31 for Saline Water Table at 4 ft. in Sand

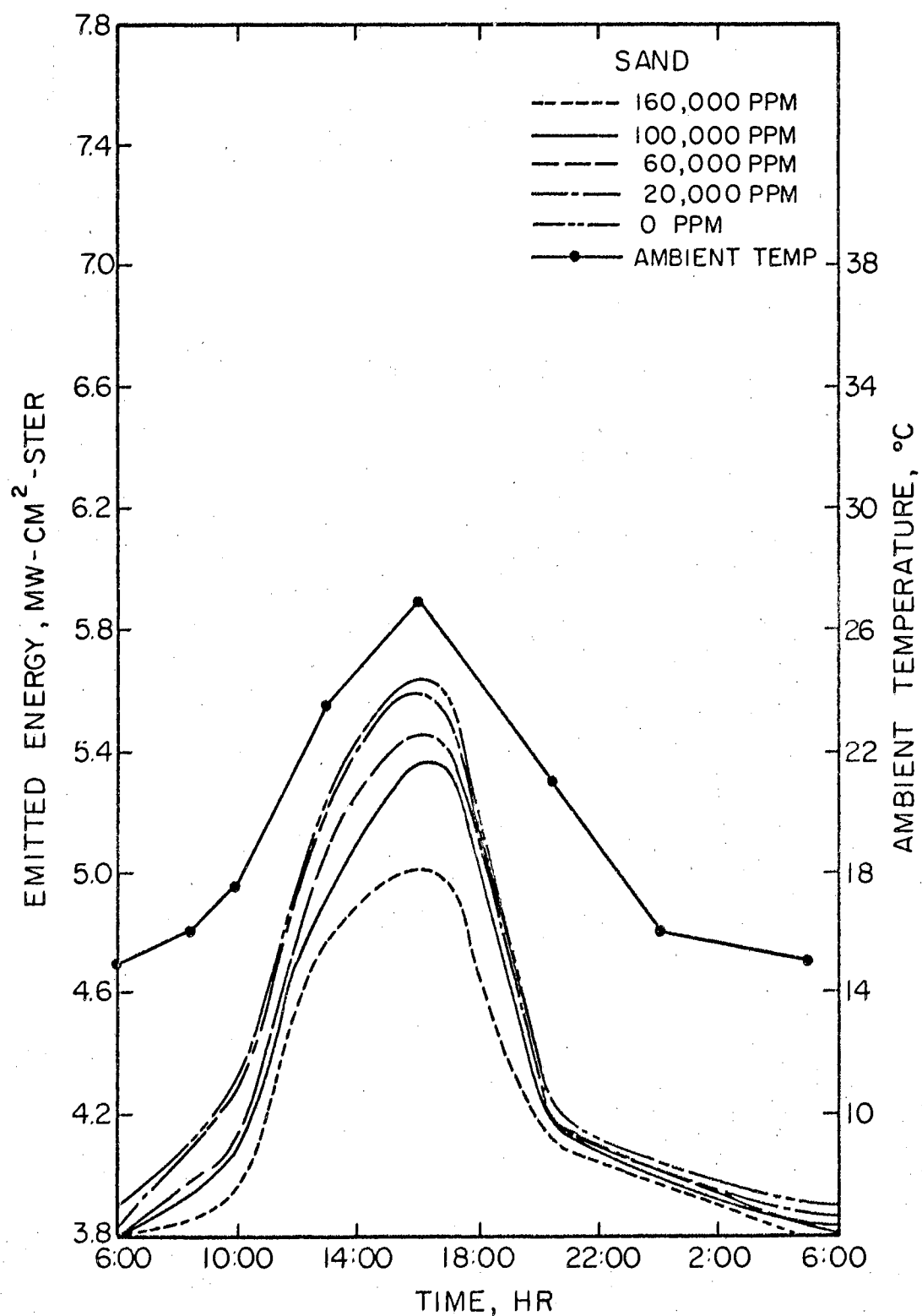


Figure 45. Emitted Energy for June 1 for Saline Water Table at 3 ft. in Sand

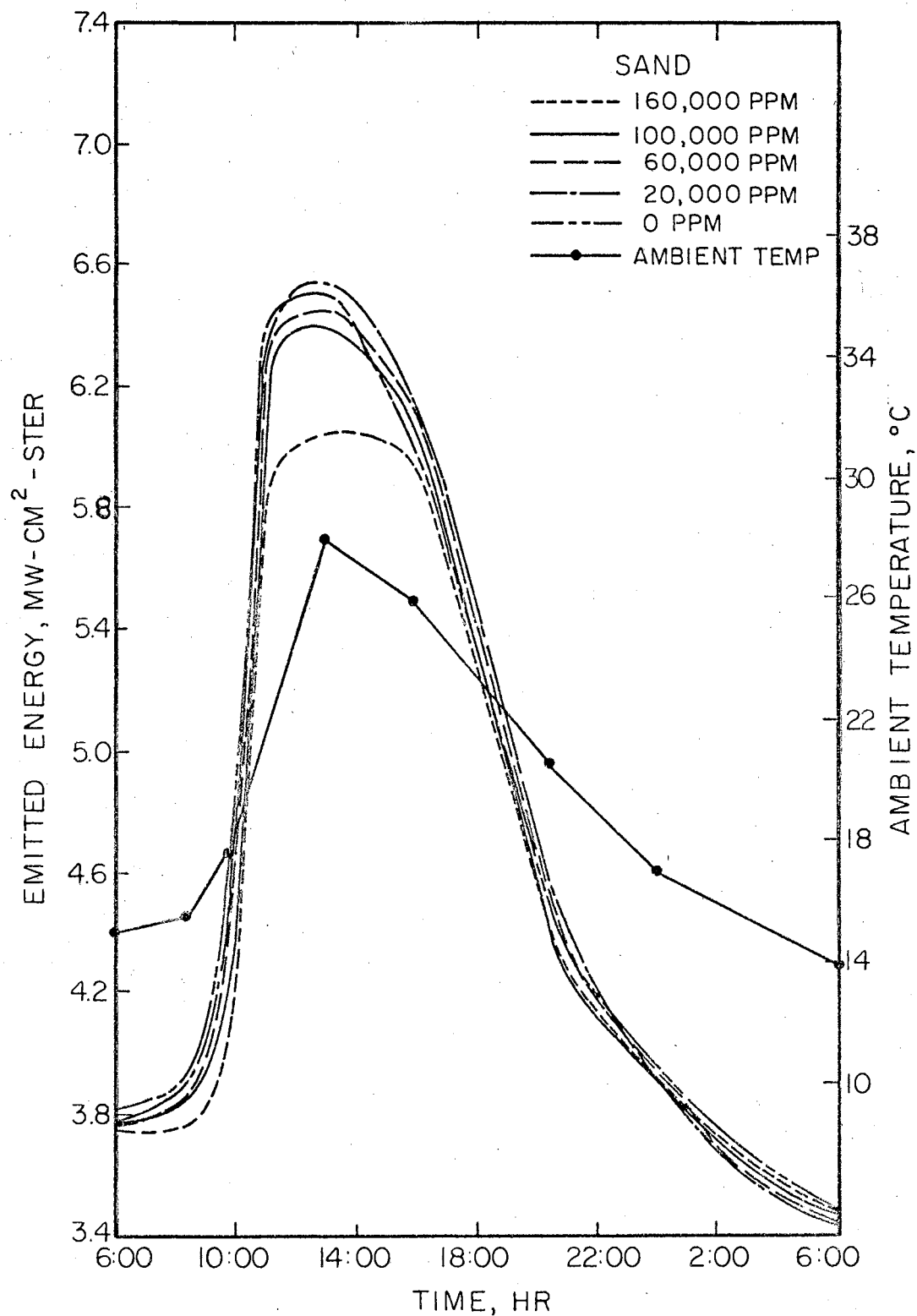


Figure 46. Emitted Energy for June 2 for Saline Water Table at 3 ft. in Sand

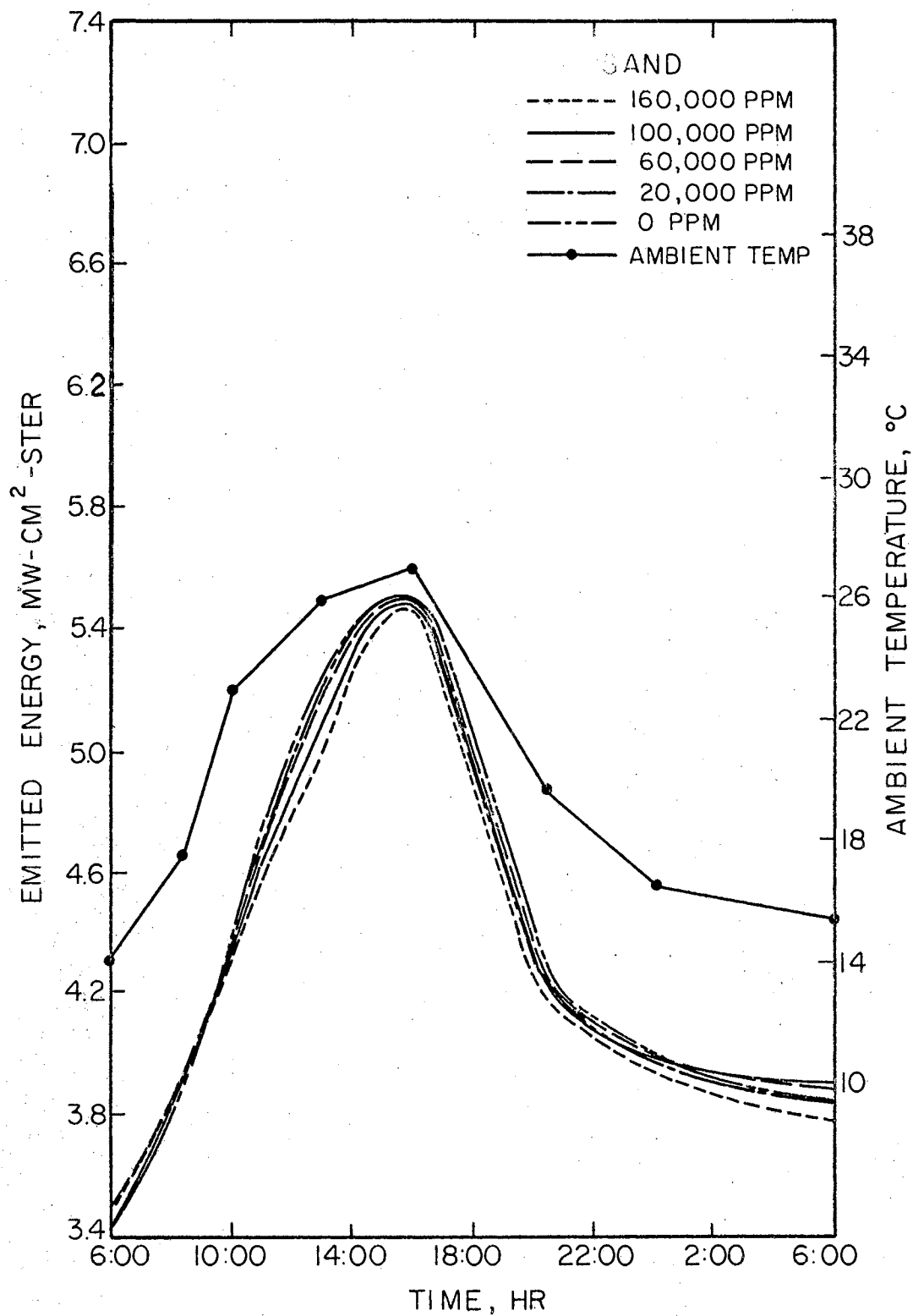


Figure 47. Emitted Energy for June 3 for Saline Water Table at 2 ft. in Sand

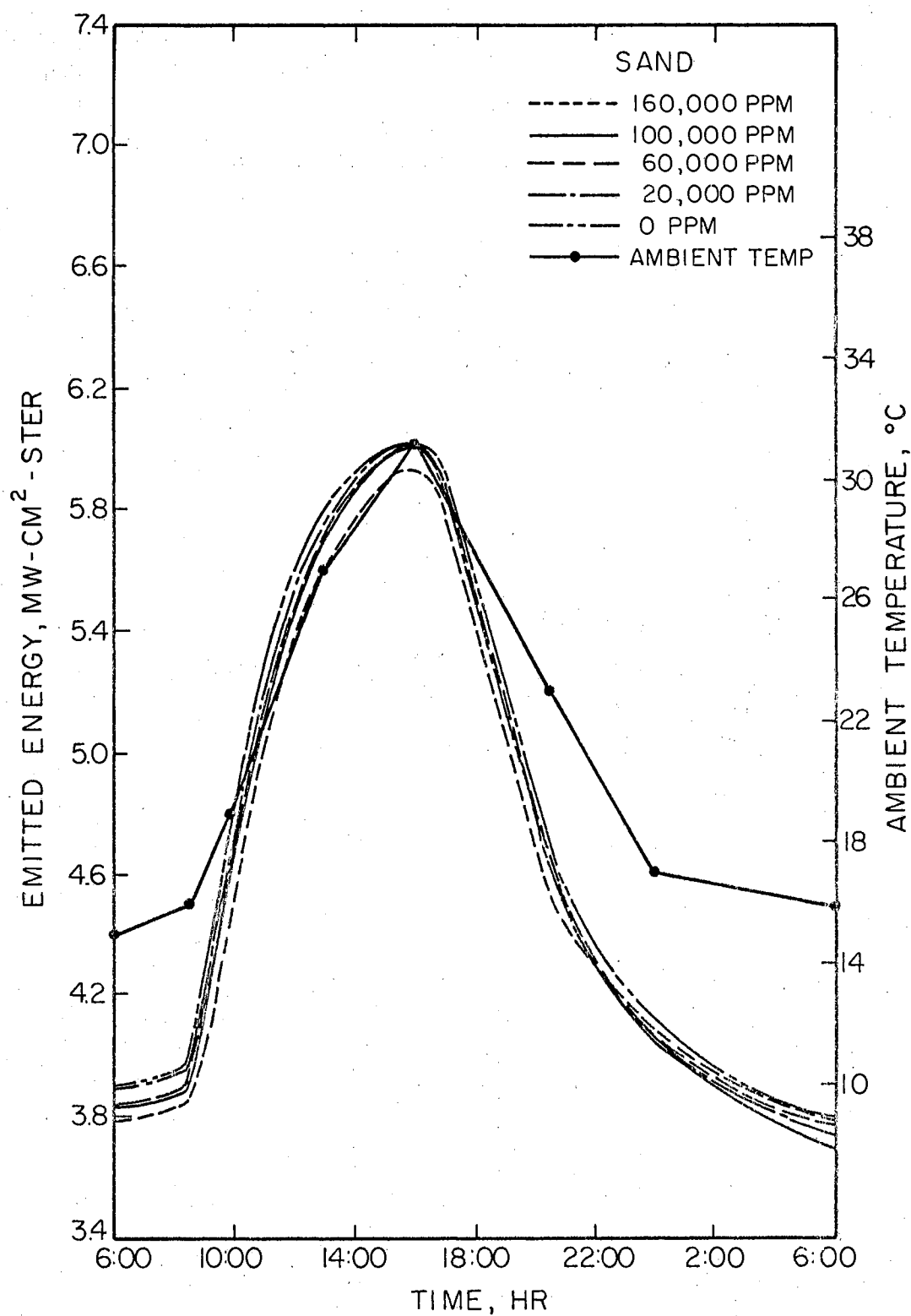


Figure 48. Emitted Energy for June 4 for Saline Water Table at 2 ft. in Sand

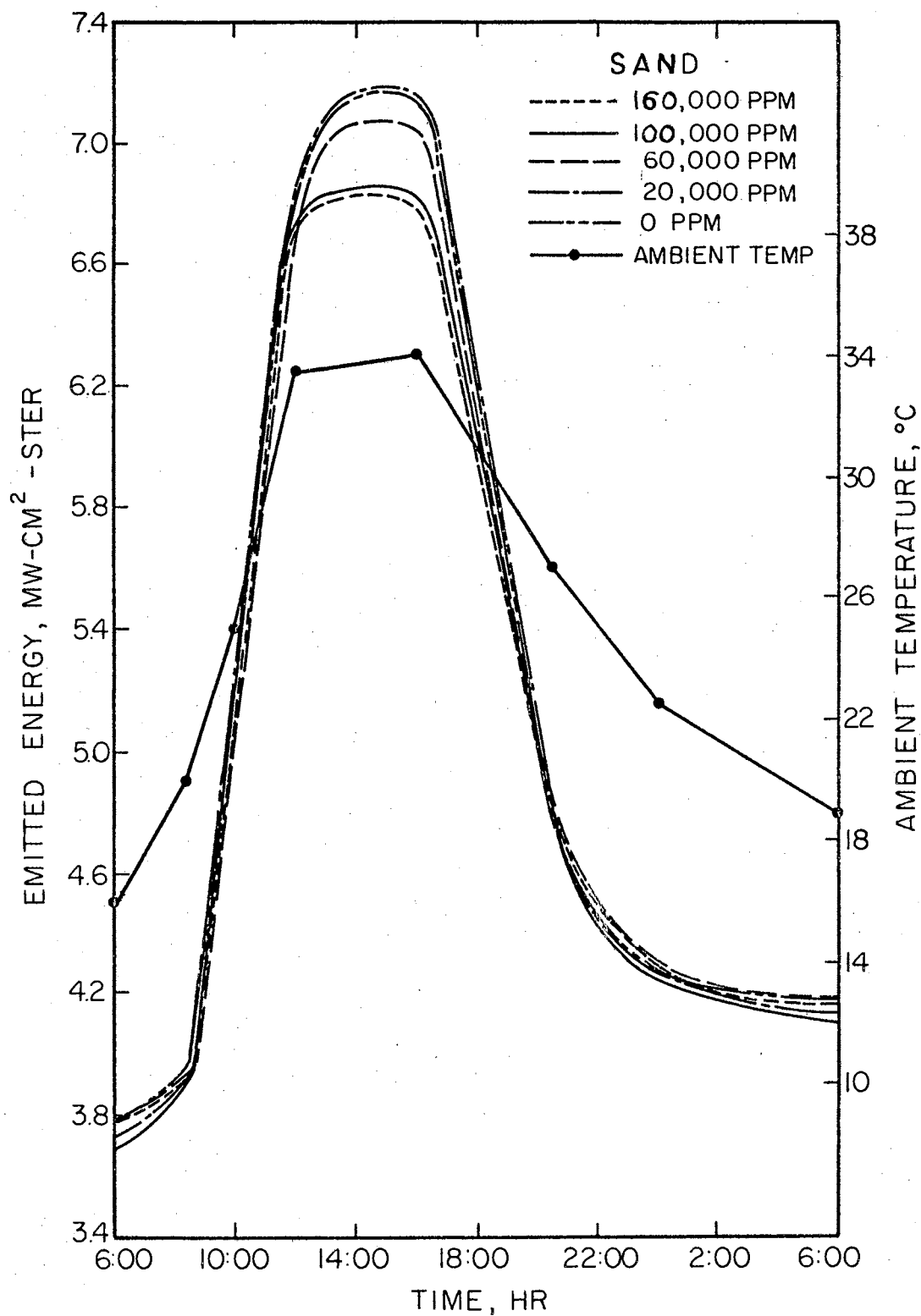


Figure 49. Emitted Energy for June 5 for Saline Water Table at 2 ft. in Sand

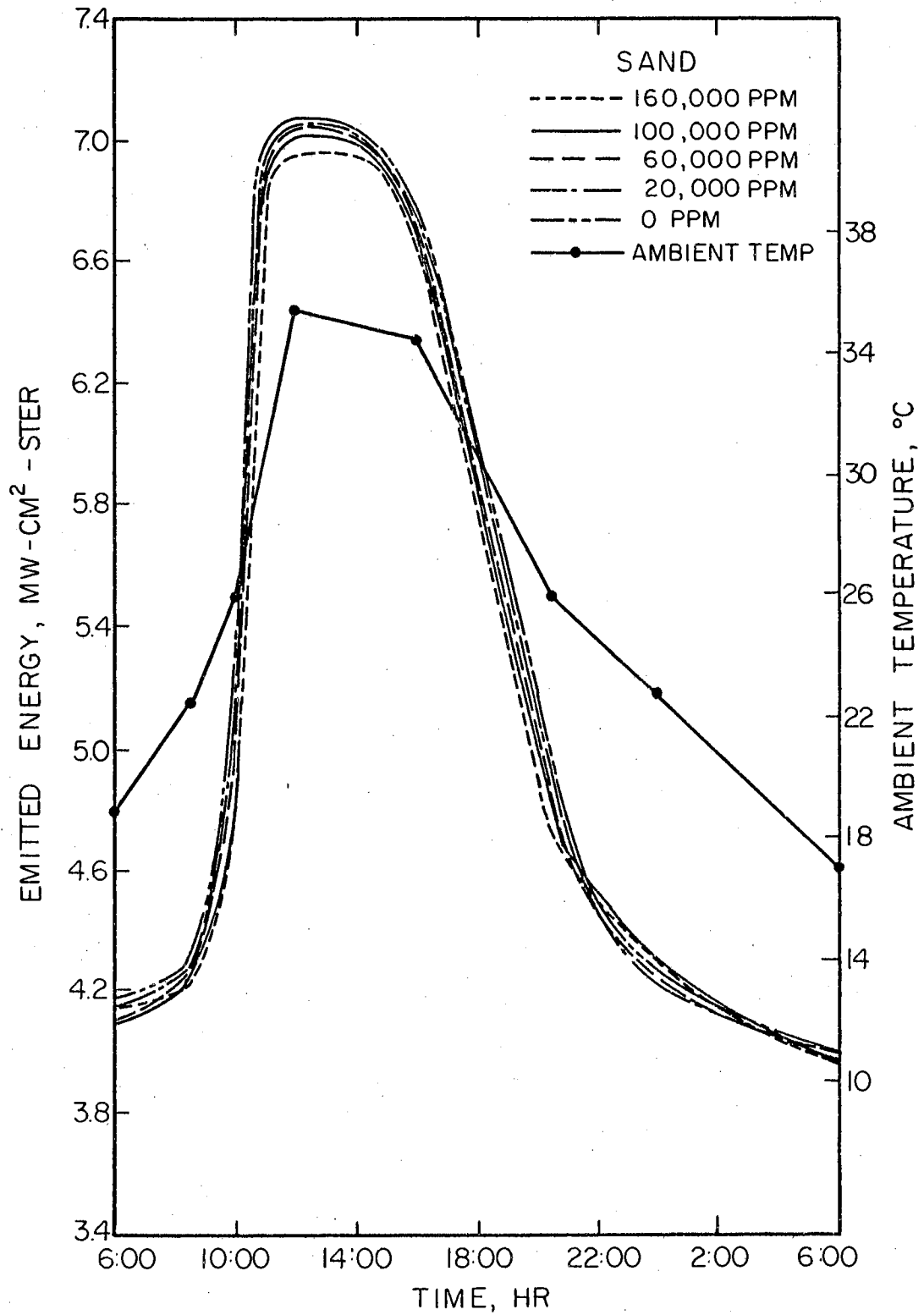


Figure 50. Emitted Energy for June 6 for Saline Water Table at 1 ft. in Sand

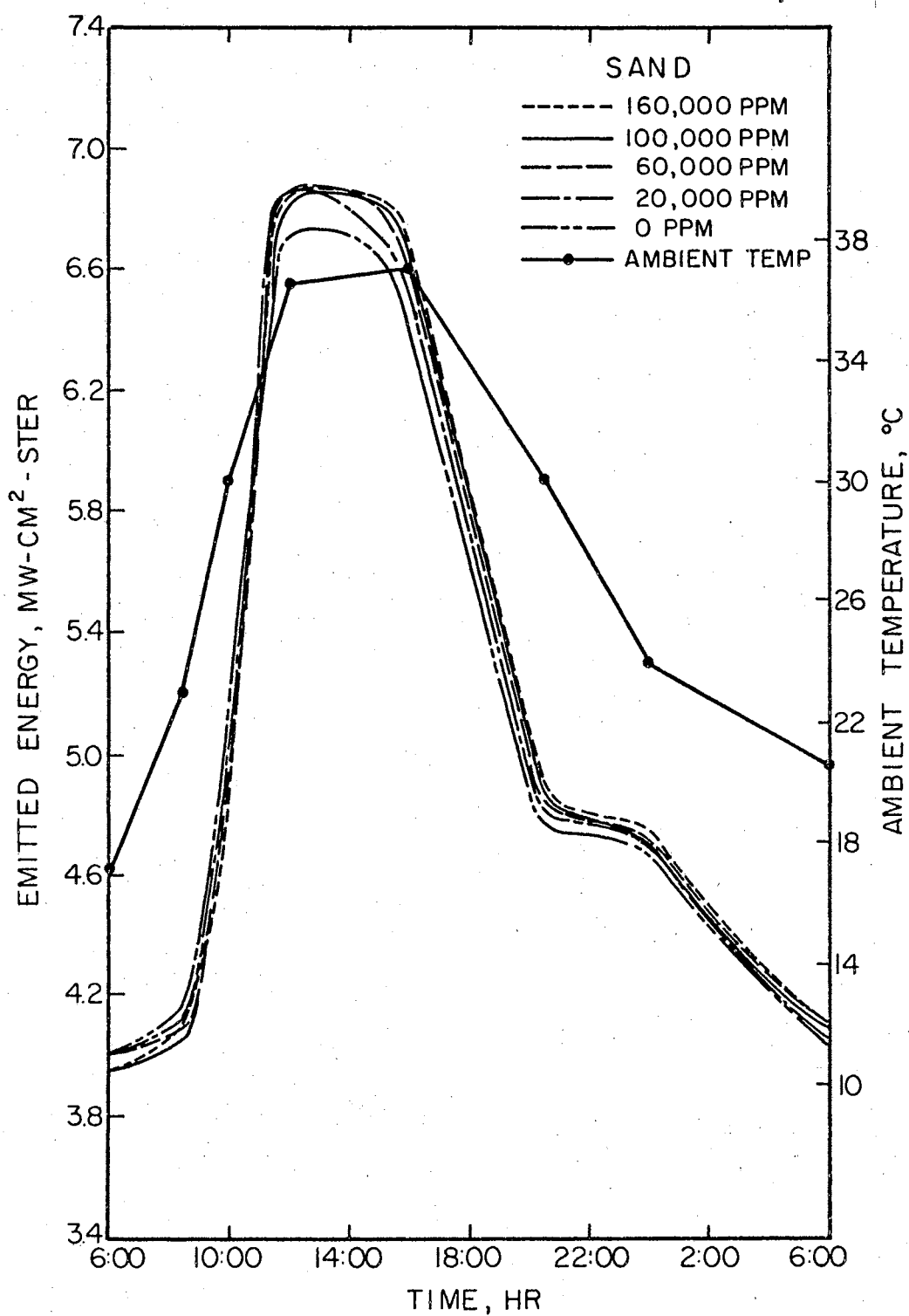


Figure 51. Emitted Energy for June 7 for Saline Water Table at 1 ft. in Sand

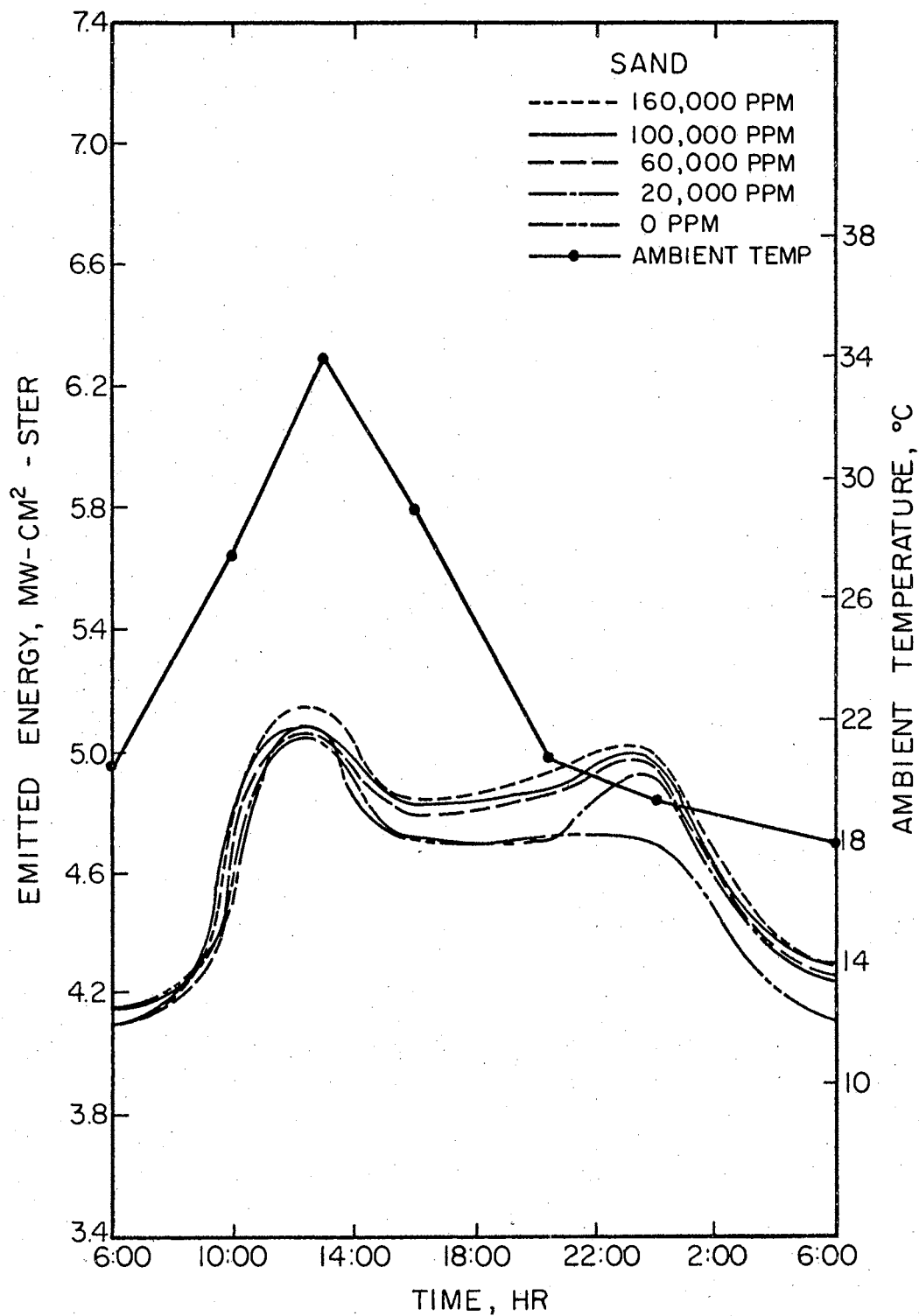


Figure 52. Emitted Energy for June 8 for Saline Water Table at 0 ft. in Sand

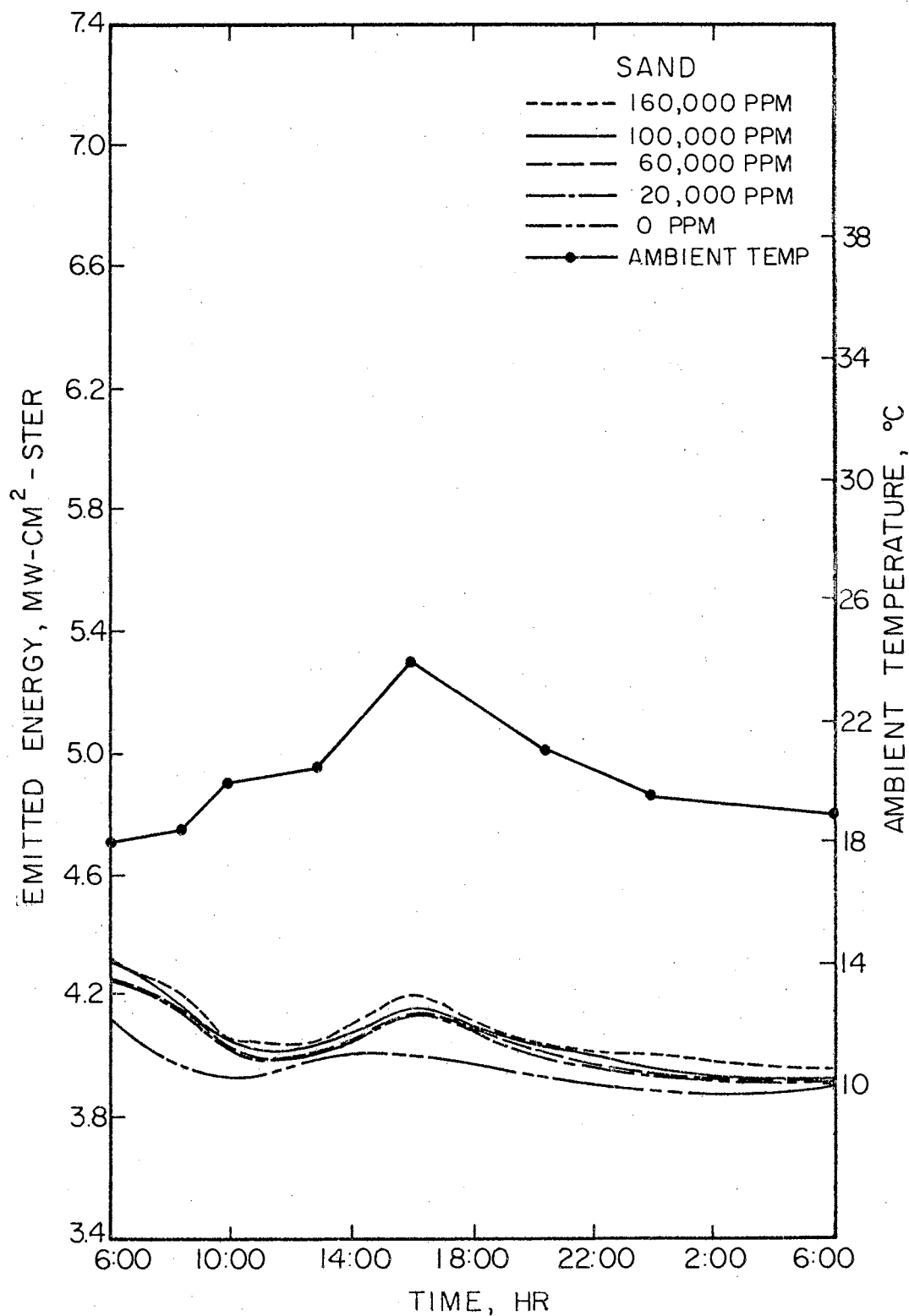


Figure 53. Emitted Energy for June 9 for Saline Water Table at 0 ft. in Sand

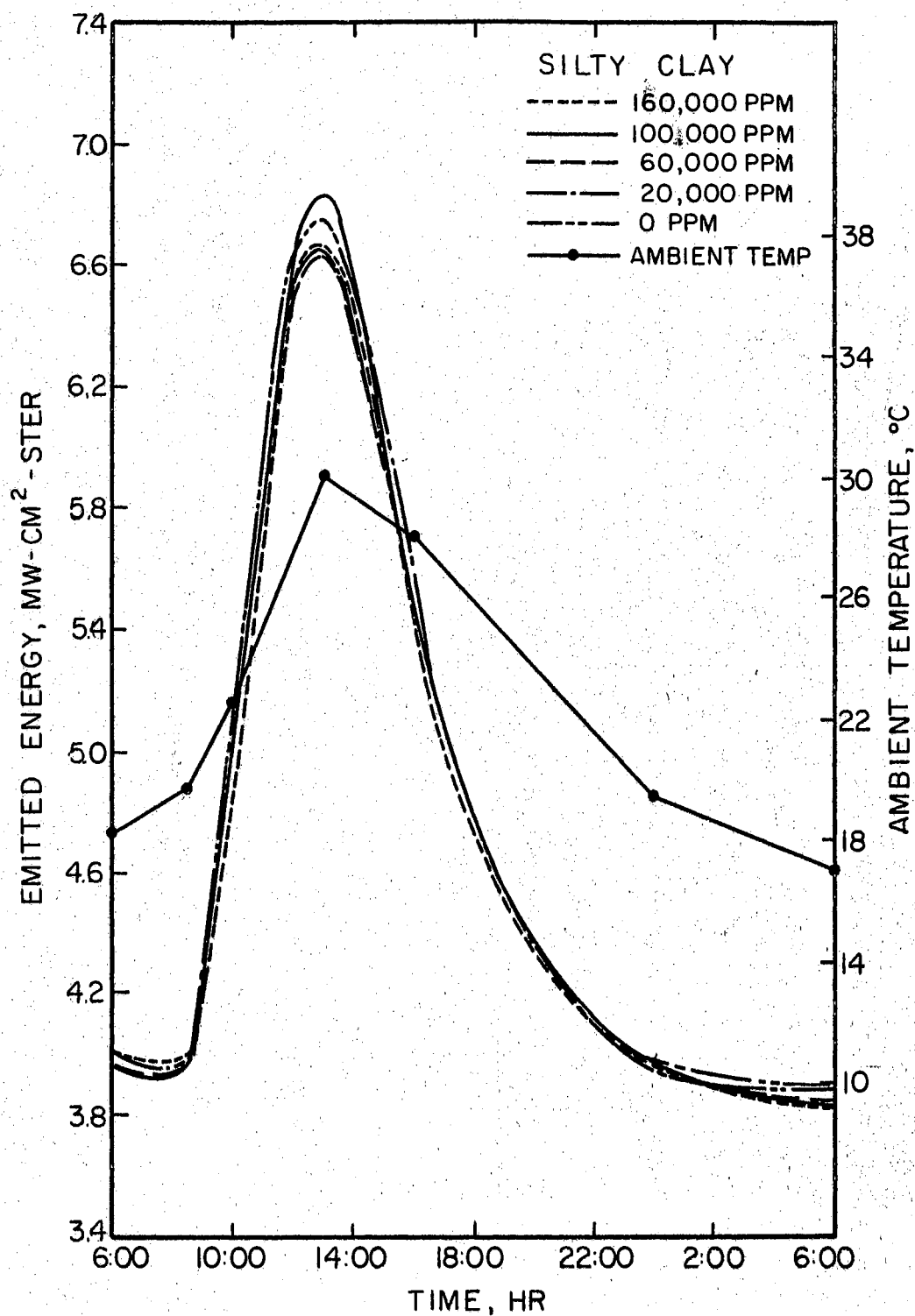


Figure 54. Emitted Energy for June 13 for Saline Water Table at 4 ft. in Silty Clay

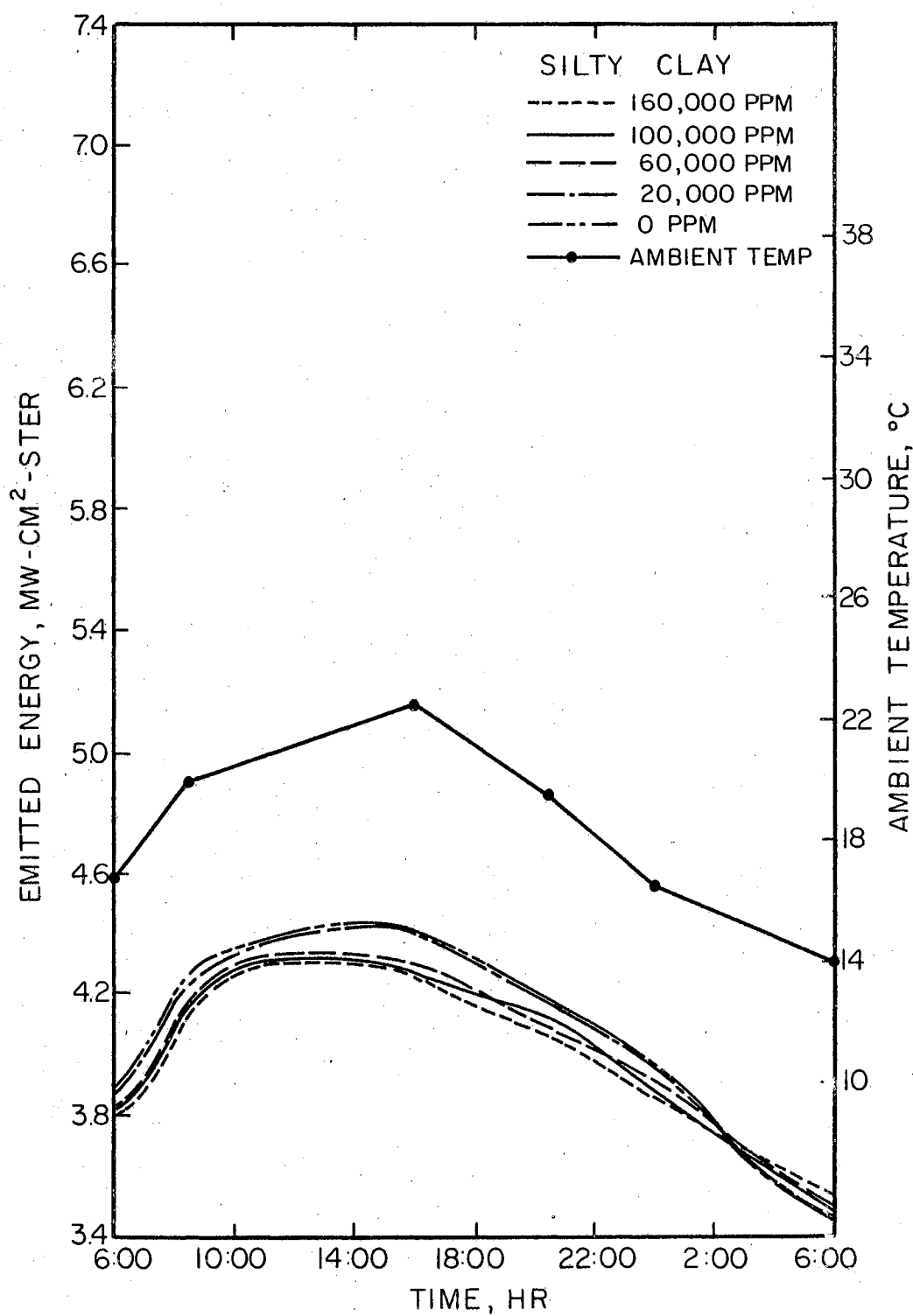


Figure 55. Emitted Energy for June 14 for Saline Water Table at 4 ft. in Silty Clay

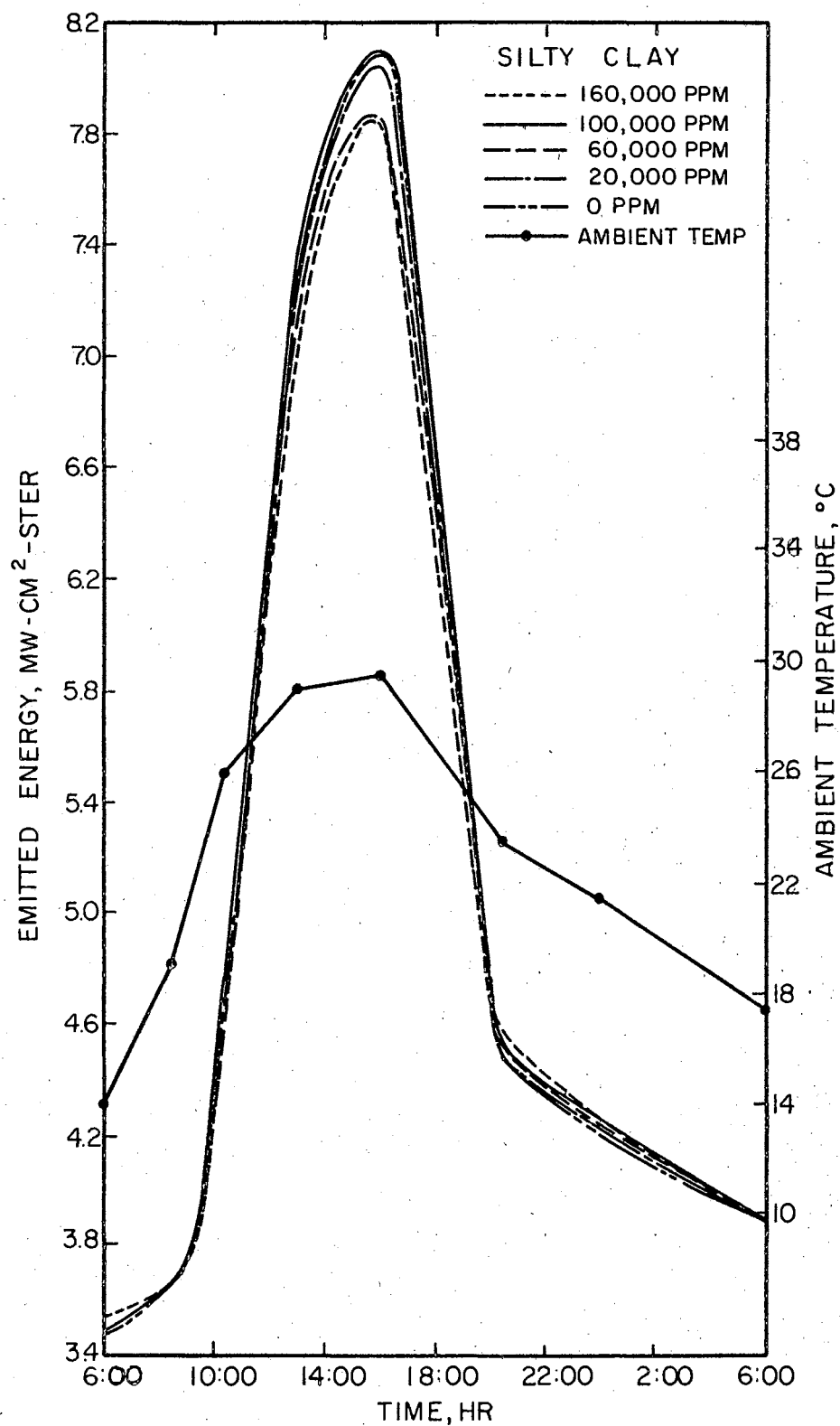


Figure 56. Emitted Energy for June 15 for Saline Water Table at 3 ft. in Silty Clay

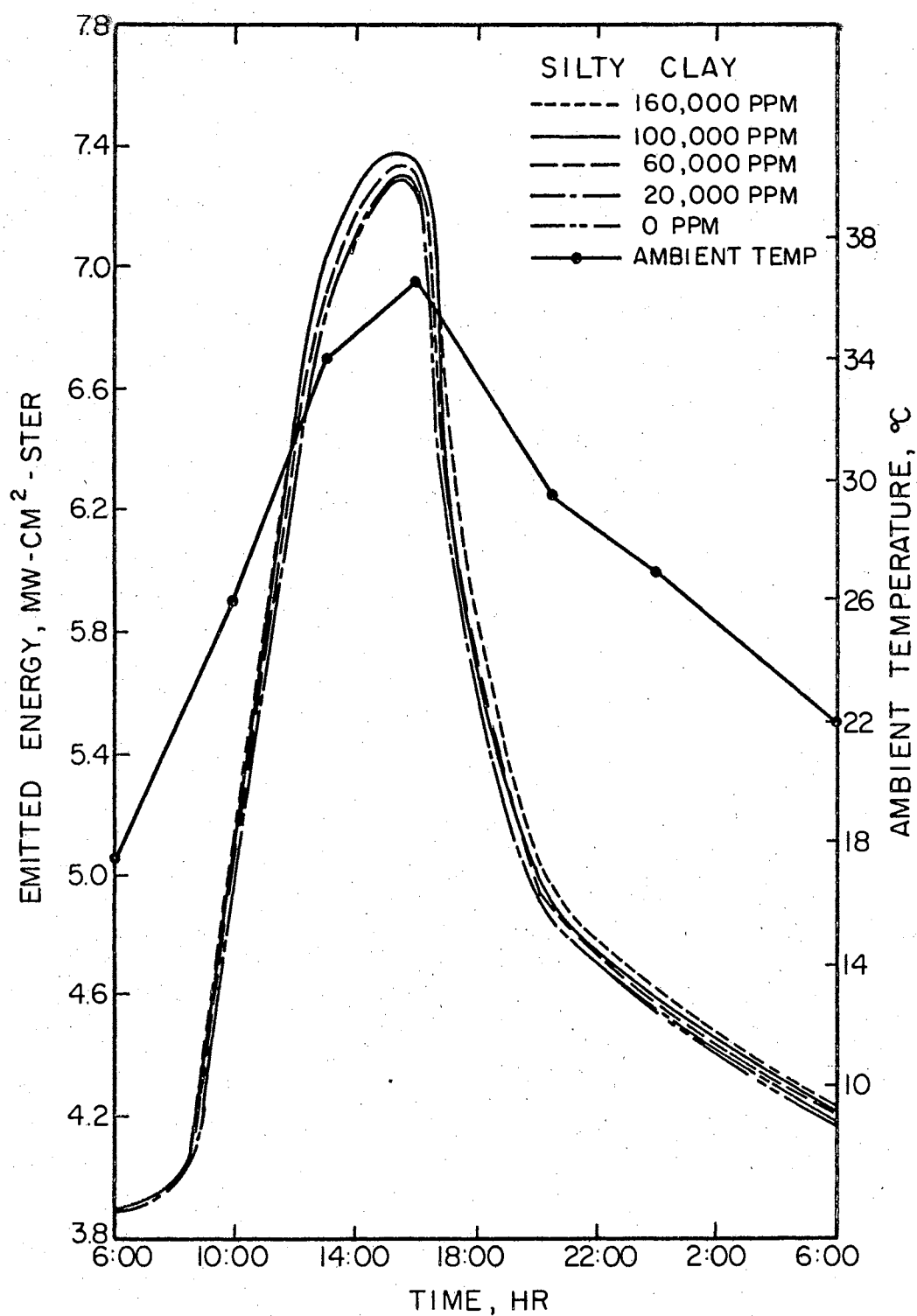


Figure 57. Emitted Energy for June 16 for Saline Water Table at 3 ft. in Silty Clay

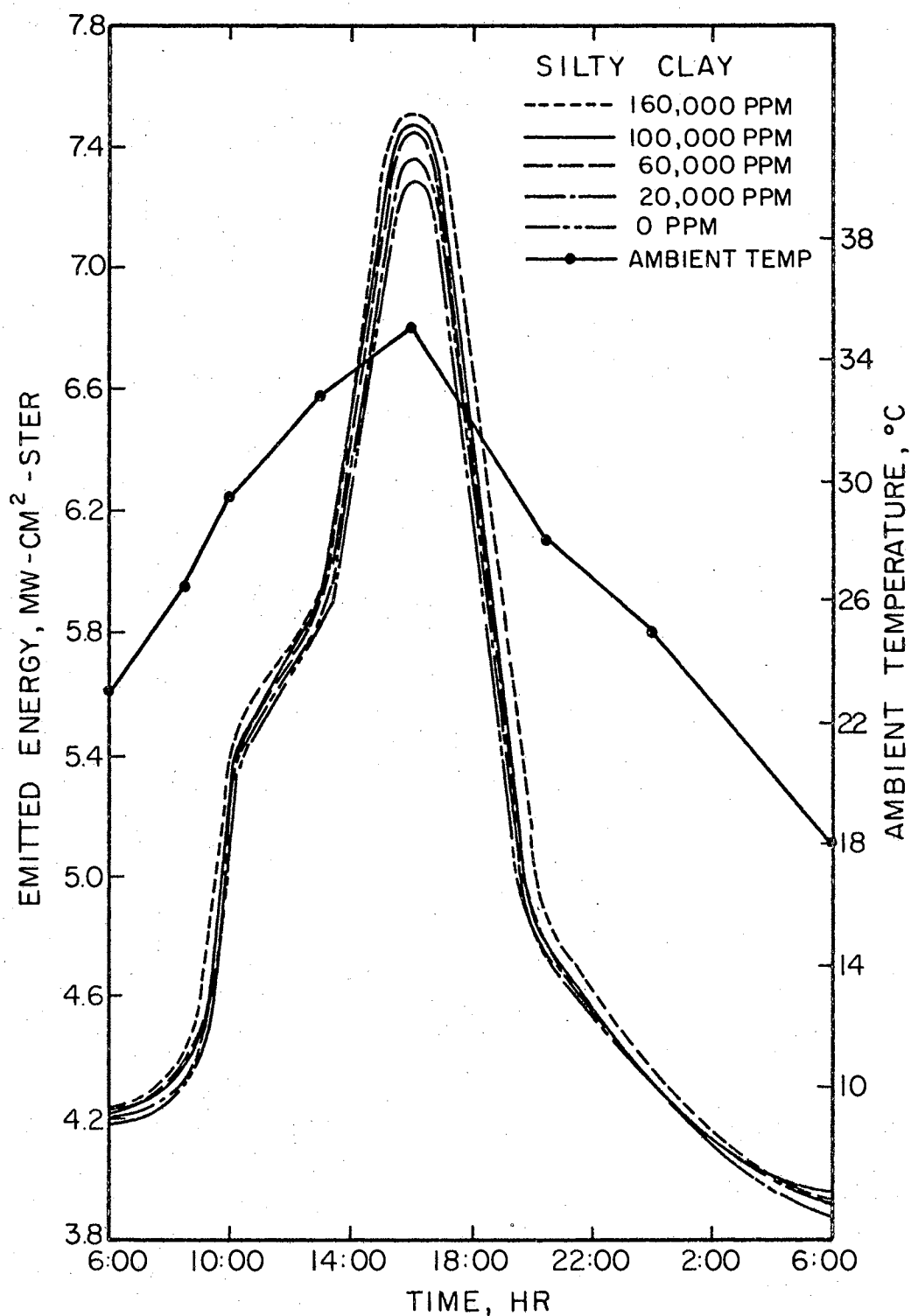


Figure 58. Emitted Energy for June 17 for Saline Water Table at 2 ft. in Silty Clay

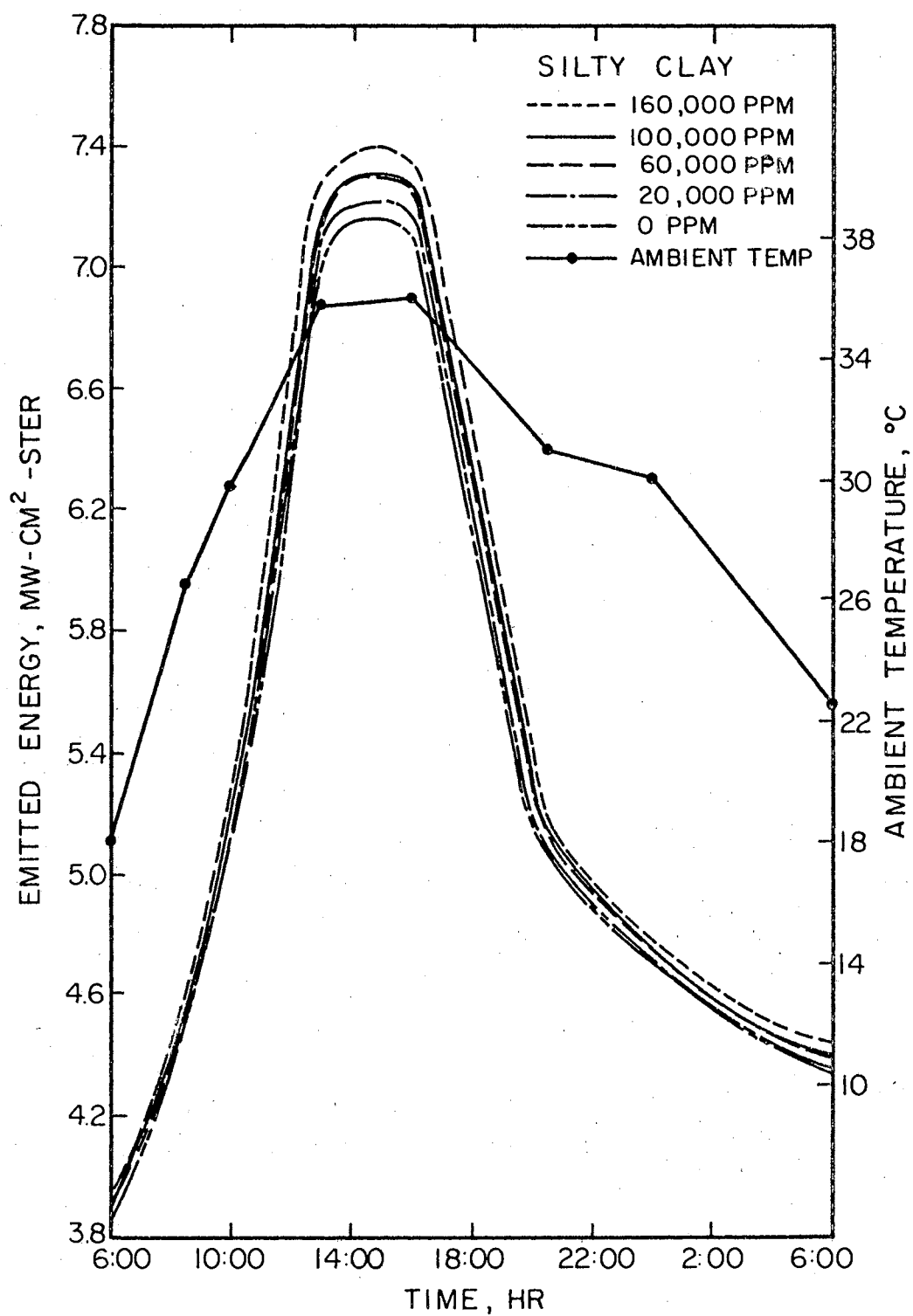


Figure 59. Emitted Energy for June 18 for Saline Water Table at 2 ft. in Silty Clay

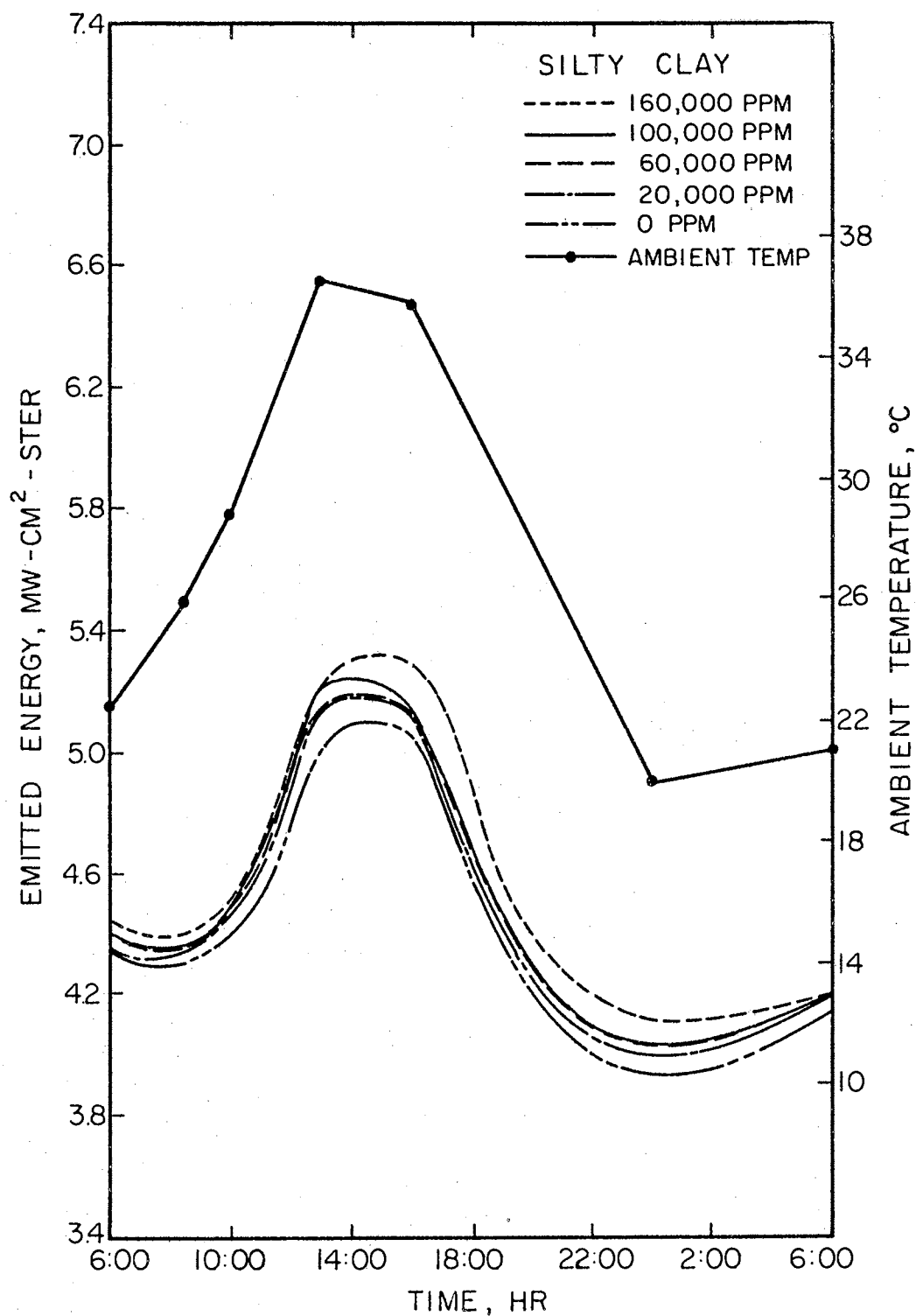


Figure 60. Emitted Energy for June 19 for Saline Water Table at 1 ft. in Silty Clay

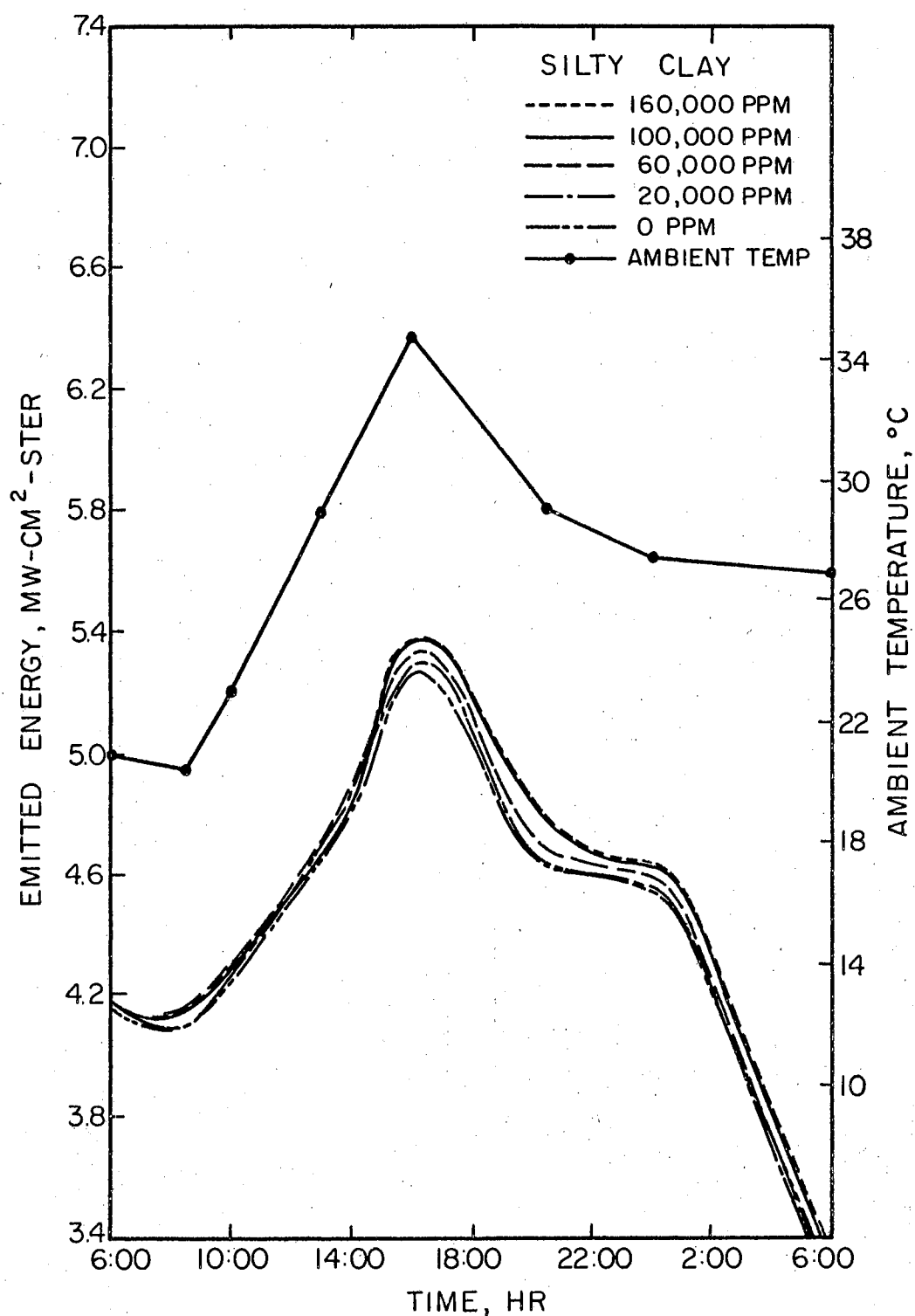


Figure 61. Emitted Energy for June 20 for Saline Water Table at 1 ft. in Silty Clay

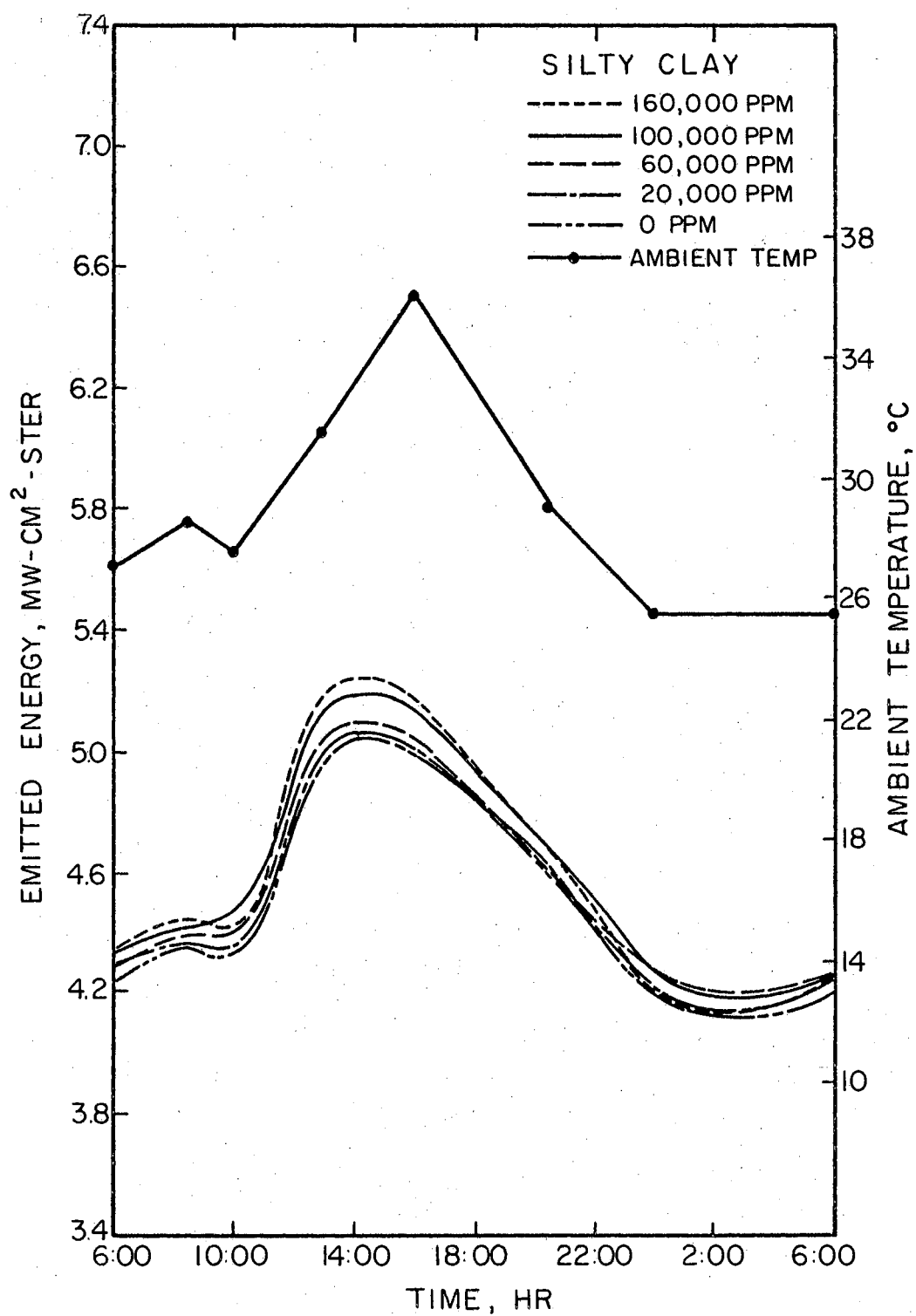


Figure 62. Emitted Energy for June 21 for Saline Water Table at 0 ft. in Silty Clay

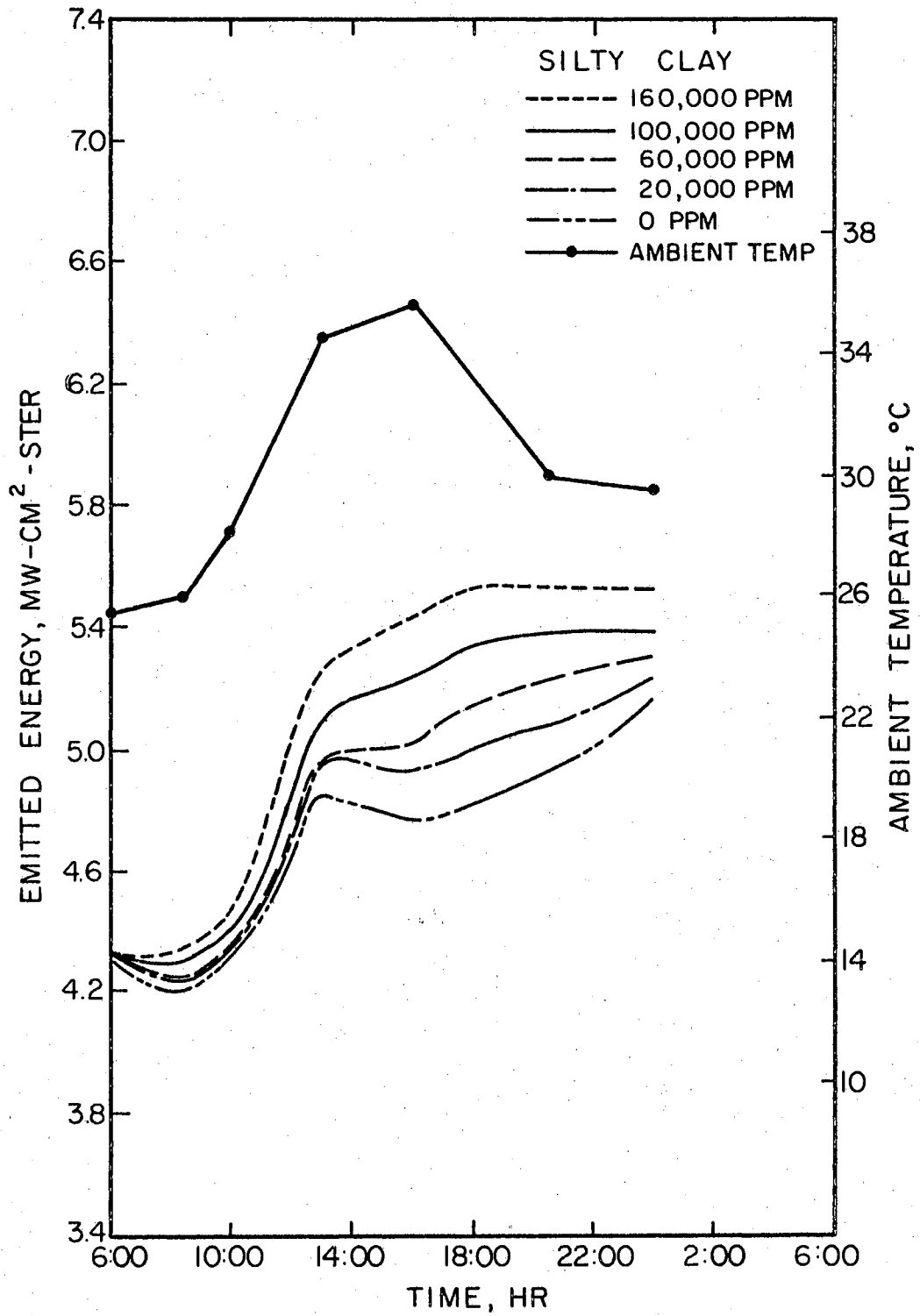


Figure 63. Emitted Energy for June 22 for Saline Water Table at 0 ft. in Silty Clay

CHAPTER VII

CONCLUSIONS AND RECOMMENDATIONS

Conclusions

This investigation was conducted (a) to study the emittance qualities of sand and silty clay containing varying salt concentrations and changing ground water tables, and (b) to study the reflectance variances of wheat and salt grass grown in soils containing varied salt concentrations. The main objects of this research were (a) to determine the best film-filter combination for measuring reflectance variances of vegetation grown in soils containing varied amounts of salt; (b) to determine if soils containing varied salt concentrations and soils with changing ground water tables could be detected by emittance measurements, and if so (c) to determine the best time to take an imagery for measuring emitted energy of soils.

From the analysis of the data discussed in the preceeding chapter the following conclusions can be drawn:

(1) The maximum difference in reflectance of both soil types of vegetation grown in soils containing varied salt concentrations occurred primarily in the near infrared region, 720-1000 millimicron wavelengths, of the electromagnetic spectrum. Therefore, Ektachrome IR film with No. 12 and No. 25 filters were used to take photographs of wheat samples. The variances in reflectance were detectable in the photographs

taken of samples in the lab. However, since a larger area and more than one layer of leaves will be photographed of vegetation in the field, these reflectance variances should be magnified. Also, the energy source for field photography will be solar energy which will result in a change of exposure time and focal length settings. Photographs for detection of salinity in soils by measuring reflectance of vegetation is best done when there is a good canopy of leaves on the plants in order to minimize ground detection. Results showed maturity to be quite instrumental in obtaining favorable reflectance qualities, however, the photographs will give best color contrasts if taken before leaves enter stages of yellowing. For this reason, detection of salinity in soils in which wheat is grown will give best results if wheat is photographed shortly before heading occurs. Aerial photography should be taken in the morning at a low altitude flight to incur a minimal amount of atmospheric haze interfering with reflectance.

(2) The maximum difference in emittances of both soil types with varied salt concentrations occurred between 2:00 and 4:00 in the afternoon. These results show the measuring of emitted energy between the atmospheric window 8 - 14 micron wavelengths by remote sensing techniques will give best detection when testing is conducted at this time. Although some of the readings taken during this investigation were erratic, the emittance variances of soils with varied salt concentrations in the field may be detectable by imagery. The field area of soil emitting the energy will be much larger and should lessen the amount of erratic readings which were detected in the soil panels. Although imagery will not give the amount of salt concentration in the soil, quantitatively, it will give the boundary of each change in salt concentration in the soil

by changing shades. After an imagery has been taken and the boundaries for the different salt concentrations have been determined, samples of soils can be taken within these particular boundaries to determine the amount of salt contained by the soil. In the field, as the soil types change so will the emitted energy change; therefore, the boundary of the soil types will also have to be determined.

(3) The maximum difference in emittance for changing non-saline water tables also occurred between 2:00 and 4:00 p.m. Therefore, detection of emittance of changing water tables by imagery may be conducted during this time. The changes in emitted energy for the different levels of water table depend upon how close the water level is to the surface and how much capillary rise will occur in the particular soil from which emittance measurements are taken. In other words, the greater the capillary rise, the greater the depth of water table that can be detected by the difference in emitted energy. An imagery will not give the depth to water table, quantitatively, but it will show the change in water table if the water level is close enough to the surface to effect the emitted energy of the soil.

(4) Differences in emittance for varied saline water tables were slight and some readings were erratic. There was an inversion of emittance when the water table reached a point where evaporation controlled the emitted energy given off by the surface. Water tables at the same level, containing different concentrations of salt, may give better detection by remote sensing techniques when the water table is closer to the surface.

(5) It should be taken into consideration that the four variables under investigation were tested separately. However, in measuring the

emittance of soils in the field, all four variables may be affecting the emittance measurements simultaneously.

Recommendations for Future Research

The following are suggestions for further research:

- (1) Different types of vegetation should be tested since the reflectance changes with different types of vegetation.
- (2) Photographs should be taken in the field to determine how well vegetation grown under field conditions can be used as detectors for salinity in soils.
- (3) Imagery should be taken of soils in the field to determine how the emitted energy of soils is affected when more than one of the variables studied in this investigation is present under field conditions.

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APPENDIX A

REFLECTANCE READINGS OF WHEAT GROWN IN SILTY
CLAY AT VARIED LEVELS OF SALINITY

Wave-length (mμ)	0 ppm Level		2,000 ppm Level		4,000 ppm Level		6,000 ppm Level		8,000 ppm Level		10,000 ppm Level	
	Max. %	Min. %	Max. %	Min. %	Max. %	Min. %	Max. %	Min. %	Max. %	Min. %	Max. %	Min. %
200	7.8	7.7	7.9	7.7	7.8	7.5	7.9	7.6	7.5	7.8	7.5	7.5
250	7.3	7.0	7.3	7.4	7.2	7.1	7.2	7.0	7.3	7.1	7.2	7.1
300	4.6	4.5	4.6	4.5	4.6	4.7	4.5	4.5	4.5	4.4	4.5	4.4
350	3.5	3.4	3.5	3.5	3.6	3.5	3.5	3.6	3.5	3.5	3.3	3.4
400	4.7	5.1	4.5	4.6	4.7	4.5	4.5	4.9	4.6	4.9	4.8	4.4
450	6.5	7.1	6.4	7.1	6.4	7.1	6.4	6.9	6.8	6.4	6.9	6.3
500	8.5	8.3	8.5	8.5	8.4	8.5	8.5	8.3	8.4	8.4	8.4	8.6
550	19.0	18.5	17.5	17.7	17.5	17.0	16.7	16.0	15.4	15.1	14.9	15.1
600	14.7	14.1	14.2	14.1	14.1	13.8	14.0	13.9	13.9	13.7	13.5	13.0
650	16.0	9.2	9.1	9.1	9.0	8.7	8.7	8.9	8.6	8.4	8.5	8.4
700	21.0	21.6	18.0	17.8	17.9	17.2	17.5	16.9	17.1	16.7	16.7	16.0
750	53.0	51.9	51.0	50.6	50.1	50.0	47.0	47.2	46.5	46.8	46.5	46.1
800	52.4	51.7	50.9	50.6	50.7	50.1	49.0	48.0	47.0	47.0	46.6	46.0
850	52.0	51.3	50.1	49.7	49.7	49.2	48.0	47.1	46.3	46.1	46.0	45.4
900	51.2	49.7	48.9	48.6	48.1	47.4	46.5	46.0	45.8	45.0	45.3	44.7
950	51.0	49.2	47.9	46.0	47.0	47.0	46.0	44.7	44.4	43.7	43.9	43.0
1000	50.4	48.6	47.5	46.0	46.1	45.6	45.1	44.0	43.8	42.6	41.6	40.5

REFLECTANCE READINGS OF SALT GRASS GROWN
AT VARIED LEVELS OF SALINITY

Wave-length (mμ)	Grown in Sand								Grown in Silty Clay							
	0 ppm Level		20,000 ppm Level		60,000 ppm Level		80,000 ppm Level		0 ppm Level		20,000 ppm Level		60,000 ppm Level			
	Max.%	Min.%	Max.%	Min.%	Max.%	Min.%	Max.%	Min.%	Max.%	Min.%	Max.%	Min.%	Max.%	Min.%	Max.%	Min.%
200	7.1	7.2	7.2	7.4	7.4	7.7	7.1	7.3	7.4	7.1	7.3	7.1	7.4	7.6		
250	5.6	6.0	5.8	6.2	6.1	6.6	6.5	6.8	6.0	5.4	5.8	5.3	5.2	5.5		
300	4.5	4.1	4.3	4.1	4.6	4.5	4.5	4.1	4.4	4.2	4.4	4.0	4.1	4.3		
350	3.2	3.0	3.5	3.1	3.5	3.0	3.6	3.2	3.4	3.1	3.2	3.0	3.5	3.1		
400	5.1	4.6	4.8	5.1	5.0	4.7	5.5	5.2	5.5	4.5	5.7	5.4	5.9	5.4		
450	7.0	6.4	7.0	6.2	6.5	6.2	6.8	7.0	7.0	6.2	7.1	6.5	7.3	7.5		
500	9.3	9.0	8.9	8.6	8.5	8.0	8.9	9.1	9.4	9.1	8.8	9.0	9.0	8.7		
550	15.3	15.0	14.9	14.7	15.1	14.8	14.2	14.6	16.2	16.4	15.5	16.0	15.1	15.8		
600	10.7	10.0	9.8	9.7	10.0	9.6	10.0	9.5	10.1	10.5	10.2	10.3	10.1	10.3		
650	8.6	8.5	7.7	7.7	7.8	7.2	7.5	7.4	8.7	8.9	8.6	8.1	8.5	8.7		
700	18.3	18.0	15.2	14.5	15.0	14.1	15.0	14.6	17.8	17.5	17.2	17.5	17.5	17.1		
750	48.0	46.5	45.9	43.3	44.0	43.2	42.5	42.0	48.2	46.0	43.9	42.4	41.9	40.7		
800	50.4	48.3	46.8	44.8	45.0	44.5	44.0	44.3	48.9	46.5	45.5	43.5	43.7	42.5		
850	49.3	45.7	46.6	45.0	45.5	44.7	43.1	43.4	50.5	48.1	46.4	44.6	44.3	42.6		
900	48.5	46.0	44.0	43.8	43.7	43.0	42.2	42.5	48.4	46.3	45.3	44.0	43.2	42.5		
950	46.8	44.6	42.2	41.7	42.0	41.1	41.0	41.2	46.2	44.7	44.0	43.2	42.5	41.0		
1000	45.4	44.2	39.7	40.1	40.0	39.2	37.7	37.0	44.4	43.0	41.7	40.0	39.0	38.1		

RELATIVE TURGIDITY OF VEGETATION

Wheat Grown in Silty Clay

<u>Salt in ppm</u>	<u>Relative Turgidity</u>
0	98.5%
2,000	96.3%
4,000	94.0%
6,000	93.5%
8,000	92.3%
10,000	90.3%

Salt Grass Grown in Sand Mixture

<u>Salt in ppm</u>	<u>Relative Turgidity</u>
0	97.2%
20,000	95.2%
60,000	92.0%
100,000	90.1%

Salt Grass Grown in Silty Clay

<u>Salt in ppm</u>	<u>Relative Turgidity</u>
0	97.2%
20,000	93.5%
60,000	91.0%

APPENDIX B

APPENDIX TERMS

Salt in ppm Parts Per Million

Ambient Temperature Degrees Centigrade

Meteorological Conditions PC - Partly Cloudy
R - Rain
C - Clear
O - Overcast

Emitted Energy ($\text{mw-cm}^2\text{-ster}$)

EMITTED ENERGY OF SAND WITH VARIANCE
IN MOISTURE AND SALINITY

Date	Time	Ambient Temp.	Met. Cond.	EMITTED ENERGY							
				Salt Content in ppm							
				0	20,000	40,000	60,000	80,000	100,000	130,000	160,000
9-21-68	5:00	14.3	C	4.70	4.71	4.71	4.71	4.72	4.77	4.77	4.71
	8:00	14.3	C	4.71	4.71	4.71	4.72	4.72	4.77	4.78	4.72
	10:00	24.1	C	4.90	4.80	4.90	4.91	4.91	4.91	4.91	4.93
	12:00	25.5	PC	5.10	4.98	5.14	5.18	5.15	5.14	5.14	5.16
	14:00	27.2	C	5.26	5.21	5.29	5.31	5.30	5.30	5.30	5.31
	16:00	27.4	C	5.42	5.32	5.35	5.37	5.37	5.29	5.29	5.29
	19:00	22.2	C	5.45	5.40	5.39	5.37	5.38	5.32	5.33	5.32
	24:00	17.0	R	5.44	5.43	5.41	5.40	5.39	5.36	5.36	5.35
9-22-68	5:00	10.8	C	5.15	5.15	5.12	5.10	5.12	5.08	5.09	5.09
	8:00	10.6	C	5.15	5.12	5.12	5.15	5.20	5.24	5.28	5.29
	10:00	17.8	C	5.56	5.60	5.60	5.61	5.61	5.62	5.63	5.64
	12:00	19.8	C	5.65	5.70	5.70	5.76	5.76	5.72	5.72	5.76
	14:00	20.2	C	5.02	5.04	5.03	5.09	5.14	5.09	5.09	5.15
	16:00	20.2	C	5.04	5.02	5.01	5.02	5.07	5.01	5.04	5.05
	19:00	14.8	C	4.78	4.78	4.79	4.80	4.80	4.78	4.80	4.81
	24:00	8.2	C	4.55	4.57	4.59	4.60	4.60	4.60	4.60	4.61
9-23-68	5:00	7.2	C	4.47	4.47	4.47	4.48	4.48	4.48	4.48	4.48
	8:00	8.0	C	4.45	4.45	4.46	4.46	4.46	4.46	4.47	4.47
	12:00	19.4	O	4.76	4.77	4.77	4.80	4.81	4.81	4.82	4.84
	14:00	19.0	O	4.85	4.86	4.86	4.88	4.89	4.90	4.90	4.91
	16:00	16.4	O	4.81	4.85	4.84	4.85	4.87	4.85	4.86	4.87
	19:00	12.4	C	4.71	4.71	4.71	4.71	4.74	4.74	4.74	4.73
	24:00	8.9	C	4.53	4.53	4.53	4.53	4.53	4.53	4.53	4.53
9-24-68	5:00	5.8	C	4.42	4.42	4.42	4.42	4.44	4.44	4.44	4.44
	8:00	5.1	C	4.36	4.36	4.36	4.37	4.37	4.38	4.37	4.40
	12:00	13.0	PC	4.81	4.84	4.85	4.85	4.85	4.81	4.82	4.84
	14:00	14.5	PC	4.86	4.86	4.86	4.88	4.90	4.92	4.91	4.92
	16:00	16.0	C	4.81	4.85	4.85	4.85	4.90	4.87	4.91	4.92
	19:00	11.4	C	4.64	4.67	4.67	4.68	4.68	4.67	4.68	4.68
	24:00	5.8	C	4.47	4.47	4.48	4.48	4.50	4.51	4.51	4.52
9-25-68	5:00	3.3	C	4.35	4.36	4.36	4.36	4.36	4.36	4.36	4.36
	8:00	4.4	C	4.34	4.34	4.34	4.34	4.34	4.34	4.35	4.34
	12:00	21.0	C	4.91	4.96	4.96	5.00	5.01	5.05	5.04	5.05
	14:00	23.0	C	5.14	5.20	5.20	5.24	5.26	5.26	5.28	5.28
	16:00	24.6	C	5.39	5.40	5.37	5.38	5.40	5.36	5.35	5.44
	19:00	17.0	C	5.36	5.39	5.37	5.37	5.37	5.37	5.36	5.40
	24:00	10.4	C	5.21	5.21	5.21	5.24	5.28	5.30	5.31	5.34
9-26-68	8:00	6.2	C	4.42	4.42	4.42	4.43	4.43	4.43	4.43	4.43
	12:00	26.0	C	5.04	5.04	5.04	5.08	5.07	5.12	5.11	5.12
	14:00	28.2	C	5.28	5.27	5.27	5.34	5.37	5.37	5.42	5.42
	16:00	30.0	C	5.20	5.24	5.22	5.26	5.27	5.26	5.34	5.34
	19:00	20.0	C	4.93	4.95	4.95	4.96	5.03	5.03	5.04	5.05
	24:00	17.8	C	4.78	4.78	4.80	4.82	4.84	4.86	4.85	4.84

EMITTED ENERGY OF SILTY CLAY WITH VARIANCE
IN MOISTURE AND SALINITY

Date	Time	Ambient Temp.	Met. Cond.	EMITTED ENERGY							
				Salt Content in ppm							
				0	20,000	40,000	60,000	80,000	100,000	130,000	160,000
9-21-68	5:00	14.3	C	4.70	4.71	4.71	4.71	4.71	4.71	4.71	4.72
	8:00	14.3	C	4.71	4.71	4.71	4.71	4.71	4.70	4.71	4.72
	10:00	24.1	C	4.87	4.89	4.90	4.92	4.93	4.92	4.95	4.94
	12:00	25.5	PC	5.08	5.10	5.07	5.12	5.16	5.16	5.18	5.18
	14:00	27.2	C	5.25	5.30	5.26	5.30	5.30	5.30	5.34	5.30
	16:00	27.4	C	5.64	5.69	5.58	5.60	5.55	5.56	5.56	5.60
	19:00	22.2	C	5.77	5.75	5.68	5.68	5.64	5.60	5.63	5.57
	24:00	17.0	R	5.52	5.49	5.49	5.50	5.50	5.50	5.40	5.50
9-22-68	5:00	10.8	C	5.37	5.37	5.35	5.35	5.30	5.30	5.26	5.21
	8:00	10.6	C	5.35	5.33	5.30	5.29	5.25	5.24	5.21	5.22
	10:00	17.8	C	5.60	5.57	5.60	5.64	5.64	5.63	5.64	5.64
	12:00	19.8	C	5.64	5.68	5.68	5.72	5.73	5.68	4.76	5.75
	14:00	20.2	C	4.89	5.03	4.96	5.04	4.95	5.00	5.09	5.16
	16:00	20.2	C	4.97	5.10	5.03	5.15	5.08	5.06	5.20	5.22
	19:00	14.8	C	4.72	4.80	4.75	4.82	4.78	4.82	4.85	4.88
	24:00	8.2	C	4.55	4.57	4.57	4.55	4.58	4.60	4.61	4.60
9-23-68	5:00	7.2	C	4.45	4.45	4.44	4.47	4.47	4.47	4.50	4.50
	8:00	8.0	C	4.42	4.43	4.42	4.43	4.45	4.44	4.45	4.46
	12:00	19.4	O	4.75	4.76	4.75	4.75	4.76	4.76	4.80	4.81
	14:00	19.0	O	4.80	4.87	4.83	4.86	4.89	4.89	4.90	4.93
	16:00	16.4	O	4.75	4.82	4.80	4.85	4.87	4.85	4.86	4.89
	19:00	12.4	C	4.65	4.70	4.67	4.72	4.72	4.72	4.75	4.75
	24:00	8.9	C	4.52	4.55	4.54	4.56	4.56	4.56	4.56	4.57
9-24-68	5:00	5.8	C	4.41	4.42	4.41	4.44	4.45	4.43	4.42	4.42
	8:00	5.1	C	4.36	4.37	4.37	4.37	4.39	4.44	4.40	4.40
	12:00	13.0	PC	4.75	4.81	4.78	4.83	4.82	4.84	4.87	4.87
	14:00	14.5	PC	4.80	4.88	4.95	4.88	4.91	4.91	4.96	4.93
	16:00	16.0	C	4.84	4.94	4.96	4.88	4.92	4.92	4.96	5.00
	19:00	11.4	C	4.60	4.65	4.63	4.67	4.70	4.67	4.70	4.72
	24:00	5.8	C	4.44	4.47	4.47	4.50	4.50	4.50	4.50	4.50
9-25-68	5:00	3.3	C	4.32	4.32	4.32	4.33	4.36	4.36	4.36	4.36
	8:00	4.4	C	4.32	4.32	4.31	4.32	4.32	4.32	4.32	4.32
	12:00	21.0	C	4.93	4.98	4.95	4.96	5.02	5.02	5.02	5.08
	14:00	23.0	C	5.12	5.27	5.21	5.30	5.26	5.34	5.35	5.31
	16:00	24.6	C	5.64	5.72	5.68	5.77	5.73	5.75	5.76	5.75
	19:00	17.0	C	5.50	5.60	5.58	5.57	5.56	5.56	5.56	5.56
	24:00	10.4	C	5.32	5.37	5.30	5.37	5.34	5.35	5.37	5.36
9-26-68	8:00	6.2	C	4.41	4.41	4.41	4.44	4.43	4.43	4.44	4.44
	12:00	26.0	C	5.04	5.10	5.08	5.10	5.09	5.10	5.10	5.15
	14:00	28.2	C	5.21	5.40	5.26	5.39	5.34	5.40	5.43	5.36
	16:00	30.0	C	5.24	5.42	5.32	5.44	5.36	5.46	5.49	5.45
	19:00	20.0	C	4.89	5.00	5.32	5.04	5.06	5.08	5.08	5.07
	24:00	17.8	C	4.76	4.82	4.79	4.85	4.85	4.86	4.88	4.87

MOISTURE VARIATION IN SOILS

Silty Clay

Salt in ppm	PERCENT OF MOISTURE CONTENT							
	Sept. 20	Sept. 21	Sept. 22	Sept. 23	Sept. 24	Sept. 25	Sept. 26	Sept. 27
0	36.10	26.0	21.5	18.24	15.748	13.2	11.248	9.748
20,000	36.00	27.752	24.25	21.00	19.748	18.2	15.748	14.248
40,000	36.25	29.0	26.748	25.0	23.748	21.5	19.248	17.8
60,000	36.12	30.248	28.748	27.0	25.75	24.3	21.748	20.0
80,000	36.00	31.748	29.0	28.5	27.25	26.5	23.75	22.25
100,000	35.88	33.248	32.0	29.248	28.248	26.748	24.5	22.5
130,000	36.05	34.248	32.248	31.248	30.748	30.248	28.748	26.5
160,000	36.12	34.0	33.0	32.0	31.20	30.0	28.5	27.0

Sand

Salt in ppm	PERCENT OF MOISTURE CONTENT							
	Sept. 20	Sept. 21	Sept. 22	Sept. 23	Sept. 24	Sept. 25	Sept. 26	Sept. 27
0	18.05	15.625	14.061	13.194	12.321	11.283	9.89	8.53
20,000	18.05	15.8	14.408	13.366	12.50	11.458	10.009	9.76
40,000	18.23	16.319	14.933	14.063	13.369	12.675	11.458	10.51
60,000	18.12	16.75	15.277	14.411	13.88	12.847	11.630	10.76
80,000	18.05	17.013	15.625	14.583	14.063	13.019	12.327	11.286
100,000	18.25	17.361	15.80	14.933	14.411	13.369	12.5	11.286
130,000	18.00	17.42	16.147	15.452	14.758	14.063	13.194	11.98
160,000	18.05	17.5	16.4	15.625	15.105	14.408	13.54	12.5

APPENDIX C

EMITTANCE MEASUREMENTS OF CHANGING WATER TABLE
FOR SILTY CLAY WITH NON-SALINE WATER

Date	Time	Ambient Temp.	Met. Cond.	E M I T T E D E N E R G Y				
				Water Level				
				0 ft.	1 ft.	2 ft.	3 ft.	4 ft.
5-9-69	5:30	10.2	C	3.36	3.39	3.36	3.39	3.37
	8:30	13.0	C	3.36	3.46	3.36	3.42	3.38
	10:00	21.0	C	3.90	4.18	4.54	5.04	5.08
	13:00	24.4	C	4.72	4.82	6.48	6.48	6.55
	16:00	26.0	C	4.51	4.67	6.16	6.36	6.49
	20:30	20.6	C	4.03	4.18	4.30	4.34	4.27
	24:00	16.4	C	3.78	3.81	3.87	3.85	3.85
5-10-69	5:30	15.8	C	3.58	3.58	3.57	3.57	3.57
	8:30	16.8	O	3.64	3.64	3.78	3.78	3.78
	10:00	17.8	C	3.90	3.93	4.30	4.36	4.36
	13:00	23.8	C	4.57	4.85	5.89	6.08	4.16
	16:00	21.4	C	4.37	4.64	5.30	5.07	5.08
	20:30	17.5	C	3.90	4.06	4.09	4.10	3.94
	24:00	14.2	C	3.69	3.68	3.46	3.43	3.42
5-11-69	5:30	12.6	C	3.52	3.52	3.40	3.48	3.38
	11:30	23.0	PC	4.10	4.50	5.30	5.39	5.49
	13:00	24.0	C	4.40	4.80	6.11	6.07	6.36
	16:00	26.8	C	4.48	4.96	6.19	6.28	6.36
	20:30	23.1	PC	4.27	4.44	4.56	4.56	4.56
	24:00	18.8	C	4.20	4.22	4.22	4.14	4.13
5-12-69	6:00	14.7	C	4.04	3.92	3.74	3.68	3.64
	8:30	18.3	C	3.90	4.08	3.84	3.88	3.84
	10:00	23.0	C	4.32	4.76	4.78	5.38	5.42
	13:00	29.4	C	5.00	5.42	7.50	7.55	7.60
	16:00	31.7	C	4.88	5.54	6.82	6.95	7.28
	20:30	22.0	C	4.30	4.58	4.60	4.56	4.56
	24:00	20.5	C	4.16	4.20	4.27	4.31	4.31
5-13-69	6:00	18.5	C	3.96	3.96	4.04	3.96	3.91
	8:30	--	R	--	--	--	--	--
	10:00	19.5	O	4.10	4.16	4.23	4.34	4.30
	13:00	26.0	O	4.40	4.72	5.15	3.30	5.35
	16:00	29.9	C	4.92	5.44	6.70	6.85	7.02
	20:30	24.1	C	4.50	4.58	4.70	4.66	4.66
	24:00	20.4	O	4.12	4.16	4.29	4.26	4.26

EMITTANCE MEASUREMENTS OF CHANGING WATER TABLE
FOR SAND WITH NON-SALINE WATER

Date	Time	Ambient Temp.	Met. Cond.	E M I T T E D E N E R G Y				
				Water Level				
				0 ft.	1 ft.	2 ft.	3 ft.	4 ft.
5-19-69	6:00	12.6	C	3.64	3.36	3.36	3.36	3.36
	8:30	18.0	C	3.84	3.50	3.50	3.50	3.50
	10:00	19.9	C	4.16	4.60	4.51	4.51	4.58
	13:00	27.7	C	5.00	6.00	6.12	6.20	6.24
	16:00	29.7	C	4.84	6.00	6.08	6.28	6.52
	20:30	24.1	C	4.50	4.54	4.58	4.58	4.58
	24:00	20.2	C	4.30	4.16	4.20	4.20	4.20
5-20-69	6:20	18.3	C	4.10	3.84	3.88	3.88	3.88
	10:00	26.7	C	4.40	5.22	5.22	5.30	5.30
	13:00	28.0	C	4.95	5.92	6.08	6.12	6.36
	16:00	29.0	PC	4.75	5.53	5.70	5.70	5.78
	20:30	25.2	C	4.50	4.58	4.63	4.60	4.65
	24:00	22.5	C	4.32	4.44	4.50	4.46	4.50
5-21-69	6:00	20.0	C	4.16	3.97	3.94	3.99	4.00
	8:30	23.0	C	4.23	4.16	4.12	4.09	4.11
	10:00	27.0	C	4.40	5.15	5.18	5.22	5.31
	13:00	32.6	C	5.04	6.66	6.82	6.82	6.90
	16:00	32.6	C	5.11	6.61	6.70	6.82	6.90
	20:30	26.2	C	4.58	4.71	4.80	4.80	4.78
	24:00	20.4	C	4.41	4.28	4.27	4.27	4.23
5-22-69	6:00	17.0	PC	4.12	4.02	4.04	4.00	4.01
	8:30	15.0	O	3.97	3.93	3.94	3.90	3.90
	10:00	16.2	O	3.97	3.97	3.97	3.93	3.94
	13:00	20.2	O	4.27	4.70	4.65	4.69	4.70
	16:00	18.0	PC	4.32	4.54	4.42	4.32	4.32
	20:30	15.6	C	4.09	4.16	4.09	3.96	3.96
	24:00	13.0	C	4.05	4.16	4.01	4.01	4.01
5-23-69	6:30	13.0	C	4.09	4.28	4.01	4.01	4.01
	8:30	14.0	C	4.16	4.23	4.09	4.09	4.09
	10:00	18.8	C	4.36	4.64	4.64	4.62	4.62
	13:00	19.0	C	4.64	5.19	5.15	5.15	5.15
	16:00	22.8	C	5.13	5.46	5.60	5.46	5.46
	20:30	18.8	C	4.71	4.77	4.78	4.85	4.77
	24:00	18.0	C	4.36	4.36	4.44	4.32	4.32
5-24-69	6:30	15.8	C	4.12	4.31	4.40	4.16	4.16
	8:30	19.4	C	5.03	5.26	5.15	5.15	5.15
	10:00	20.8	C	5.07	5.52	5.52	5.60	5.60
	13:00	29.4	C	6.23	6.66	6.66	7.08	7.24
	16:00	25.0	C	5.66	5.85	6.00	6.15	6.23
	20:30	19.2	C	4.81	5.00	5.04	5.00	5.00
	24:00	18.4	C	4.28	4.50	4.58	4.50	4.50

APPENDIX D

EMITTANCE MEASUREMENTS OF CHANGING WATER TABLE
FOR SAND WITH SALINE WATER

Water Level	Date	Time	Ambient Temp.	Met. Cond.	E M I T T E D E N E R G Y				
					Salt Content in ppm				
					160,000	100,000	60,000	20,000	0
4 ft.	5-30-69	6:00	20.6	PC	4.00	3.96	4.03	3.96	4.02
		8:30	24.6	PC	4.44	4.44	4.47	4.49	4.54
		10:00	27.6	C	4.84	5.22	5.32	5.46	5.56
		13:00	33.4	C	6.74	6.96	6.90	7.08	7.08
		16:00	35.4	C	6.90	7.05	7.08	7.18	7.18
		20:30	28.4	C	4.96	5.01	5.00	5.00	5.13
		24:00	24.0	C	4.40	4.36	4.39	4.37	4.39
4 ft.	5-31-69	6:00	21.0	PC	4.08	4.06	4.07	4.06	4.06
		8:30	26.0	PC	4.54	4.55	4.53	4.54	4.54
		10:00	25.0	O	4.54	4.54	4.54	4.58	4.55
		13:00	32.6	C	7.05	6.96	6.90	7.00	6.88
		16:00	34.8	C	7.87	8.00	7.96	7.96	8.00
		20:30	28.0	C	4.78	4.74	4.75	4.78	4.77
		24:00	26.5	C	4.58	4.56	4.58	4.58	4.58
3 ft.	6-1-69	6:00	15.0	PC	3.80	3.80	3.90	3.94	3.90
		8:30	16.0	C	3.84	3.94	3.98	4.10	4.10
		10:00	17.5	C	3.96	4.09	4.10	4.30	4.30
		13:00	23.5	C	4.78	4.92	5.08	5.22	5.22
		16:00	26.0	C	5.00	5.38	5.46	5.60	5.64
		20:30	24.0	C	4.10	4.15	4.15	4.14	4.14
		24:00	16.0	C	3.90	3.94	3.90	3.94	3.94
3 ft.	6-2-69	6:00	15.0	C	3.76	3.80	3.80	3.86	3.90
		8:30	15.5	C	3.76	3.84	3.84	3.90	3.90
		10:00	17.5	C	4.10	4.36	4.50	4.58	4.58
		13:00	28.0	C	6.04	6.40	6.48	6.56	6.52
		16:00	26.0	C	5.96	6.00	5.98	5.98	5.92
		20:30	20.5	C	4.36	4.36	4.44	4.48	4.46
		24:00	17.0	C	3.90	3.90	3.95	3.90	3.98
2 ft.	6-3-69	6:00	14.0	O	3.48	3.46	3.44	3.44	3.48
		8:30	17.5	O	3.96	3.96	3.94	3.96	3.93
		10:00	23.0	O	4.32	4.40	4.40	4.40	4.27
		13:00	26.0	PC	5.01	5.12	5.18	5.22	5.22
		16:00	27.0	C	5.46	5.47	5.49	5.52	5.48
		20:30	19.8	C	4.18	4.19	4.22	4.20	4.20
		24:00	16.5	C	3.92	3.96	3.96	3.98	4.00
2 ft.	6-4-69	6:00	15.0	C	3.78	3.84	3.84	3.90	3.90
		8:30	16.0	C	3.88	3.88	3.88	3.93	3.93
		10:00	19.0	C	4.51	4.57	4.64	4.71	4.71
		13:00	27.0	C	5.60	5.70	5.70	5.74	5.79
		16:00	31.0	C	5.92	6.00	6.00	6.00	6.00
		20:30	23.0	C	4.51	4.57	4.57	4.48	4.57
		24:00	17.0	C	4.10	4.03	4.03	4.03	4.10

Water Level	Date	Time	Ambient Temp.	Met. Cond.	E M I T T E D E N E R G Y				
					Salt Content in ppm				
					160,000	100,000	60,000	20,000	0
2 ft.	6-5-69	6:00	16.0	C	3.78	3.68	3.78	3.74	3.78
		8:30	20.0	C	3.93	3.92	3.94	3.96	3.96
		10:00	25.0	C	5.03	5.07	5.03	5.00	4.98
		13:00	33.4	C	6.82	6.82	6.70	6.74	6.70
		16:00	34.0	C	6.78	6.82	7.04	7.14	7.14
		20:30	27.0	C	4.81	4.92	4.84	4.73	4.77
		24:00	22.5	C	4.26	4.26	4.30	4.20	4.22
1 ft.	6-6-69	6:00	19.0	C	4.14	4.09	4.14	4.15	4.18
		8:30	22.5	C	4.22	4.22	4.26	4.25	4.26
		10:00	26.0	C	4.85	4.92	4.92	5.07	5.07
		13:00	35.5	C	6.96	7.00	7.03	7.03	7.07
		16:00	34.5	C	6.65	6.65	6.70	6.74	6.79
		20:30	26.0	C	4.68	4.71	4.76	4.80	4.81
		24:00	22.8	C	4.29	4.29	4.24	4.24	4.26
1 ft.	6-7-69	6:00	17.0	C	3.94	3.96	4.00	4.00	4.00
		8:30	23.0	C	4.10	4.02	4.07	4.10	4.14
		10:00	29.0	C	4.96	4.92	4.96	5.00	5.04
		13:00	35.5	C	6.86	6.82	6.86	6.82	6.70
		16:00	36.0	C	6.70	6.66	6.63	6.56	6.39
		20:30	29.0	C	4.88	4.84	4.83	4.80	4.76
		24:00	24.0	C	4.75	4.72	4.72	4.70	4.67
0 ft.	6-8-69	6:00	20.6	C	4.10	4.10	4.10	4.14	4.14
		10:00	27.5	C	4.84	4.82	4.68	4.68	4.58
		13:00	34.0	C	5.12	5.08	5.08	5.04	5.08
		16:00	29.0	PC	4.82	4.82	4.78	4.70	4.70
		20:30	20.8	PC	4.92	4.84	4.86	4.70	4.72
		24:00	19.5	O	5.04	5.00	5.00	4.96	4.70
0 ft.	6-9-69	6:00	18.0	O	4.30	4.30	4.26	4.26	4.10
		8:30	18.5	O	4.20	4.16	4.16	4.14	3.96
		10:00	20.0	PC	4.06	4.02	4.02	4.00	3.92
		13:00	20.5	PC	4.06	4.02	4.02	4.02	4.00
		16:00	24.0	C	4.20	4.16	4.16	4.14	4.00
		20:30	21.0	C	4.02	4.02	4.00	4.00	3.93
		24:00	19.5	C	4.00	3.96	3.96	3.96	3.88
		6:00	19.0	O	3.95	3.92	3.92	3.91	3.91

EMITTANCE MEASUREMENTS OF CHANGING WATER TABLE
FOR SILTY CLAY WITH SALINE WATER

Water Level	Date	Time	Ambient Temp.	Met. Cond.	E M I T T E D E N E R G Y				
					Salt Content in ppm				
					160,000	100,000	60,000	20,000	Q
4 ft.	6-13-69	6:00	18.2	C	3.98	3.96	3.96	4.00	4.00
		8:30	19.8	C	4.00	3.96	3.96	3.96	4.00
		10:00	22.5	C	4.78	4.89	4.93	5.12	5.12
		13:00	30.0	C	6.66	6.82	6.61	6.75	6.64
		16:00	28.0	C	5.46	5.56	5.56	5.60	5.50
		20:30	--	R	--	--	--	--	--
		24:00	19.5	PC	3.93	3.96	3.96	3.93	3.96
4 ft.	6-14-69	6:00	17.0	PC	3.80	3.80	3.82	3.88	3.88
		8:30	20.0	PC	4.12	4.16	4.16	4.22	4.28
		10:00	--	R	--	--	--	--	--
		13:00	--	R	--	--	--	--	--
		16:00	22.5	PC	4.26	4.26	4.30	4.40	4.40
		20:30	19.5	C	4.04	4.12	4.08	4.16	4.16
		24:00	16.5	C	3.84	3.88	3.92	3.96	3.96
3 ft.	6-15-69	6:00	14.0	C	3.52	3.48	3.48	3.46	3.44
		8:30	19.1	C	3.66	3.66	3.66	3.64	3.64
		10:00	26.0	C	4.74	4.74	4.74	4.70	4.68
		13:00	29.0	C	7.04	7.36	7.24	7.24	7.14
		16:00	29.5	C	7.76	8.00	7.76	7.96	8.00
		20:30	23.5	C	4.56	4.51	4.48	4.48	4.46
		24:00	21.5	C	4.26	4.26	4.26	4.22	4.21
3 ft.	6-16-69	6:00	17.5	C	3.90	3.90	3.90	3.90	3.86
		8:30	22.5	C	4.10	4.06	4.06	4.06	4.06
		10:00	26.0	C	5.12	5.12	5.08	5.08	5.00
		13:00	34.0	C	7.00	7.00	6.90	6.86	6.86
		16:00	36.5	C	7.36	7.36	7.32	7.28	7.28
		20:30	29.3	C	4.92	4.88	4.88	4.84	4.84
		24:00	27.0	C	4.64	4.60	4.60	4.56	4.56
2 ft.	6-17-69	6:00	22.0	C	4.22	4.22	4.19	4.19	4.17
		8:30	25.5	C	4.40	4.36	4.36	4.32	4.32
		10:00	28.5	C	5.42	5.38	5.38	5.34	5.30
		13:00	31.8	PC	5.92	5.92	5.88	5.84	5.80
		16:00	34.0	C	7.50	7.48	7.44	7.36	7.28
		20:30	27.0	O	4.83	4.74	4.74	4.71	4.73
		24:00	24.0	C	4.36	4.34	4.34	4.36	4.34
2 ft.	6-18-69	6:00	18.0	C	3.96	3.96	3.92	3.90	3.86
		8:30	26.5	C	4.60	4.60	4.56	4.54	4.54
		10:00	29.8	C	5.12	5.08	5.08	5.00	5.00
		13:00	35.8	PC	7.28	7.14	7.14	7.10	7.00
		16:00	36.0	C	7.36	7.28	7.28	7.18	7.10
		20:30	31.0	C	5.18	5.14	5.12	5.08	5.08
		24:00	30.0	C	4.78	4.77	4.74	4.68	4.68

Water Level	Date	Time	Ambient Temp.	Met. Cond.	E M I T T E D E N E R G Y				
					Salt Content in ppm				
					160,000	100,000	60,000	20,000	0
1 ft.	6-19-69	6:00	22.5	C	4.44	4.40	4.40	4.36	4.36
		8:30	26.0	C	4.40	4.35	4.35	4.36	4.36
		10:00	28.9	PC	4.50	4.48	4.44	4.44	4.40
		13:00	36.5	PC	5.22	5.22	5.14	5.12	5.00
		16:00	35.8	PC	5.30	5.14	5.14	5.12	5.06
		20:30	--	R	--	--	--	--	--
		24:00	20.0	PC	4.10	4.04	4.04	4.00	3.93
1 ft.	6-20-69	6:00	21.0	O	4.20	4.20	4.20	4.20	4.16
		8:30	20.5	O	4.14	4.14	4.14	4.10	4.10
		10:00	23.0	O	4.30	4.30	4.30	4.30	4.27
		13:00	29.0	O	4.66	4.66	4.68	4.68	4.68
		16:00	34.8	PC	5.38	5.38	5.34	5.30	5.26
		20:30	29.0	C	4.74	4.74	4.68	4.64	4.64
		24:00	27.5	C	4.62	4.62	4.60	4.56	4.54
0 ft.	6-21-69	6:00	27.0	O	4.36	4.33	4.30	4.30	4.26
		8:30	28.5	C	4.44	4.40	4.40	4.36	4.36
		10:00	27.5	O	4.40	4.44	4.39	4.34	4.30
		13:00	31.5	PC	5.18	5.12	5.04	5.00	4.96
		16:00	36.0	PC	5.18	5.14	5.04	5.00	5.00
		20:30	29.0	C	4.68	4.68	4.64	4.64	4.64
		24:00	25.5	C	4.30	4.26	4.26	4.20	4.20
0 ft.	6-22-69	6:00	25.5	C	4.34	4.34	4.34	4.34	4.30
		8:30	26.0	C	4.34	4.30	4.26	4.22	4.20
		10:00	28.0	C	4.44	4.40	4.36	4.36	4.34
		13:00	34.5	C	5.26	5.08	4.96	4.96	4.84
		17:00	35.5	C	5.42	5.22	5.00	4.92	4.78
		20:30	30.0	C	5.52	5.38	5.22	5.08	4.92
		24:00	29.5	C	5.52	5.38	5.30	5.22	5.16

VITA⁴⁷

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